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**MAP OF UPPER DEVONIAN AND YOUNGER
PALEOZOIC ROCKS, MACKENZIE CORRIDOR**

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MAP OF UPPER DEVONIAN AND YOUNGER PALEOZOIC ROCKS, MACKENZIE CORRIDOR

This map summarizes data on Upper Devonian, Carboniferous and Permian rocks for the Mackenzie Corridor - an area from latitude 60°N to the Arctic coast (but excluding the Mackenzie Delta), from the exposed Precambrian on the east to longitude 136° on the west. Rock types, approximate thickness, and correlation are shown on the inset chart. Sources for field data are shown on the inset index map. Subsurface interpretations are based on my own work south of 65°N , on company well reports and Canadian Stratigraphic Service Ltd. lithologs north of 65°N . Regional summaries of parts of the Corridor area, from a petroleum assessment point of view, are in de Wit et al. (1973), Gilbert (1973), Kunst (1973) and Law (1971). Concepts of depositional history have been discussed and documented by Braman (1981) for the northern area, and by Richards (1983) for Lower Carboniferous strata in the south.

In order to avoid an impossibly cluttered map, the emphasis varies from area to area. In the south, the emphasis is on sub-Cretaceous geology and on facies trends. Isopach maps of various units in this area appear in Belyea (1971) and Williams (1977, 1981). In the Mackenzie Mountains, where most of the pre-Permian strata are of the starved basin facies, the map merely shows distribution of preserved rocks. In the north, the area of highest petroleum potential, drill stem tests and isopachs are shown.

South of the 60th parallel, correlative rocks host large reserves of oil and gas. For example: Upper Devonian Leduc reefs, Nisku and Wabamun carbonates, Mississippian carbonates, and Mississippian and Permian sandstones. This map should help to explain why upper Paleozoic rocks of the Mackenzie Corridor, in contrast to those south of latitude 60°N , offer such a low potential for petroleum. In the Eagle Plains, west of the map area, several small gas fields and one oil field occur in Viséan (Mississippian) clastics.

Some of the more interesting items depicted by the map are listed below.

Southern part of the map

- 1) Most of the mapped sediments belong to a regressive suite – the waning stages of the Kaskaskia Sequence of Sloss (1963). This is most obvious from column GH of the inset chart (essentially a west to east cross-section). The various shallow water carbonate units step progressively farther west through time: Twin Falls, Jean Marie, Kakisa, uppermost Kotcho limestone tongue and, lastly, the shallow water facies of the Flett Formation.
- 2) All mappable facies belts trend north by northwest. This pattern – westward prograding carbonate units over westward dipping clinothem within the clastic facies – is consistent with earlier sedimentation patterns in the Hay River/Twin Falls formations (Williams, 1977) and during the Givetian (Williams, 1986). The sedimentary model assumes a distant source area for the clastic material, delivered, via the Franklinian Trough (Embry and Klovan, 1976), by long shore or long bank currents. This pattern persisted through Devonian and most, if not all, of Carboniferous time.
- 3) There was a pronounced tectonic sag west of the Liard River; this is best shown on the inset cross-sections. The clastic fill of this sag is a shallowing-upward, coarsening-upward sequence: deep water Besa River shales at the base, with tongues of deep water cherty limestone, siltstone or sandstone, culminating in the shallow water, in part nonmarine Mattson sandstones (see Richards, 1983). Most of the Mattson sands reached this area from the northeast or north; the inferred western source for some of the sandstone (lower inset cross-section) is deduced from the outcrop pattern of lowest Mattson sandstone tongues on the La Biche map (Douglas and Norris, 1959), but this requires further checking in

the field. Although no faults are drawn on the inset cross-sections, the sag could be interpreted as a north-trending graben, or half-graben.

- 4) The pattern of sub-Cretaceous geology, assuming an erosion surface of low relief, indicates a broad, east-west trending syncline. This feature approximately coincides with what was, in Devonian time, the Tathlina Arch (Belyea, 1971). This is an interesting parallel with the post-Devonian reversal of the Peace River Arch.
- 5) Areas where the upper Paleozoic section contains a suitable facies at sufficient depths to offer petroleum potential are fairly small. A Desan-type play (Debolt equivalents lie within the Flett carbonates) exists between longitudes 121°W and 123°W at shallow depths (log section 3). The Jean Marie play (log section 2) occupies part of the same belt. The thick Mattson sandstones (log section 4) are exposed in the large Laramide structures west of 123°W ; prospective traps would be limited to structural depressions. The up-dip eroded edge of the Mattson sandstones (a Tattoo-type play) offers some potential; over most of this belt the Paleozoic sands are directly overlain by a porous basal Cretaceous sandstone (Williams, 1978).

The Mackenzie Plain (between the Mackenzie and Franklin Mountains, north to $\sim 64^{\circ}\text{N}$)

- 1) Although thickness data are too sparse to construct a useful isopach map, it would appear that this belt is a northern continuation of the tectonic sag so evident farther south. Isopach maps of several older Paleozoic units show that the Mackenzie Plain has had a long history as a negative tectonic belt (reports in preparation). West of this downwarp, in the Selwyn Basin, the Earn Group (Gordey et al., 1982) is a deep water, starved, shaly facies. The chert conglomerates within this package probably indicate local uplifts, cannibalism of earlier Paleozoic strata and, possibly, some western highlands.

- 2) A late Frasnian Jean Marie-like limestone is present within the eastern part of the Mackenzie Plain; whether this limestone is a precise correlative of the Jean Marie or of the Kakisa limestone is unknown. The western depositional limit of this limestone follows the same north by northwest trend as do facies belts farther south. The outcrops were described as reef limestones by Douglas and Norris (1961, 1963); they are biostromal, not biohermal, in geometry.
- 3) There are only two conceivable potential reservoirs: the reefs or the sandstones that occur high in the section; neither display reservoir quality in outcrops.
- 4) In the plains east of the Franklin Mountains, the basal part of the Fort Simpson shale is all that survived pre-Cretaceous erosion. This section offers no significant petroleum potential.

The northern area

- 1) This area, sometimes called the Peel Basin or Peel Plateau, lies within the right-angled reentrant formed by the junction of the Franklinian and Cordilleran geosynclines. The Franklinian Geosyncline ended with the Ellesmerian Orogeny (Late Devonian - Early Carboniferous). The main effects of this Orogeny lie northwest and northeast of the map area, from western Alaska to the Arctic Islands (Douglas, 1970). Within the map area its effects are seen in the pre-Permian lacuna (column A of the inset chart) and in the lobes of coarse clastics, up to pebble sized, in latest Devonian - early Carboniferous fill deposits.

The Cordilleran Geosyncline ended with the Columbian (Late Jurassic to mid-Cretaceous) and Laramide (Late Cretaceous and Tertiary) orogenies. These orogenies influenced the preservation, but not the sedimentary history

of upper Paleozoic rocks. The Eskimo Lakes Arch (Young et al., 1976) is a Mesozoic feature; this area was probably strongly negative (an Ellesmerian foredeep) through Devonian and earliest Carboniferous time.

- 2) The small area with Carboniferous limestone, represented by column D of the inset chart, is representative of a stratigraphic section that occurs in the Eagle Plain, west of the map area. Rocks of this facies do not occur, and probably never were deposited, east of the Richardson Mountains.
- 3) The isopached unit, which includes the Canol Formation, is a shallowing-upward, coarsening-upward sequence. The deep water, starved-basin Canol shales are succeeded by Imperial shale, silty shale, sandy shale with sandstone turbidites, and, finally, by shallow water marine to nonmarine sandstone. From the study by Braman (1981) it is evident that most of the basin fill prograded into the area from the east or northeast. The youngest part of the fill also received a contribution from rising Ellesmerian highlands northwest of the map area. One of the main depocentres was in the area of the northern Richardson Mountains (see Fig. 16 of Norris, 1985).
- 4) The position of the Cretaceous/Paleozoic boundary is uncertain in most wells, highly uncertain in many [the type section of the Tuttle Formation as defined in the subsurface, includes an unknown thickness of Cretaceous section (Pugh, 1983, p. 44)]. In most areas there is no clearly recognizable basal Cretaceous sandstone. Even with palynological control, the boundary is not always clear (see discussion by McGregor in Norford et al., 1970, p. 8).
On the Trail River map the Cretaceous/Paleozoic contact is mapped as a near vertical fault - the Trevor Fault (Norris, 1981d). This cannot be correct. Shell Peel River L-19 (~66°45'N, 135°15'W) and I-21 (~66°10'N, 134°15'W), both of which lie about 10 km east of the alleged fault, encountered Paleozoic rocks at

or very near the surface (see well reports, also report by McGregor in Norford et al., 1970, p. 4-7). The Trevor Fault may exist, as mapped; it is not, however, along the Mesozoic/Paleozoic contact.

- 5) In the southeastern part of the Wind River map (column C of the inset chart) there is a strong unconformity between Devonian and Mississippian beds. This is probably a local phenomenon, unrelated to the Ellesmerian Orogeny. It probably is, rather, a manifestation of a paleoarch that later became the Bonnet Plume Basin (a report on Cretaceous history, in preparation).
- 6) A search for petroleum in upper Paleozoic rocks of this area must concentrate on finding reservoir sands within the flysch or basinal facies. The sandstones of the eastern, or platform facies are, for the most part, too shallow to offer favorable prospects. Also, along the eastern belt of preserved Imperial Formation the basal Cretaceous sandstone is porous. The fact that this basal Cretaceous sandstone is commonly oil stained or oil saturated (Yorath and Cook, 1981, p. 45) probably indicates that oil has migrated through and escaped from Imperial sandstones. Oil impregnated Imperial sandstones are known in the Norman Wells area (Tassonyi, 1969, p. 147).

The petroleum potential of the flysch facies is difficult to assess. Imperial and Tuttle sandstone units are characteristically poorly sorted, lithic and clay rich. However, the entire section contains a high proportion of organic matter (plant remains) and reservoir beds do exist, albeit rarely. Small gas flows were obtained in the Russell H-23 well (~70°N, 130°W) from two Imperial sands. Between 1965 and 1967 Shell Canada drilled nine shallow tests in the Peel River area (~66°N-67°N, 134°W). Of these, four encountered thin reservoir sands with fair permeabilities (judging from water recoveries), three of the four reservoirs yielded traces of gas.

- 7) There is as yet no published comprehensive analysis of the thermal maturity of the Imperial Formation. However, a large volume of these sediments must lie within the oil window. From the fact that there is oil at Norman Wells, it can be assumed that the Canol Formation in that general area lies within the oil window. In the well report on Shell Peel River H-37 (~66°35'N, 135°W) it is stated that "...by 8595' the sediments are mature...and by 9000' expulsion should be complete". The base of the Imperial Formation in the well is at 8860' (2700 m). Norris and Cameron (1986) describe an outcrop of bitumen that occurs in the upper part of the Imperial Formation (~67°50'N, 133°45'W). Both the bitumen and host sediments indicate a maturity within the lower part of the oil window. The base of the Imperial in this area lies at a depth of about 1500 m. In the Kugalik N-02 well (~68°30'N, 131°30'W) the Canol shale (881 m) has been heated well beyond the oil window (Gunther and Meijer Drees, 1977).

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List of Illustrations

Map of upper Paleozoic sediments, Mackenzie Corridor, scale ~1:1,270,000.

Log section 1, Upper Devonian formations, near 60°, 119°.

Log section 2, Upper Devonian formations, near 60°, 122°.

Log section 3, Flett Formation, near 60°, 123°.

Log section 4, Flett to Mattson transition, near 60°, 124°.