

LEGEND

Coloured legend blocks indicate map units that appear on this map

QUATERNARY

SURFICIAL DEPOSITS

POST LAST GLACIATION

NONGLACIAL ENVIRONMENTS

ORGANIC DEPOSITS: peat; 1 to 2 m thick; formed by the accumulation of vegetation in poorly drained depressions (swamps and bogs); produces flat, wet terrain

Bog peat: Sphagnum or forest peat formed in an ombrotrophic environment; may be treed or treeless with a cover of ericaceous shrubs; hummocky, wet terrain, in places underlain by ground ice, O¹h; undifferentiated bog and fen deposits, O

Fen peat: peat derived from sedges and partially decayed shrubs in a eutrophic environment; forms relatively open peatlands with a mineral rich water table that persists seasonally near the surface; often covered with low shrubs and sometimes a sparse layer of trees

COLLUVIAL DEPOSITS: mass wasting debris <1-100 m thick; nonsorted to

poorly sorted, massive to stratified debris deposited by direct, gravity-induced movement

Landslide debris: active and inactive landslides, undivided, C; hummocky

Talus (scree): accumulation of angular boulders below cliffs; usually forming fans or

C_{RG}
Rock glaciers: inactive rock glaciers; rock debris deformed by the flow of buried or interstitial ice, forming ridges and furrows

Colluvial veneer: thin cover of rock debris <1 m; usually the product of solifluction acting on lower slopes

ALLUVIAL DEPOSITS: sorted gravel, sand, and organic detritus deposited by

Fluvial deposits: sorted gravel and sand >1 m thick; forming active flood plains with meander channels and scroll marks, Ap; alluvial fan deposits, poorly sorted gravel and sand >1 m thick, Af; numerous subparallel alluvial channels covering gentle to moderate slopes, Ac; low terraces with meander scars with active and inactive channels, primarily along the Liard River, AI; undifferentiated, A

Alluvial veneer: deposits too thin to mask the underlying surface, <1 m thick

Fluvial terraces: inactive; alluvial fan terraces, Aft

LACUSTRINE DEPOSITS: sand, silt and minor clay deposited in a former lake; generally overlain by organic deposits; exposed by recent fluctuations in lake levels

POSTGLACIAL OR LATE WISCONSINAN
PROGLACIAL AND GLACIAL ENVIRONMENTS

GLACIOLACUSTRINE DEPOSITS: fine sand, silt, and clay, deposited in glacier-dammed lakes in valleys or along the margin of the retreating Laurentide Ice

Glaciolacustrine deposits: sediment > 1 m thick; usually overlain by organic deposits in lowlands with level topography; hummocky topography, Lh

Lv Glaciolacustrine veneer: discontinuous cover of lacustrine sediment, usually <1m

GLACIOFLUVIAL DEPOSITS: proglacial outwash, gravel and sand with minor diamictons deposited in front of the ice margin, usually 1-10 m thick; forming distal outwash terraces Gt; proximal kettled outwash terraces, Gtk; glaciofluvial fan terraces, Gft

Till blanket: > 1 m thick; forming undulating topography that may be fluted and drumlinized in places; contains Canadian Shield erratics; moraines or crevasse

Till veneer: < 1 m thick and discontinuous; underlying bedrock topography is

TILL: nonsorted debris deposited directly by glaciers; matrix is sandy to clayey and contains striated clasts of various lithologies, including many Canadian Shield erratics

fillings forming a ridged topography, Tr; hummocky moraine, Th

PRE-QUATERNARY BEDROCK

R Sedimentary bedrock, R; two prominent ridges of the Liard Range expose Paleozoic to Mesozoic rocks forming northward trending cliffs; extensively gullied outcrops exposed along river banks, Rg

Rw Rubble covered bedrock surfaces

NOTE: In areas where the surficial cover forms a complex pattern, the area is coloured according to the dominant unit and labelled in descending order of cover

Geological boundary (defined, gradational)		/-
Small swamp or bog		∌€
Landslide (arrow indicates the direction of movement)		\bigcirc
Abandoned meltwater channel or channel occupied by an underfit stream (large, small and direction of flow inferred, small and direction of flow not inferred)		1/
Lateral meltwater channel, barb on the uphill side	🗸	~
Escarpment		TITITI
Kettle		The same
End moraine		
Minor moraine or crevasse filling		//
ce moulded form in bedrock and/or till (direction of flow inferred, not inferred)		//
Striae		7
Cirque; cirques; peaks and sharp ridges formed by glacial erosion		
Radiocarbon date (uncorrected)	Date Lab. no.	Material Elev. (m)

NOTES ON GLACIAL HISTORY:

Fisherman Lake map area was glaciated during the Late Wisconsinan glaciation (ca. 25 000-10 000 years ago). Wood fragments between reworked till and glaciolacustrine sediment radiocarbon dated 32.7 ka BP (I-3187, Millar 1968), which provide a maximum age for the advance. Although the area was probably first glaciated by montane ice from the west, the last major ice sheet to cover the area was the continental Laurentide Ice Sheet. The Laurentide Ice Sheet flowed into the region from the northeast, carrying with it distinctive erratics from the Canadian Shield, which are now found in till mantling the valley floors and on several mountain summits in the area. Just to the east of the map area, the flow of this ice sheet was diverted southward along the eastern flank of the Liard Range, but South of Mount Coty, the ice flowed to the west and then northwest into the map area. This northward deflection in flow was caused by the Laurentide Ice Sheet coming in contact with montane ice west of here. North of Fisherman Lake, the flow continued northward, up the mountain slopes. Well-preserved glacial striae on a ridge crest at 1200 m above sea level, show a northwestward ice flow into the Kotaneelee River valley.

During deglaciation, the Liard River system was blocked by the retreating Laurentide Ice Sheet and glacial lakes were created in many of the mountain valleys, including the valley occupied by Fisherman Lake. As the ice margin retreated further eastward these lakes enlarged onto the plains forming glacial Lake Liard. Pond sediments a few kilometres east of Fisherman Lake record continuous sedimentation since at least 9590 BP, with permafrost invading the site during the last 300-500 years (GSC-1890, and other dates; Matthews 1980). Several other radiocarbon dates in the area record episodic forest fires (e.g. GSC-717; McNeely, 1989; exact location unknown), as well as human occupation dating back several thousand years (Millar, 1968).

Landslides are common in the area. Interbedded recessive shales and thick sandstones exposed along cliffs in the Liard Range are prone to failures causing landslides ranging in size from a few 10's m² to several km². Active and relic landslides suggest that mass wasting has been occurring throughout the postglacial time and is on going. For example, an active landslide/debris flow occurred 5 km northeast of Fisherman Lake, on a west-facing mountain flank. Initial failure occurred near the ridge crest prior to August 1991, and it now extends about 3.5 km downvalley but is mostly <200 m wide. A second debris flow occurred in ca. 1995, on the same ridge, 5 km north of the former landslide.

REFERENCES

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