Summary Geological Report for the McConnells Jest Property, Yukon Territory

Collaborative Authors: *Fraser Kirk, Daniela Marcoux, Hanson Wong & Carrie Wong*

Project Overseen By: *Andrew Randell, PGeo*

24th April 2015

Contents

1.0 Summary

The McConnell's Jest property, located in the central Yukon, lies 65 km northeast of Mayo, Y.T. Access to the property is limited to a 25 minute helicopter trip from the Mayo airstrip. The Hansen-McQuesten Lake road, which lies east of the property provides foot and skidoo access.

The McConnell Property consists of 172 contiguous (quartz) claims, owned 100% by Bill Koe-Carson, and covers an area of approximately 3,371 hectares. The property was under option with Golden Predator Canada Corp. but that agreement was terminated due to unrelated legal issues.

The property itself has seen limited exploration activity. The property was first staked by United Keno Hill Mines in the 1960’s, termed the “Zed” (Z) claims (Minfile#: 106D055). The property was subsequently staked in 2010 by Bill Koe-Carson. Golden Predator Canada Corp. oversaw soil and rock sampling programs in 2011 and 2012.

Mapping took place in 1961 (Green, 1972) by L. Green and the Geological Survey of Canada (GSC) as part of a helicopter-supported party known as the Operation Ogilvie (Minfile# 106D055). Directly south of the property, local topographic map sheet 105M was remapped by Roots (1997) of the GSC, and in 2003 the GSC released a geological compilation that included this area (Golden Predator Canada Corp., 2013).

The property underwent glaciation during the McConnell glaciation (>23,000 years ago; Bond, 1999). It has been demonstrated through a number of field seasons that the ground is covered by basal till. The lack of a distinct soil anomaly is attributed to the presence of loess deposits on the property.

The McConnell pluton is one of a series of Cretaceous plutons that have been included in the Tintina Gold Province. The Tintina Gold Province (e.g. Hart, 2004) is a belt of Au-deposits in the Northern Cordillera of Yukon and Alaska that are bounded by two dextral transpressional fault systems - the Denali Fault (southern limit) and the Tintina Fault (northern limit). The belt extends beyond the northern tip of the Tintina Fault, into the Selwyn Basin in the Yukon, and into the western-most NWT. Within the Tintina Gold Province a series of highly prospective Intrusion-Related Gold (IRGS) deposits form the Tombstone Gold Belt, the most significant host to IRGS globally.

Paleozoic clastic rocks of the upper Devonian and Mississippian Earn Group underlie the majority of the McConnell property. These metasedimentary sequences were formed in a submarine fan and channel deposit setting and subsequently deformed during Cordilleran tectonics. The 7 x 2.5 km McConnell pluton intrudes the Earn Group. The pluton is a mid-Cretaceous Tombstone suite granodiorite intrusion which occupies a large portion of the property.

Two major mineral properties lie adjacent to McConnells Jest, the Dublin Gulch IRGS deposit to the west and the Keno Hill silver district to the south east. Many similarities exist between McConnells Jest and Dublin Gulch (6.3 M oz indicated and inferred, Wardrop Engineering Inc., 2011) and so the IRGS model has been adopted to describe mineralisation.

The basis of this report is to describe and draw conclusions from all previous exploration activity on the property. This report will address current understanding and look to provide new insight into the nature of mineralisation.

Additionally, a series of cost-effective strategies in order to better understand and expand mineral resources are proposed. All conclusions at this time should be considered preliminary.

2.0 Hive Introduction

This report has been compiled, researched and written by SGDS Hive, which is a collaborative team of graduates and professionals. The information in this report has been fact checked by Andy Randell P.Geo who also oversaw the project.

Note that this report does not constitute a National Instrument 43-101 as no site visit was conducted. Instead, this is a summary report that has consolidated data and made conclusions and recommendations from the information available through the claim holder, government and other sources. The report does however follow the framework of a NI 43-101 layout to ensure maximum coverage and completion of the information.

3.0 Location & Access

3.1 Accessibility

The property is located in the central Yukon and lies 65 km northeast of Mayo, Y.T. on map sheet 106D03 and 105M14 at 479500m E and 7100000m N in NAD83 Zone 8 (Fig 3.1). The closest sizeable town is Mayo, located on the Stewart River, approximately 65 km to the southwest. Mayo is accessible from Whitehorse via a 460 km all-weather road and is also serviced by the Mayo airport, which is located just to the north of Mayo. Access to the property is limited to a 25 minute helicopter trip from the Mayo airstrip. The Hansen-McQuesten Lake road, which lies east of the property provides foot and skidoo access.

3.2 Infrastructure

Mayo has a population of approximately 450 and offers accommodation, fuel, a nursing station, and earth-moving contractors. The Government of Yukon maintains a 1,400 m gravel airstrip, suitable for charter flights, about 3 km north of Mayo. There are no scheduled air services to Mayo. Local resources in terms of manpower, rental equipment, materials, and supplies are very limited.

A broader range of services is available in Whitehorse, Yukon, located about six hours by road to the south of the property. Whitehorse has a population of 22,815 (National Household Survey, 2011) and has regularly scheduled air service to Vancouver, Edmonton, Calgary, and Fairbanks.

Electrical transmission lines from a hydroelectric facility near Mayo extend to the villages of Elsa and Keno City, about 20 km south of the property.

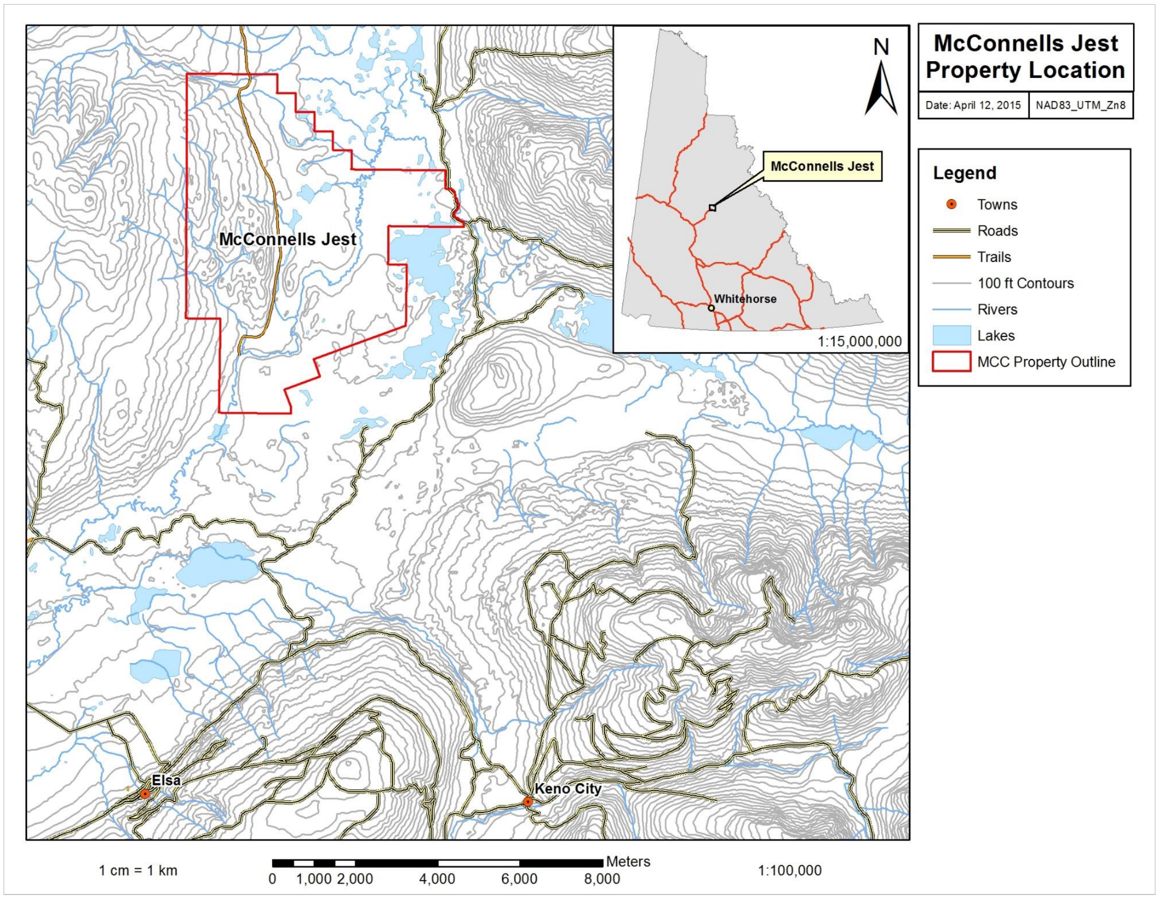

Figure 3.1 *Map showing the boundary of the McConnells Jest claim block. Inset shows position within the Yukon Territory.*

4.0 Claim Information

Claim Name	Number(s)	Grant Number	Registered Owner
McConnells Jest	1 to 40	YD16701 to YD16740	Bill Koe-Carson - 100%
McConnells Jest	41 to 52	YD54701 to YD54712	Bill Koe-Carson - 100%
McConnells Jest	53 to 56	YD54713 to YD54716	Bill Koe-Carson - 100%
McConnells Jest	57 to 120	YD54717 to YD54780	Bill Koe-Carson - 100%
McConnells Jest	121 to 125	YD61470 to YD61474	Bill Koe-Carson - 100%
McConnells Jest	126 - 172	YD126853 - YD126899	Bill Koe-Carson - 100%

Table 4.1 *Claim information for the McConnell's Jest property. Full details can be found in Appendix 6.*

5.0 Physiography and Climate

5.1 Physiography

The property is situated just southwest of the Davidson Range and McQuesten Lake. Topographically, the property lies in the bottom of the McQuesten Valley and is characterized by rolling hills and plateaus; elevation ranges from 640 m to 920 m above sea level. Relief on the property is moderately steep due to creek incising and hills that rise rapidly over lithology changes (Golden Predator Canada Corp., 2011). The property underwent glaciation during the McConnell glaciation (>23, 000 years ago; Bond, 1999), and it has been demonstrated through a number of field seasons that the ground is covered by basal till (Golden Predator Canada Corp., 2013). Outcrops are rare, generally less than two percent of the surface area, and are limited to ridge tops and creek walls. Vegetation on the property consists of stunted spruce on north facing slopes and narrow valley floors, as well as slope alder. South facing slopes contain both coniferous trees and areas of deciduous aspen, poplar and birch. Patchy permafrost occurs on north-facing slopes (Golden Predator Canada Corp., 2011). It should be noted that an estimated 50% of the area has been burned in previous forest fires.

5.2 Climate

The central Yukon is characterized by a subarctic continental climate with cold winters and warm summers. The mean annual temperature for the area is approximately -3°C, with an annual range of 63.5°C. January is the coldest month, July the warmest. Average temperatures in the winter are between -15 and -20 degrees Celsius (°C) but can reach -60°C. The summers are moderately warm with average temperatures in July around 15°C. Annual precipitation ranges from 375 to 600 mm, about half of which falls as snow, which starts to accumulate in October and remains into May or June.

Because of its northern latitude, winter days are short with the sun low on the horizon such that north-facing slopes can experience ten weeks without direct sunlight around the winter solstice. Conversely, summer days are very long, especially in early summer around the summer solstice. Exploration and mining work can be carried out year-round.

6.0 Property History

There are no historical quartz claims recorded in the immediate vicinity of the property. However, based on the Minfile occurrence report for “Zed” (106D 055), United Keno Hill Mines had some ground over the current McConnells Jest claims. Throughout the 1960s and 1970s, United Keno Hill Mines carried out grid soil sampling and prospecting on its claim groups including over the Zed occurrence. No significant mineralization was recorded, and no assessment report was filed.

The area was regionally mapped (1:250,000) by L. Green (1972) of the GSC, and by C. Roots (1997) of the GSC, who remapped topographic map sheet 105 M located to the immediate south. In 2003, Gordey and Makepeace (Gordey and Makepeace, 2003) of the of the GSC released a geological compilation which included the area (MinFile#: 106D 055, 2008).

The ground remained unclaimed until Bill Koe-Carson staked the ground in 2010.

7.0 Geology

7.1 Regional Geology

The McConnell pluton is one of a series of Cretaceous plutons that have been included in the Tintina Gold Province (see Fig. 7.1). The Tintina Gold Province (e.g. Hart, 2004) is a belt of Au-deposits in the Northern Cordillera of Yukon and Alaska that are bounded by two dextral transpressional fault systems - the Denali Fault (southern limit) and the Tintina Fault (northern limit). The belt extends beyond the northern tip of the Tintina Fault, into the Selwyn Basin in the Yukon, and into the western-most NWT. Within the Tintina Gold Province a series of highly prospective Intrusion-Related Gold (IRGS) deposits (Pink belt in Fig. 7.1) form the Tombstone Gold Belt, the most significant host to IRGS globally.

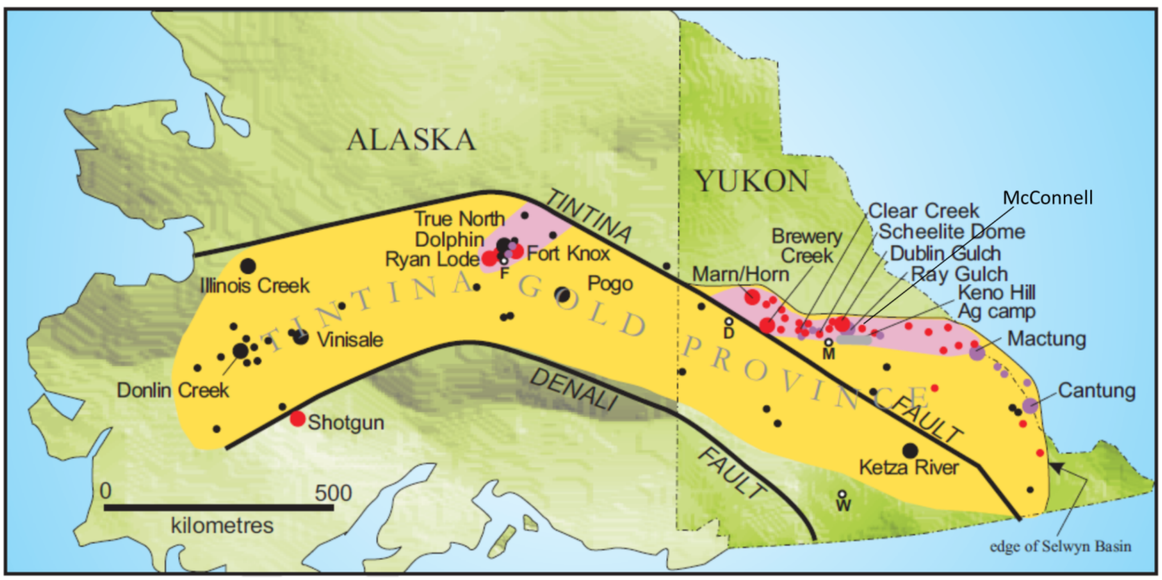

Figure 7.1 *Map of the Tintina Gold Province for the Yukon Territory and Alaska. Mineral deposits are shown as large circles, mineral occurrences are shown as small circles. Gold deposits in red dots could be considered under the same genetic model as McConnells Jest. Deposits marked in black dots are not of the same origin, or are ambiguous in origin. Deposits marked in purple are skarn deposits (predominantly W-bearing). Within the TGP is the Tombstone Gold Belt (TGB; marked in pink) which is the major host to gold deposits in the Yukon and Alaska. Mair et al. (2006) suggested ~450 km of post-formation offset along the Tintina Fault, causing displacement of the Fairbanks district. Population centres marked are Whitehorse (W), Mayo (Mayo), Dawson (D) and Fairbanks (F). After Hart (2007).*

The property lies on the north central margin of the Selwyn Basin tectonic province. The Selwyn Basin, a passive margin sequence, was deposited on the north-western margin of North America during the late

Precambrian through Middle Jurassic (Abbott et al. 1986; Gordey and Anderson, 1993). The McConnell property is underlain by metasedimentary rocks of the Earn Group of the Selwyn Basin (see Fig. 7.2). The Earn Group consists of a series of metasedimentary and meta-volcanic rocks originally deposited during the Devonian to Mississippian. Metasedimentary rocks are commonly grey to black shales, metamorphosed to phyllite, with subordinate chert, siltstone, sandstone, limestone, bedded barite, baritic limestone, and chert-pebble conglomerate. A chlorite-muscovite phyllite unit is proposed to be a metamorphosed felsic volcanic rock (Murphy, 1997). The depositional environment of the Earn Group was a deep marine basin disrupted by faults to cause periods of coarser clastic influx (Abbott et al. 1986).

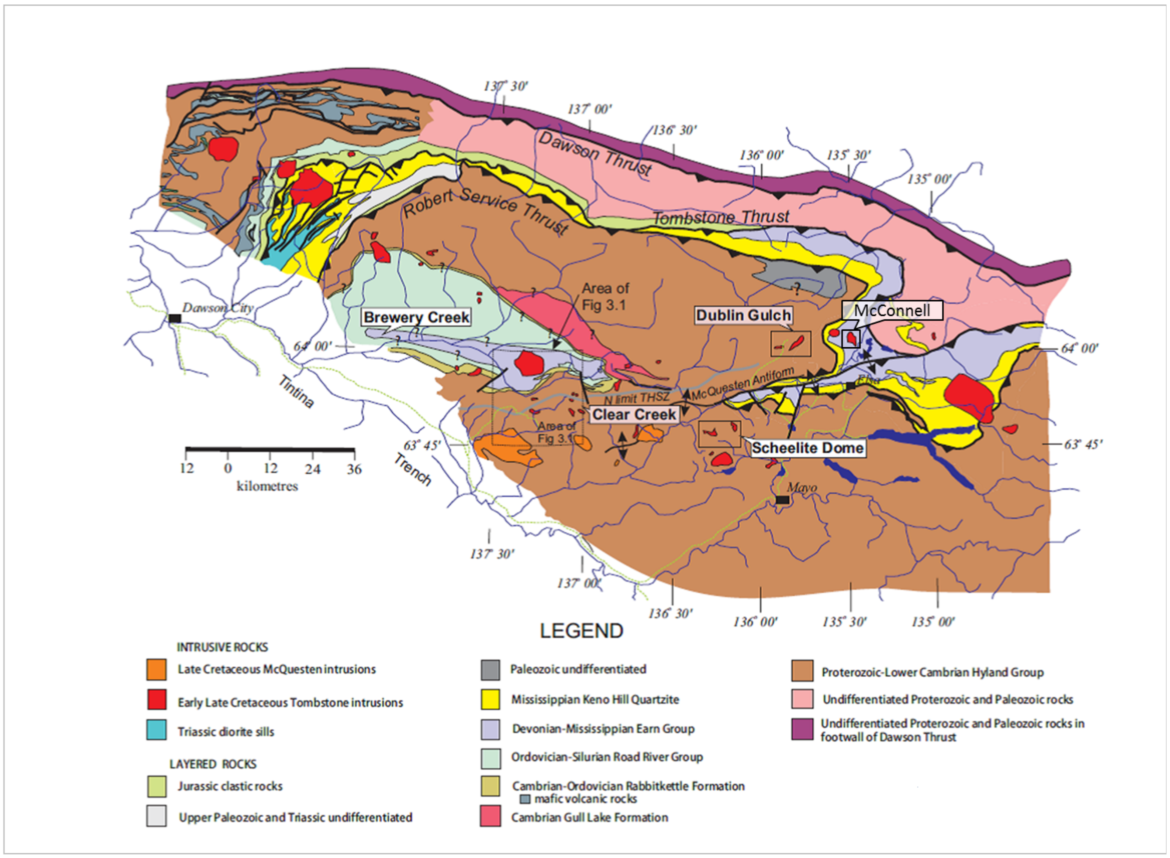

Figure 7.2 *Regional geology of the north western Selwyn Basin showing distribution of plutons, stratigraphic units and structural features. The McConnell pluton is hosted by the Devonian-Mississippian Earn Group. Modified after Scott Wilson Mining (2010).*

Deformation within the Selwyn Basin is associated with the Cordilleran Orogeny. Metamorphism is typically of lower greenschist facies. The formation of a series of folds and three thrust sheets initiated in the Jurassic as the localized effects of Cordilleran convergence began (Mair et al. 2006) The Dawson Thrust, the Tombstone Thrust, and the Robert Service Thrusts disrupt the stratigraphy of the basin (see Fig. 7.3) and may have formed structural conduits for magma during ascent through the crust.

The Tombstone Gold Belt has been divided into a number of suites based on the age, location, morphology and geochemical properties of plutons. In the north-central Selwyn Basin these are the Tombstone Suite (94 Ma – 89 Ma), the Mayo Suite (96 Ma – 93 Ma), and the Tungsten Suite (98 Ma – 94

Ma) (Rasmussen, 2013). Although no date exists for the McConnell pluton, its proximity to Dublin Gulch and the Roop Lakes Stock (94.0 Ma and 92.8 Ma; Selby et al. 2003; Roots, 1997) suggests an age of c. 93 Ma and a classification within the Mayo Suite. The Mayo Suite is characterized by 1-5 km^2 (east)/ 20-80 km^2 (west), single phase to weakly composite plutons, which are alkalic-calcic to calcic and chiefly composed of quartz monzonite, trending east-west along the northern margin of the Selwyn Basin (Hart et al. 2004).

Metamorphic cooling ages (^{40}Ar-^{39}Ar; Mair et al. 2006) indicate that plutonism in the area took place around 10 Ma after the cessation of Cordilleran collisional tectonics. Plutonism took place around 500 km inboard from the active subduction of the Farallon plate beneath North America, indicating that Andean-style subduction related plutonism is not the source of melting. Melting for plutonism took place in the sub continental lithospheric mantle due to mantle upwelling (and associated heat flow) after delamination (Mair et al. 2011).

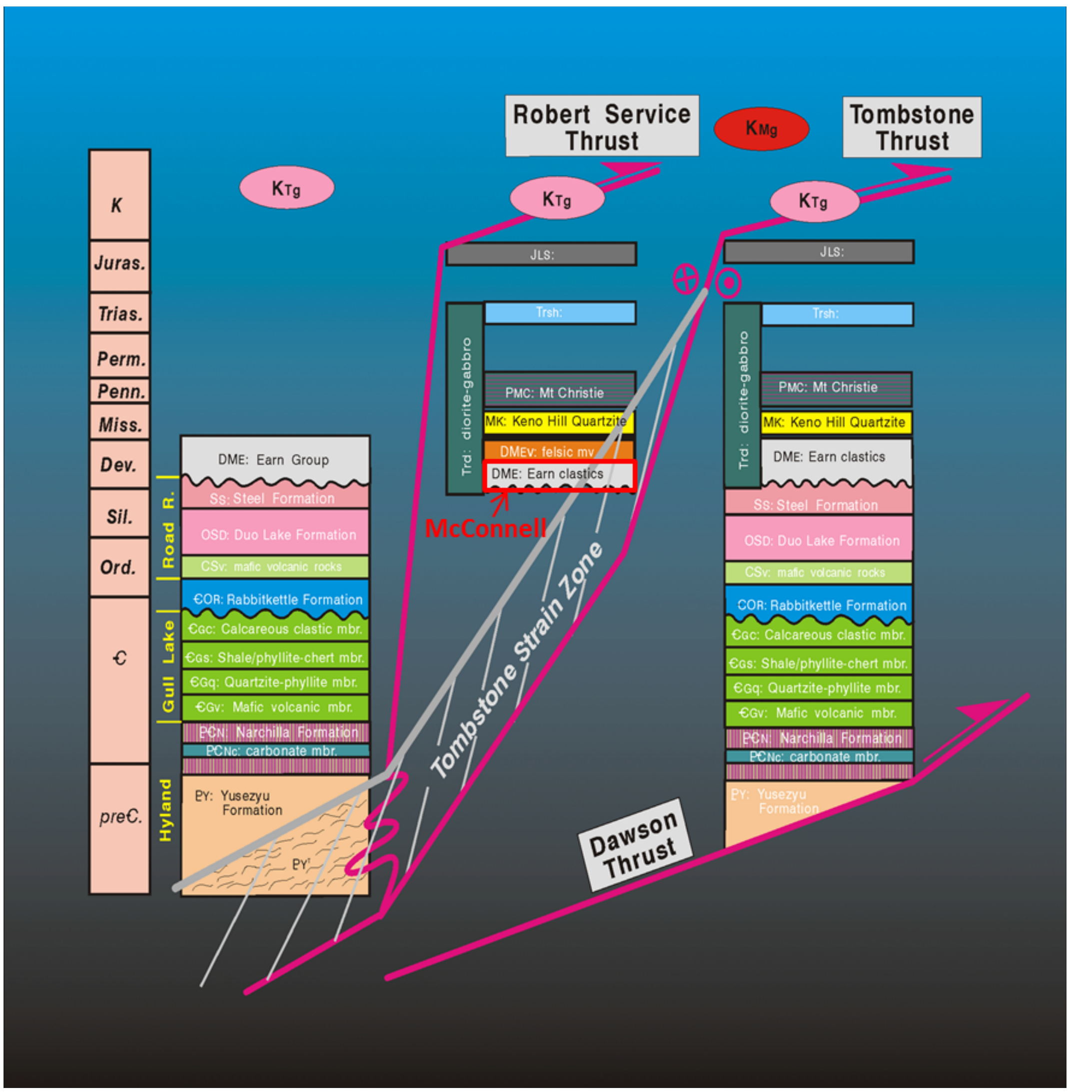

__Figure 7.3__ Stratigraphy of the Selwyn Basin and thrust stacking, modified from Murphy (1997). The McConnell pluton was intruded outside of- but proximal (<10 km) to- the Tombstone Strain Zone.

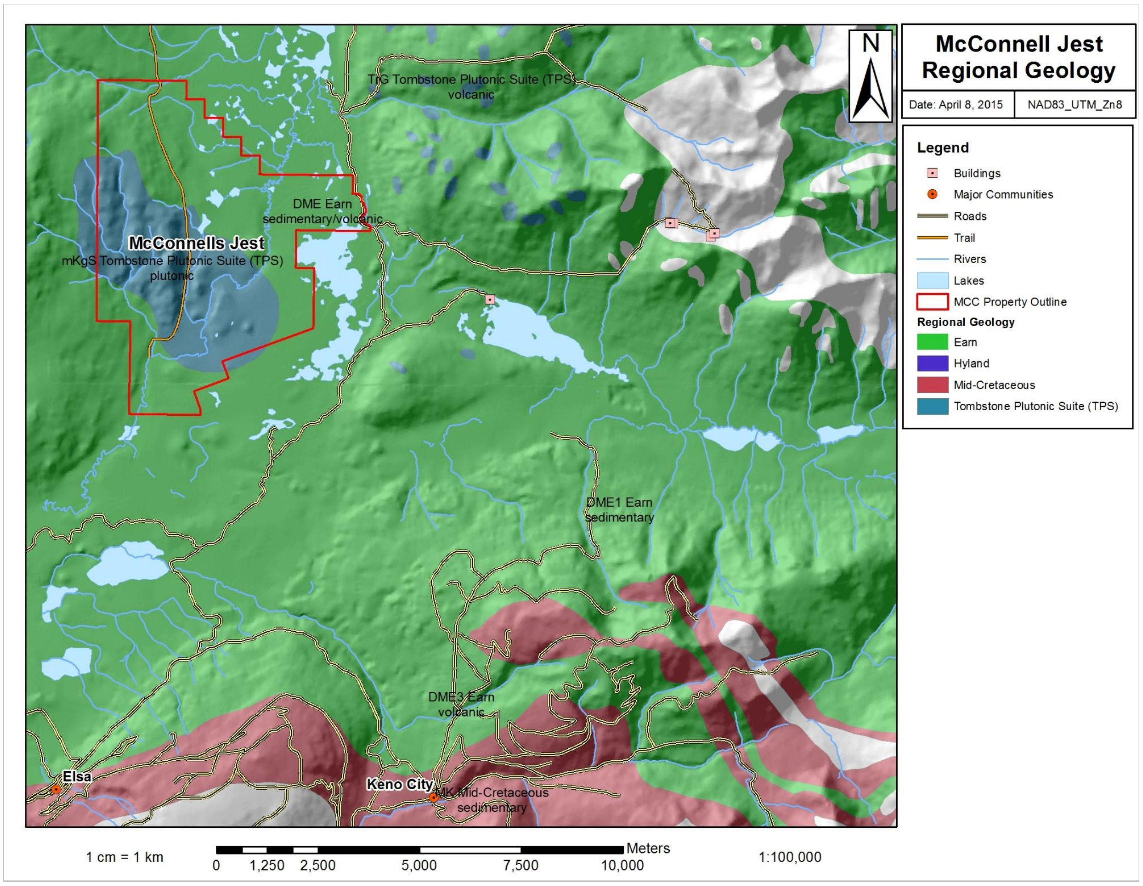

Figure 7.4 *Regional Geology of the McConnells Jest area. A mid-Cretaceous pluton intrudes metasedimentary horizons of the Earn Group.*

7.2 Property Geology

Paleozoic clastic rocks of the upper Devonian and Mississippian Earn Group underlie the majority of the McConnell property (Fig. 7.4). These metasedimentary sequences were formed in a submarine fan and channel deposit setting and subsequently deformed during Cordilleran tectonics. Rock types in the Earn Group are predominantly siliceous shales and cherts with interbeds of arenites and wackes, chert pebble conglomerates, siltstones and barite with rare limestone (Murphy, 1997). A quartz mica schist is the most commonly mapped expression of the Earn Group on the property to date (Koe-Carson, 2010).

The 7 x 2.5 km McConnell pluton intrudes the Earn Group. The pluton is a mid-Cretaceous Tombstone suite granodiorite intrusion which occupies a large portion of the property. The pluton trends ~120 degrees along its long axis. Due to the little amount of historic exploration and drilling, very little is known about contacts, structures and for the most part mineralization.

7.3 Mineralization

Though the property is under-explored, an initial classification within the intrusion-related gold system (IRGS) deposit model is suggested. Hart (2005) provides the following seven points as a summary of defining points for Intrusion-Related Gold Systems:

1. Metaluminous, sub-alkalic intrusion of intermediate to felsic composition, which are transitional between ilmenite and magnetite series.

2. Carbonic hydrothermal fluids responsible for mineralization;

3. A metal assemblage that variably combines gold with elevated Bi, W, As, Mo,Te, and/or Sb and low concentrations of base metals;

4. A low sulphide mineral content, mostly <5 vol%, with a reduced ore mineral assemblage that typically comprises arsenopyrite, pyrrhotite and pyrite and lacks magnetite or hematite;

5. Spatially-restricted, commonly weak hydrothermal alteration;

6. A tectonic setting well inboard of inferred or recognized convergent plate boundaries;

7. A location in magmatic provinces best or formerly known for tungsten and/or tin deposits.

McConnells Jest satisfies the regional geologic requirements of this model and early exploration work suggests that this model is the most appropriate. A partially aligned stockwork of quartz veins, with potassic alteration selvedges, has been identified in multiple locations. Additionally, arsenopyrite veins associated with higher gold grades, with more pervasive clay alteration selvedges, are also present within the current sample set. The proximity and similarity of McConnells Jest to the adjacent Dublin Gulch, 6.3 M oz Au (Wardrop Engineering Inc., 2011), which is a holotypic example of an IRGS, further strengthens the case for this classification. If further exploration clarifies the proposed IRGS classification then gold can be genetically related to the intrusion of the McConnell pluton. Further, the areas of higher potential on the property will be those portions of the pluton which are unroofed. Several examples of elevated tungsten exist in the sample set and suggest that exploration for a skarn deposit would be justified.

7.4 Statistics

7.4.1 Data Preparation

Data was prepared for statistical analysis by replacing values below detection with a zero value. Given the small sample set, values above detection were reset to the value representing the upper detection limit for inclusion. Data transformations were performed on an element by element basis in order to normalize the distribution of values, in order to satisfy the requirement of normally distributed data for statistical techniques. Data were either normalized using a natural logarithm ("_3" suffix in figures and tables) or a double natural logarithm ("_4" suffix in figures and tables) transformation.

7.4.2 Correlations

A linear correlation was performed for 26 elements using the Pearson's Product Moment Correlation coefficient (Pearson, 1896). The Spearman's Rank Correlation Coefficient (Spearman, 1904), values are ranked and then a Pearsons Correlation is performed on the ranked values. Ranking of the values can be useful for data which, even when transformed, is not entirely normally distributed. Results vary between -1 and 1; a value of 1 represents a perfect positive correlation and a value of -1 represents a perfect negative correlation. A full set of correlation results are available in table 7.1 (Pearson) and 7.2 (Spearman).

Gold shows statistically significant (≥0.70) correlations with Zn (0.98) using Pearson's method and, K (0.83) and Ti (0.70) using Spearman's method. The association of Au with Zn may be due to co-precipitation of Au and Sphalerite ((Zn,Fe)S) during late-stage Au-Ag-Pb-Zn veins. The correlation between Au and K may point to a strong association of gold within veins which have potassic (K-rich) alteration. The presence of hydrothermal rutile (TiO_2) in association with gold and/or gold bearing minerals may be the cause for the correlation of Au with Ti. These associations should be confirmed with a mineralogical investigation of the property and should be considered as hypotheses only.

	Au_3	Ag_3	Al_3	As_3	Ba_3	Bi_3	Ca_3	Cd_3	Co_4	Cr_3	Cu_3	Fe_3	K_3	Mg_3	Mn_3	Mo_3	Na_3	Ni_4	Pb_3	Sb_4	Sc_3	Sr_3	Ti_3	V_3	W_4	Zn_4
Au_3		0.00	0.00	0.00	0.02	0.00	0.11	0.00	0.04	0.00	0.00	0.00	0.66	0.05	0.38	0.00	0.00	0.04	0.00	0.09	0.00	0.09	0.30	0.08	0.23	0.98
Ag_3	0.47		0.38	0.00	0.00	0.00	0.95	0.00	0.00	0.00	0.00	0.02	0.00	0.76	0.33	0.09	0.11	0.78	0.00	0.00	0.79	0.00	0.00	0.39	0.77	0.19
Al_3	-0.32	0.06		0.00	0.00	0.98	0.00	0.05	0.00	0.00	0.91	0.81	0.00	0.00	0.43	0.06	0.00	0.01	0.05	0.78	0.00	0.00	0.98	0.00	0.04	0.01
As_3	0.39	0.34	-0.21		0.13	0.00	0.42	0.01	0.00	0.00	0.00	0.00	0.27	0.00	0.07	0.67	0.00	0.09	0.00	0.07	0.00	0.98	0.00	0.00	0.11	0.94
Ba_3	-0.16	0.27	0.73	0.11		0.01	0.11	0.00	0.00	0.00	0.07	0.12	0.00	0.14	0.15	0.14	0.00	0.05	0.49	0.90	0.00	0.00	0.32	0.02	0.25	0.00
Bi_3	0.53	0.65	0.00	0.50	0.20		0.16	0.00	0.00	0.01	0.00	0.00	0.00	0.04	0.06	0.77	0.38	0.73	0.00	0.01	0.33	0.06	0.01	0.00	0.78	0.01
Ca_3	-0.11	0.00	0.21	-0.06	0.11	-0.10		0.00	0.88	0.84	0.00	0.17	0.19	0.00	0.00	0.19	0.00	0.83	0.27	0.02	0.00	0.00	0.53	0.78	0.55	0.09
Cd_3	0.26	0.61	0.14	0.20	0.33	0.41	0.22		0.36	0.03	0.00	0.24	0.00	0.26	0.03	0.16	0.02	0.73	0.62	0.49	0.24	0.00	0.01	0.02	0.68	0.06
Co_4	0.14	0.21	0.21	0.31	0.27	0.30	0.01	0.06		0.28	0.00	0.00	0.00	0.04	0.18	0.00	0.04	0.00	0.00	0.32	0.00	0.08	0.76	0.13	0.47	0.01
Cr_3	-0.57	-0.27	0.50	-0.39	0.34	-0.17	-0.01	-0.16	0.08		0.00	0.00	0.02	0.01	0.00	0.00	0.00	0.00	0.00	0.02	0.00	0.02	0.00	0.00	0.00	0.84
Cu_3	0.27	0.37	-0.01	0.54	0.13	0.55	-0.26	0.21	0.36	-0.24		0.00	0.22	0.04	0.10	0.20	0.01	0.72	0.00	0.63	0.45	0.09	0.11	0.33	0.40	0.00
Fe_3	0.22	0.17	-0.02	0.46	0.11	0.31	-0.10	-0.08	0.40	-0.24	0.63		0.72	0.01	0.00	0.05	0.00	0.06	0.00	0.99	0.01	0.42	0.10	0.03	0.12	0.00
K_3	-0.03	0.27	0.70	0.08	0.76	0.22	0.09	0.34	0.23	0.16	0.09	-0.03		0.86	0.76	0.51	0.00	0.34	0.83	0.73	0.00	0.00	0.19	0.38	0.61	0.14
Mg_3	-0.14	-0.02	0.34	-0.32	0.10	-0.14	0.47	-0.08	0.15	0.19	-0.14	0.20	0.01		0.00	0.31	0.01	0.00	0.57	0.38	0.00	0.00	0.05	0.00	0.39	0.00
Mn_3	-0.06	-0.07	0.06	0.13	0.10	-0.13	0.55	0.16	0.09	-0.22	0.12	0.37	-0.02	0.23		0.21	0.04	0.39	0.67	0.00	0.00	0.00	0.22	0.25	0.00	0.00
Mo_3	-0.32	-0.12	0.13	-0.03	0.11	0.02	-0.09	-0.10	0.20	0.51	0.09	0.14	-0.05	0.07	-0.09		0.00	0.00	0.27	0.02	0.22	0.53	0.02	0.82	0.01	0.44
Na_3	-0.33	0.11	0.75	-0.27	0.59	0.06	0.22	0.17	0.14	0.63	-0.17	-0.24	0.55	0.20	-0.14	0.24		0.01	0.08	0.00	0.00	0.00	0.75	0.00	0.00	0.82
Ni_4	-0.15	-0.02	0.18	-0.12	0.14	-0.02	0.02	-0.02	0.32	0.29	0.03	0.13	-0.07	0.30	-0.06	0.22	0.20		0.56	0.00	0.00	0.28	0.05	0.00	0.01	0.17
Pb_3	0.28	0.23	-0.14	0.25	-0.05	0.48	-0.08	0.04	0.20	-0.20	0.33	0.45	0.02	-0.04	0.03	0.08	-0.13	0.04		0.00	0.22	0.96	0.40	0.05	0.65	0.00
Sb_4	0.12	0.22	0.02	-0.13	0.01	0.19	-0.17	-0.05	0.07	0.17	-0.03	0.00	-0.02	0.06	-0.33	0.16	0.20	0.23	0.21		0.93	0.47	0.61	0.05	0.01	0.46
Sc_3	-0.21	0.02	0.59	-0.27	0.40	-0.07	0.44	0.08	0.22	0.33	-0.05	0.17	0.33	0.66	0.38	0.09	0.47	0.27	-0.09	-0.01		0.00	0.60	0.00	0.47	0.00
Sr_3	-0.12	0.20	0.41	0.00	0.41	0.14	0.75	0.31	0.12	0.16	-0.12	0.06	0.31	0.45	0.28	0.04	0.50	0.08	0.00	-0.05	0.53		0.02	0.24	0.44	0.01
Ti_3	0.07	0.24	0.00	0.20	0.07	0.18	0.05	0.18	-0.02	-0.22	0.11	0.12	0.09	0.14	-0.09	-0.16	0.02	-0.14	-0.06	0.04	0.04	0.16		0.08	0.02	0.32
V_3	-0.12	-0.06	0.43	-0.41	0.16	-0.21	-0.02	-0.17	0.11	0.23	-0.07	0.15	0.06	0.59	-0.08	0.02	0.22	0.38	-0.14	0.14	0.58	0.08	0.12		0.10	0.01
W_4	-0.09	-0.02	0.15	-0.11	0.08	0.02	-0.04	0.03	0.05	0.31	-0.06	-0.11	0.04	0.06	-0.21	0.18	0.21	0.18	-0.03	0.19	0.05	0.05	-0.17	0.12		0.90
Zn_4	0.00	0.09	0.19	0.01	0.21	0.17	0.12	0.13	0.18	0.01	0.28	0.40	0.10	0.25	0.34	0.05	0.02	0.10	0.49	-0.05	0.26	0.19	-0.07	0.18	-0.01	

Table 7.1 *Pearson's Correlation Coefficient for detected elemental concentrations at McConnells Jest. Points of higher correlation (>0.59) are colour coded. Blue = 0.60 to 0.69, Green = 0.70 to 0.79, Orange = 0.80 to 0.89, Red = >0.90.*

	Au_3	Ag_3	Al_3	As_3	Ba_3	Bi_3	Ca_3	Cd_3	Co_4	Cr_3	Cu_3	Fe_3	K_3	Mg_3	Mn_3	Mo_3	Na_3	Ni_4	Pb_3	Sb_4	Sc_3	Sr_3	Ti_3	V_3	W_4	Zn_4
Au_3		0.00	0.00	0.00	0.55	0.00	0.14	0.00	0.03	0.00	0.00	0.01	0.83	0.01	0.62	0.00	0.00	0.01	0.02	0.59	0.00	0.38	0.70	0.13	0.02	0.41
Ag_3	0.49		0.08	0.00	0.00	0.00	0.94	0.00	0.00	0.13	0.00	0.18	0.00	0.64	0.15	0.53	0.08	0.78	0.49	0.01	0.61	0.00	0.00	0.44	0.90	0.93
Al_3	-0.25	0.12		0.00	0.00	0.42	0.00	0.06	0.00	0.00	0.75	0.15	0.00	0.00	0.45	0.03	0.00	0.01	0.43	0.77	0.00	0.00	0.59	0.00	0.12	0.00
As_3	0.43	0.31	-0.20		0.24	0.00	0.50	0.00	0.00	0.00	0.00	0.00	0.52	0.00	0.00	0.21	0.00	0.13	0.02	0.01	0.00	0.98	0.08	0.00	0.00	0.64
Ba_3	-0.04	0.31	0.72	0.08		0.00	0.04	0.00	0.00	0.00	0.13	0.04	0.00	0.20	0.01	0.96	0.00	0.18	1.00	0.99	0.00	0.00	0.60	0.09	0.24	0.00
Bi_3	0.57	0.64	0.06	0.46	0.21		0.29	0.00	0.00	0.09	0.00	0.00	0.00	0.01	0.19	0.90	0.46	0.50	0.00	0.15	0.36	0.01	0.11	0.00	0.89	0.25
Ca_3	-0.10	0.00	0.21	-0.05	0.15	-0.08		0.01	0.90	0.78	0.00	0.04	0.38	0.00	0.00	0.11	0.00	0.77	0.25	0.10	0.00	0.00	0.39	0.96	0.19	0.36
Cd_3	0.38	0.62	0.13	0.22	0.30	0.47	0.18		0.58	0.07	0.01	0.12	0.00	0.15	0.10	0.02	0.13	0.38	0.24	0.22	0.68	0.00	0.02	0.03	0.44	0.57
Co_4	0.16	0.23	0.28	0.30	0.30	0.27	0.01	0.04		0.38	0.00	0.00	0.00	0.02	0.11	0.01	0.06	0.00	0.01	0.60	0.00	0.18	0.70	0.23	0.55	0.02
Cr_3	-0.56	-0.11	0.51	-0.41	0.28	-0.12	0.02	-0.13	0.06		0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.02	0.04	0.00	0.00	0.54
Cu_3	0.30	0.33	0.02	0.55	0.11	0.50	-0.31	0.17	0.38	-0.22		0.00	0.12	0.04	0.15	0.25	0.00	0.23	0.00	0.11	0.10	0.07	0.43	0.14	0.67	0.00
Fe_3	0.17	0.09	0.10	0.41	0.15	0.20	-0.14	-0.11	0.40	-0.25	0.67		0.25	0.00	0.00	0.29	0.00	0.01	0.00	0.24	0.04	0.66	0.15	0.02	0.21	0.00
K_3	0.02	0.30	0.73	0.05	0.83	0.23	0.06	0.33	0.24	0.21	0.11	0.08		0.48	0.51	0.38	0.00	0.47	1.00	0.87	0.00	0.00	0.16	0.07	0.20	0.08
Mg_3	-0.18	-0.03	0.38	-0.33	0.09	-0.18	0.43	-0.10	0.16	0.21	-0.14	0.22	0.05		0.04	0.08	0.00	0.00	0.96	0.41	0.00	0.00	0.01	0.00	0.11	0.00
Mn_3	0.04	-0.10	0.05	0.21	0.18	-0.09	0.47	0.12	0.11	-0.28	0.10	0.34	0.05	0.15		0.00	0.07	0.72	0.30	0.00	0.00	0.01	0.03	0.17	0.00	0.00
Mo_3	-0.33	-0.04	0.15	-0.09	0.00	0.01	-0.11	-0.17	0.18	0.54	0.08	0.08	-0.06	0.12	-0.23		0.02	0.00	0.95	0.00	0.03	0.95	0.04	0.91	0.00	0.49
Na_3	-0.30	0.13	0.71	-0.34	0.55	0.05	0.24	0.11	0.13	0.57	-0.22	-0.21	0.53	0.24	-0.13	0.16		0.01	0.47	0.01	0.00	0.00	0.42	0.00	0.06	0.50
Ni_4	-0.18	0.02	0.18	-0.11	0.10	-0.05	-0.02	-0.06	0.29	0.29	0.08	0.18	-0.05	0.26	-0.03	0.28	0.18		0.76	0.04	0.00	0.46	0.23	0.00	0.03	0.51
Pb_3	0.17	0.05	-0.06	0.17	0.00	0.31	-0.08	-0.08	0.18	-0.22	0.30	0.43	0.00	0.00	0.07	0.00	-0.05	0.02		0.06	0.66	0.75	0.34	0.29	0.59	0.00
Sb_4	0.04	0.18	0.02	-0.19	0.00	0.10	-0.12	-0.09	0.04	0.20	-0.11	-0.08	-0.01	0.06	-0.31	0.20	0.17	0.15	0.14		0.58	0.57	0.83	0.25	0.00	0.60
Sc_3	-0.25	0.04	0.61	-0.32	0.37	-0.07	0.42	0.03	0.20	0.42	-0.12	0.15	0.34	0.68	0.28	0.15	0.55	0.28	-0.03	0.04		0.00	0.28	0.00	0.11	0.00
Sr_3	-0.06	0.23	0.41	0.00	0.38	0.18	0.73	0.29	0.09	0.16	-0.13	0.03	0.31	0.44	0.19	0.00	0.46	0.05	0.02	-0.04	0.53		0.00	0.21	0.61	0.03
Ti_3	0.03	0.20	0.04	0.12	0.04	0.11	0.06	0.16	-0.03	-0.14	0.06	0.10	0.10	0.18	-0.16	-0.14	0.06	-0.08	-0.07	0.01	0.08	0.21		0.05	0.09	0.68
V_3	-0.11	-0.06	0.41	-0.42	0.12	-0.23	0.00	-0.15	0.09	0.20	-0.10	0.17	0.13	0.57	-0.10	0.01	0.24	0.29	-0.07	0.08	0.53	0.09	0.14		0.14	0.02
W_4	-0.17	0.01	0.11	-0.21	-0.08	-0.01	-0.09	-0.05	0.04	0.36	-0.03	-0.09	-0.09	0.11	-0.21	0.28	0.13	0.15	-0.04	0.24	0.11	-0.04	-0.12	0.10		0.65
Zn_4	-0.06	-0.01	0.25	-0.03	0.21	0.08	0.07	0.04	0.16	0.04	0.26	0.38	0.12	0.25	0.29	0.05	0.05	0.05	0.52	-0.04	0.28	0.15	-0.03	0.16	0.03	

Table 7.2 *Spearman's Rank Correlation Coefficient for detected elemental concentrations at McConnells Jest. Points of higher correlation (>0.59) are colour coded. Blue = 0.60 to 0.69, Green = 0.70 to 0.79, Orange = 0.80 to 0.89, Red = >0.90.*

7.4.3 Principal Component Analysis (PCA)

The suite of elements from assay data was split into two sets, 1) an "alteration" set (Al, Ba, Ca, Cr, K, Mg, Mn, Na, Sc,Sr, Ti, V), and 2) an "ore mineralization" set (Au, Ag, As, Bi, Cd, Co, Cu, Fe, Mo, Ni, Pb, Sb, W, Zn) in order to most clearly portray these complementary aspects of the sample lithologies. Data used in the PCA was selected and transformed using the same criteria as for the element vs element correlations. The PAST software of Hammer et al. (2001) was used to carry out the PCA. These associations should be confirmed with a mineralogical investigation of the property and should be considered as hypotheses only.

Alteration Set (Fig 7.5): Three distinct grouping of elements are present:

1. Ti: Ti shows a large divergence from the rest of the elements within the alteration grouping. This may point to, as stated above, a hydrothermal rutile (TiO_2) phase.

2. Magmatic (Al, Ba, Ca, K, Mg, Mn, Na, Sc, Sr, V): The close grouping of these elements is most likely explained by their dominant residence within the host rocks to mineralisation.

3. Cr: The proximity to the other magmatic elements may point to the residence of Cr within intermediate dykes mapped on the property.

There is no clear pattern of association of any of these elements with high grade gold values in Fig 7.5.

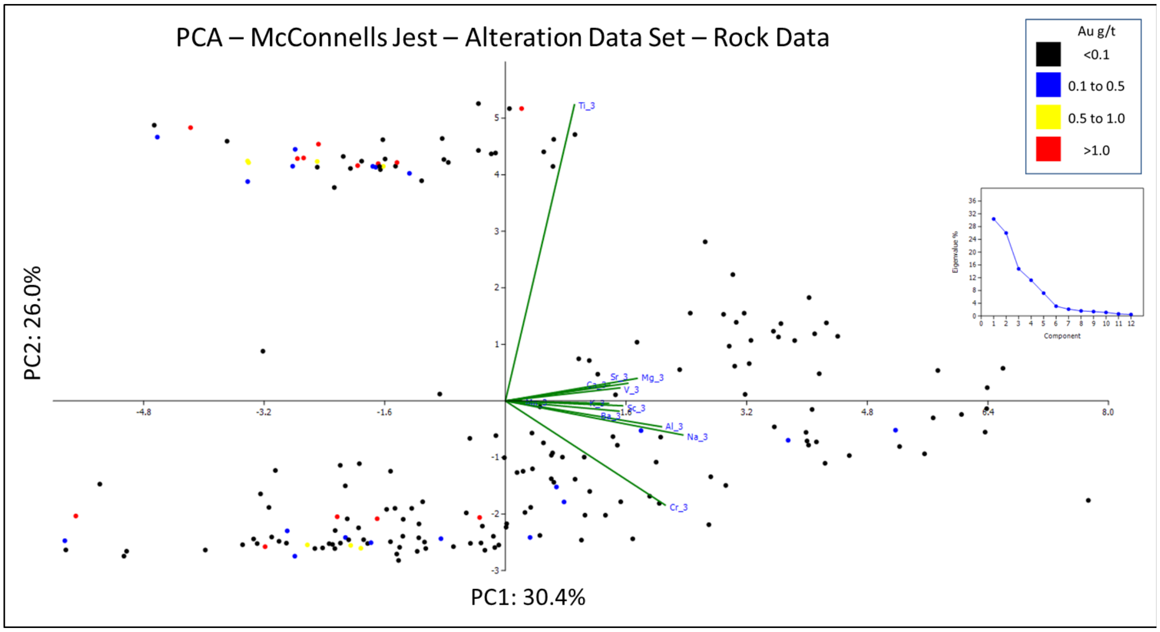

Figure 7.5 *PCA Scatter plot for the alteration set of elements at McConnells Jest.*

Ore Mineralisation Set (Fig 7.6): Five distinct groupings of elements are present:

1. W, Ni, Sb: A number of Ni-Sb minerals could be the result of this association. Minor volumes of scheelite ($CaWO_4$) are common in quartz veins in IRGS and are the most probably residence of W in this group.

2. Mo, Pb, Zn: The grouping of these elements may point to a mineralisation stage with molybdenite (MoS_2), galena (PbS) and sphalerite ((Zn,Fe)S).

3. Fe, Bi, Cu, Ag: The sulfosalt mineral tetrahedrite ($(Ag,Cu,Fe)_{12}Sb_4S_{13}$) and bismuth minerals may be the dominant residence of these minerals. It should be noted that Sb, an important component of tetrahedrite, is not associated with this group.

4. Cd, Au: Cd can occur as an impurity in sphalerite. Given the linear correlation of Au with Zn, the association of native gold with sphalerite is suggested here.

5. As: Arsenopyrite (FeAsS) veins are common on the property. Although these generally return elevated Au assay grades, the trajectory of the As component is not towards high Au values. This may indicate that arsenopyrite is a host to gold but is not genetically related to the same fluid which precipitated gold.

Groups 3 and 4 trend towards high grade gold values (red dots) on Fig. 7.6, suggesting that these may be either i) two discrete Au mineralizing events or ii) sub-stages of a broader, single Au mineralisation event.

An investigation into the hydrothermal paragenesis at McConnells Jest is required to test the above hypotheses.

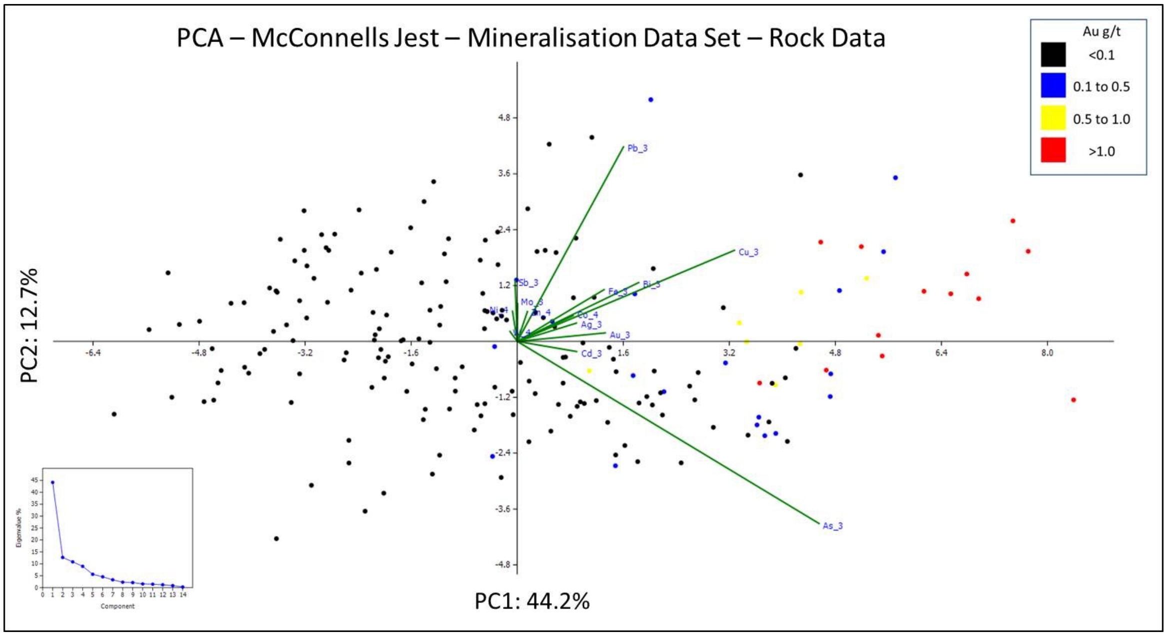

Figure 7.6 *PCA Scatter plot for the mineralisation set of elements at McConnells Jest.*

7.5 Surficial Geology

The property underwent glaciation during the McConnell glaciation (>23,000 years ago; Bond, 1999). It has been demonstrated through a number of field seasons that the ground is covered by basal till. Sampling of the property has been carefully completed, where possible, at sufficient depths (i.e. > 0.50 m) to avoid sampling possible shallow loess deposits. Generally, samples were collected at shallower depths (i.e. <0.3 m) in subalpine terrain with limited overburden or in boulder fields, as this was sufficient to avoid sampling glacial material (Golden Predator Canada Corp., 2013).

7.6 Adjacent Properties

Two major mineral properties lie adjacent to McConnells Jest, the Dublin Gulch gold deposit to the west and the Keno Hill silver district to the south east (see Fig. 7.7). The Dublin Gulch and Keno Hill properties are owned by Victoria Gold Corp. and Alexco Resources Corp., respectively.

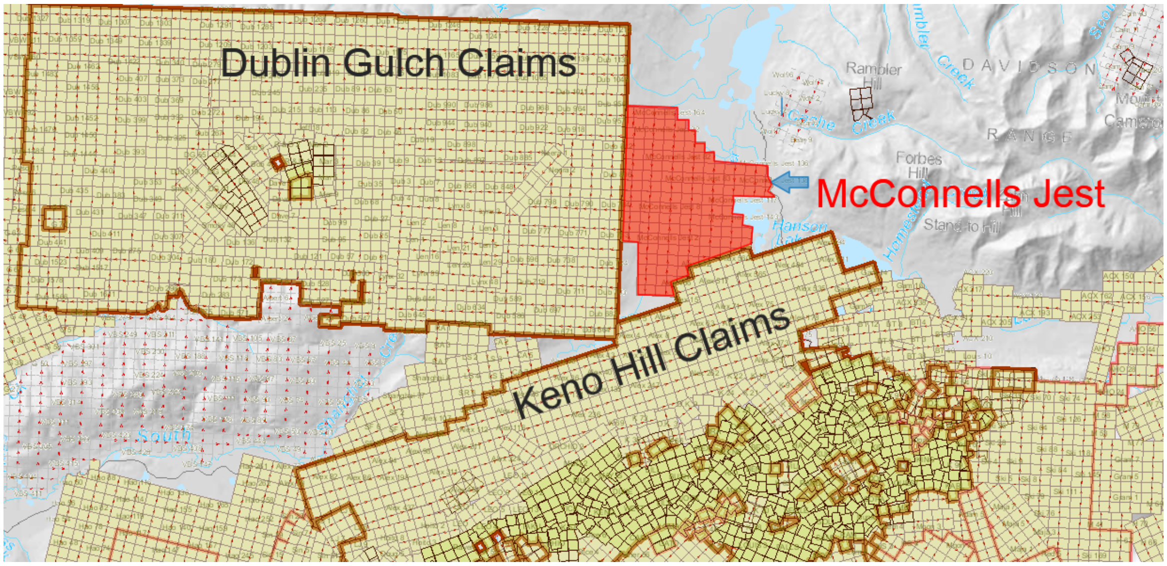

Figure 7.7 *Claim blocks adjacent to the McConnells Jest property.*

Victoria Gold Corp. acquired the Dublin Gulch property in June, 2009 through the acquisition of StrataGold Corp. Victoria Gold Corp. holds 1,912 quartz claims, 10 quartz leases, and 1 federal Crown grant on the Dublin Gulch property. Currently, Dublin Gulch is an advanced-stage gold exploration project with around 630 diamond drill holes to date and a global resource of 6.3 M oz (4.8 M oz – 222 Mt @ 0.68g/t Au indicated; 1.5 M oz - 78 Mt @ 0.60 g/t Au inferred). The Eagle Zone, which has a 2.3 M oz (proven and probable) gold reserve contained within it (Wardrop Engineering Inc., 2011, 2012) is the most significant zone of mineralization. The property also hosts the Olive Zone - a recent gold exploration target, the Wolf (Mar) tungsten skarn (Indicated: 12.7 Mt @ 0.31 % WO_3, 86.2 M lbs contained WO_3; Inferred: 1.3 Mt @ 0.30 % WO_3, 8.9 M lbs contained WO_3; SRK Consulting, 2008) and the Rex-Peso silver prospect (Probable: 0.14 Mt @ 716 g/t Ag, 3.7 % Pb; Hitchins and Orssich, 1995). Mineralization within the Eagle and Olive Zones at Dublin Gulch is considered a holotypic example of an intrusion-related gold system (IRGS) (e.g. Lang and Baker, 2001). It is likely that any significant gold mineralization at McConnells Jest would align with this deposit type, but cannot be confirmed until further work on the property has been conducted.

Alexco Resources Corp. acquired the Keno Hill property in February, 2006 through the purchase of the assets of the bankrupt United Keno Hill Mines Limited. The property was purchased using a wholly-owned subsidiary, Elsa Reclamation and Development Company. Alexco Resources Corp. holds 695 quartz mining leases and 871 quartz mining claims on the Keno Hill property to the south of McConnells Jest. The Keno Hill silver district is a polymetallic silver-lead-zinc vein district that has been mined since the Klondike gold rush of the 1890's. Historical production up to 1989 for the Keno Hill district is 117.5M oz Ag, 710M lbs Pb and 436M lbs Zn (Cathro, 2006). From 2006 to 2012, Alexco Resources drilled 405 diamond drill holes on the property. The most prominent zone in the district at current is the Bellekeno silver mine (Indicated: 365,000 t @ 659 g/t Ag, 5.3 % Pb, 5.3 % Zn; Inferred: 243,000 @ 428 g/t Ag, 4.1 % Pb, 5.1 % Zn – Alexco Resources Corp., 2012). Historical production (since 1919) for the Bellekeno mine is 7.9 M oz Ag (SRK Consulting, 2014). Commercial production of the Bellekeno mine began on January 1, 2011 until operations were temporarily suspended in August, 2013. The nominal rate of production for the Bellekeno mine during this period was 250 tonnes per day. Production is expected to begin again in Q3 2015. The

Lucky Queen, Flame & Moth, Onek and Bermingham occurrences provide additional silver, lead and zinc resources. The deposit model of Hantelmann (2013) to describe mineralization at Bellekeno is unlikely to occur at McConnells Jest, but cannot be entirely ruled out until further work has taken place.

8.0 Exploration Programs

In 2010 and with the assistance of a YMIP grant, Bill Koe Carson staked the property and collected 12 stream samples, 44 soil samples and 28 rock samples (Bourne, 2011). Stream sample MJSED-004 returned 11.7 ppm Au, MJSED-006 returned 0.558 ppm Au and MJSED-009 returned 0.305 ppm Au (Figure 2-2). Of the 28 rock samples, two had weakly anomalous gold assays in the 0.1 to 0.2 ppm range, however several samples contained anomalous pathfinder elements, for example 3722 ppm arsenic and 98 ppm bismuth in sample MJR-24.

In 2011, Golden Predator optioned the ground and contracted All-In Exploration Inc. (Whitehorse, Yukon) to complete the collection of 380 soil samples from a grid covering the western section of the property. Samples were collected every 50m along east-west oriented lines 200m apart, to a total of 19.6km.

The results from this survey outlined several multi-element geochemical anomalies, with sporadic highs (up to 208 ppb Au) and a cluster of elevated values (10 to 17 ppb Au) in the northeast quadrant. This anomaly is around 400m in length and lie within 500m of the anomalous stream sediments collected in 2010. This anomaly is associated with elevated levels of copper and arsenic.

In the southwest part of the grid, there is a strong cluster of arsenic anomalies that are associated with the highest gold result (208 ppb). There were also elevated levels of silver and bismuth in the same quadrant.

Work continued in 2012 with a short field program undertaken again by Golden Predator.

A total of 74 rock samples were taken over the 3 day program on the McConnell claims. The program was undertaken by three geologists and an experienced prospector who has worked with Golden Predator for a number of years. Focus was put on intrusive rocks and sedimentary rocks proximal to those intrusions, as well as rocks which hosted sheeted quartz veins.

2012 work resulted in a number of interesting anomalous targets which warrant follow up work (figures 8 and 9). Most notable is AA064560, a bedrock sample from a quartz-arsenopyrite breccia/vein which assayed over 25 g/t Au. With an orientation of 112 degrees azimuth, and a 38 degree dip, the sample shows a similarity to Dublin Gulch style structural extensional veining. In addition, a number of samples assaying over 0.3 g/t Au were discovered, and a soil sample which assayed 1.47 g/t Au at the north of the property were also discovered.

Figure 8.1 *Example of scorodite vein at surface. Sample AA064560 has an assay grade >25g/t Au.*

It is understood that no work was undertaken in 2013 due to economic setbacks. Golden Predator did not return to the site, and the claim owner could not raise the cash to return to the property that year.

In 2014, the claim owner did return to the property, and with a small team collected 102 rock samples from across the property, in particular in the two anomalous areas previously identified by Golden Predator in their soil program.

These rocks samples returned values up to 28.8ppm Au (sample 14474), with an additional 16 samples returning grades in excess of 0.5ppm. Many of the samples were from sheeted veins or scorodite exposures within the two anomalous zones. These zones were subsequently named Bullion Blister (in the west) and Pink Mountain (in the east).

Pink Mountain has an abundance of sheeted vein systems, and covers an area approximately 500m by 375m (although this remains open on three sides). The grades here slightly lower but more consistent, around 1g/t Au.

Bullion Blister hosts many of the scorodite veins in oxidised rock, and as such has returned the highest assays, including the 28.8g/t.

9.0 Geochemical Analytical Procedures

Geochemical analyses for samples from 2010 to 2014 are summarized in Table 9.1. A more detailed description of the analytical techniques is presented below, split by the year of analysis.

Certificate Number	Lab	Type	Total Samples	Received	Completed	Method
10-360-02341	INSP	Soil	2	26th July 2010	11th August 2010	Pd-1AT-ICP, Ag-1AT-GV, Au-1AT-AA, Ag-4A-OR, Pt-1AT-ICP
10-360-00307	INSP	Rock	9	28th October 2010	10th December 2010	30-4A-TR
10-360-00308	INSP	Moss	1	24th September 2010	4th October 2010	30-4A-TR
10-360-00309	INSP	Rock	28	30th September 2010	12th October 2010	30-4A-TR
10-360-03010	INSP	Soil	44	30th September 2010	19th October 2010	30-4A-TR, Au-1AT-AA
10-360-03200	INSP	Pulp	37	13th October 2010	18th October 2010	Au-1AT-AA
WHI1101802	ACME	Soil	320	24th October 2011	5th December 2011	ACM 1DX15
WHI1101803	ACME	Soil	58	24th October 2011	5th December 2011	ACM 1DX15
12Y640856	AGAT	Rock	78	-	19th October 2012	AGAT 201074
12Y640884	AGAT	Soil	242	-	19th October 2012	AGAT 201074
WHI14000057	ACME	Rock	102	24th July 2014	13th August 2014	FA430, AQ200, FA530
WHI14000057M	ACME	Metallic Screen	4	24th September 2014	8th October 2014	FS651, FA550-Au

Table 9.1 *Geochemical Analytical Procedures for 2010 to 2014.*

9.1 2010

All samples from the 2010 field season were sent to Inspectorate Laboratories, Whitehorse, YT., Canada. A total of 37 rock (certificates: 10-360-00307 – 9 samples, 10-360-00309 – 28 samples), 44 soil (certificate: 10-360-03010) were analyzed for 30 elements using inductively coupled plasma emission spectroscopy (ICP-ES) package “30-4A-TR”. A 4-acid aqua regia digestion was performed on a 0.5 g split of the sample and subsequently analyzed using ICP-ES.

46 soil (certificates: 10-360-02341 – 2 samples, 10-360-03010 – 44 samples) and 37 pulp samples (certificate: 10-360-03200) were analyzed for gold using the “Au-1AT-AA” fire assay package. A lead collection fire assay fusion was made from 30 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed by atomic absorption spectroscopy (AAS).

Platinum, Palladium and Silver were also tested for in 2 soil samples (certificate: 10-360-02341) using the “Pt-1AT-ICP”, “Pd-1AT-ICP” and, “Ag-1AT-GV” and “Ag-4A-OR”, respectively. For Platinum and Palladium, A lead collection fire assay fusion was made from 50 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed by ICP-ES. For silver in the “Ag-1AT-GV” package, a lead collection fire assay fusion was made from 30 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed by gravimetric analysis. For “Ag-4A-OR” the sample was decomposed using a 4-acid digestion and analyzed for silver using AAS.

9.2 2011

All samples from the 2011 field season were sent to ACME Analytical Laboratories Ltd. in Whitehorse, YT., Canada. A total of 320 soil (certificate: WHI1101802) and 61 rock (certificate: WHI1101803) samples were analyzed using the “ACM 1DX15” package for 36 elements. Samples were dried at 60ºC, 100g of the sample was then sieved using an 80 mesh. Sample splits of 0.5 g are leached in hot modified aqua regia and analyzed using inductively coupled plasma mass spectrometry (ICP-MS).

9.3 2012

All samples from the 2012 field season were sent to AGAT ISO 9001 certified lab in Whitehorse Y.T, Canada.. A total of 242 soil (certificate: 12Y640884) and 78 rock (certificate: 12Y640856) samples were analyzed using the “AGAT 201074” package.

Analysis was by aqua-regia digestion and a mass spectrometer finish with a 52 metal analysis package. What follows are excerpts from the AGAT laboratory mining geochemistry package.

Samples were dried at 60 degrees centigrade, crushed to the point of 75% passing through a 2mm mesh, then split with a Jones riffler splitter or rotary split. The sample was then pulverized to the point of 85% passing through a 75 micrometer mesh. Finally, samples were screened after drying, shaken on an 80 mesh sieve with the positive fraction stored and the negative fraction sent to the laboratory for analysis. This concludes the preparation portion of sampling.

Prepared samples are digested with aqua regia for one hour using temperature controlled hot blocks. Resulting digests are diluted with de-ionized water. Sample splits of 1 gram or routinely used. These 1 gram samples are then ran through a mass spectrometer. Perkin Elmer 7300DV and 8300DV ICP-OES (Optical Emission Spectroscopy) and Perkin Elmer Elan 9000 and NexION ICP-MS (Mass Spectrometer) are used in analysis. Inter-Element Correction (IEC) techniques are used to correct for any spectral interferences (Golden Predator Canada Corp., 2013).

It should be noted that determination of gold by this method is semi-quantitative due to small sample size. Samples with arsenic above detection (>10,000 ppm) were re-run using AAS.

9.4 2014

All samples from the 2014 field season were sent to ACME Labs in Whitehorse, YT., Canada. A total of 102 rock (certificate: WHI14000057) samples were analyzed using the “FA430” package for gold and “AQ200” package for a further 36 elements. Using the “FA430” package, a lead collection fire assay fusion was made from 30 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed by atomic absorption spectroscopy (AAS). For “AQ200”, Sample splits of 0.5 g are leached in hot modified aqua regia and analyzed using inductively coupled plasma mass spectrometry (ICP-MS). Gold samples >10 g/t were re-run using the “FA-530” package, where a lead collection fire assay fusion was made from 30 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed for by gravimetric analysis.

A sub-set of four samples (certificate: WHI14000057M) from the original 102 were selected for additional metallic screening to test for gold using the “FS651” package. Samples underwent metallic pulverizing and a 500 g sample split was screened to 106 µm. Gravimetric analysis was performed on the plus fraction and instrumentation on the minus fraction. Two of the samples >10 g/t gold were re-run using the “FA550-Au” package; a lead collection fire assay fusion was made from 50 g of the sample for total sample decomposition. The resulting silver dore was digested in acid and analyzed by gravimetric analysis.

9.5 QA/QC

All analyses were tested for accuracy and precision using a series of standardized materials, preparation duplicates and procedural blanks at the corresponding laboratories under their respective internal quality control protocol. Field blanks were inserted into soil analyses for 2011 and 2012 at a frequency of ~1 per 50 samples. Table 9.2 lists the standards used for each certificate.

Certificate Number	Standard(s)
10-360-02341	STD-ME-6
10-360-00307	STD-ME-6
10-360-00308	No standard used
10-360-00309	STD-ME-6
10-360-03010	STD-ME-8, STD-OREAS-45P-4A, STD-Oxi67
10-360-03200	STD-Oxi67
WHI1101802	STD DS8
WHI1101803	STD DS8
12Y640856	Standard used but not referenced
12Y640884	Standard used but not referenced
WHI14000057	STD AGPROOF, STD DS10, STD OREAS45EA, STD OXD108, STD OXI121, STD OXN117, STD SP49
WHI14000057M	STD AGPROOF, STD OXD108, STD OXI121, STD OXN117, STD OXP91, STD SP49, STD SQ70

Table 9.2 *Summary of geochemical standards used by laboratories.*

10.0 Exploration Results

The work so far seems to indicate that the McConnells Jest property is geologically very similar to that of Dublin Gulch, and thus has the potential to contain a significant deposit.

10.1 Geochemical Evidence

Through consolidating the geochemical data collected between 2011 and 2014, it becomes clear that there are several zones of interest (Fig. 10.1), although it should be remembered that the entire property has not been covered, and so there is significant potential for other mineral zones.

The main area of interest lies along the SW-NE trending lineations (refer to section 10.2), where multiple elements are in higher proportions, including gold, silver, bismuth, arsenic and lead. These also coincide with veining observed and recorded in field notes.

Around the edges of the pluton, there are several silver and silver-lead anomalies, which could represent either distal (and therefore cooler) systems, or overprinting from Keno Hill style mineral veins.

On the southwestern side of the pluton, there seems to be some correlation to elevated tungsten levels, although the number of samples collected from this area are low so it is tough to draw conclusions.

This assemblage of elements is in line with the an Intrusion Related Gold System, and spatially bears resemblance to the geochemical distributions of Dublin Gulch.

Although levels in the soil samples were low, this is to be expected with the glacial cover across the property. Samples directly from veins show much higher results and are more indicative of grade.

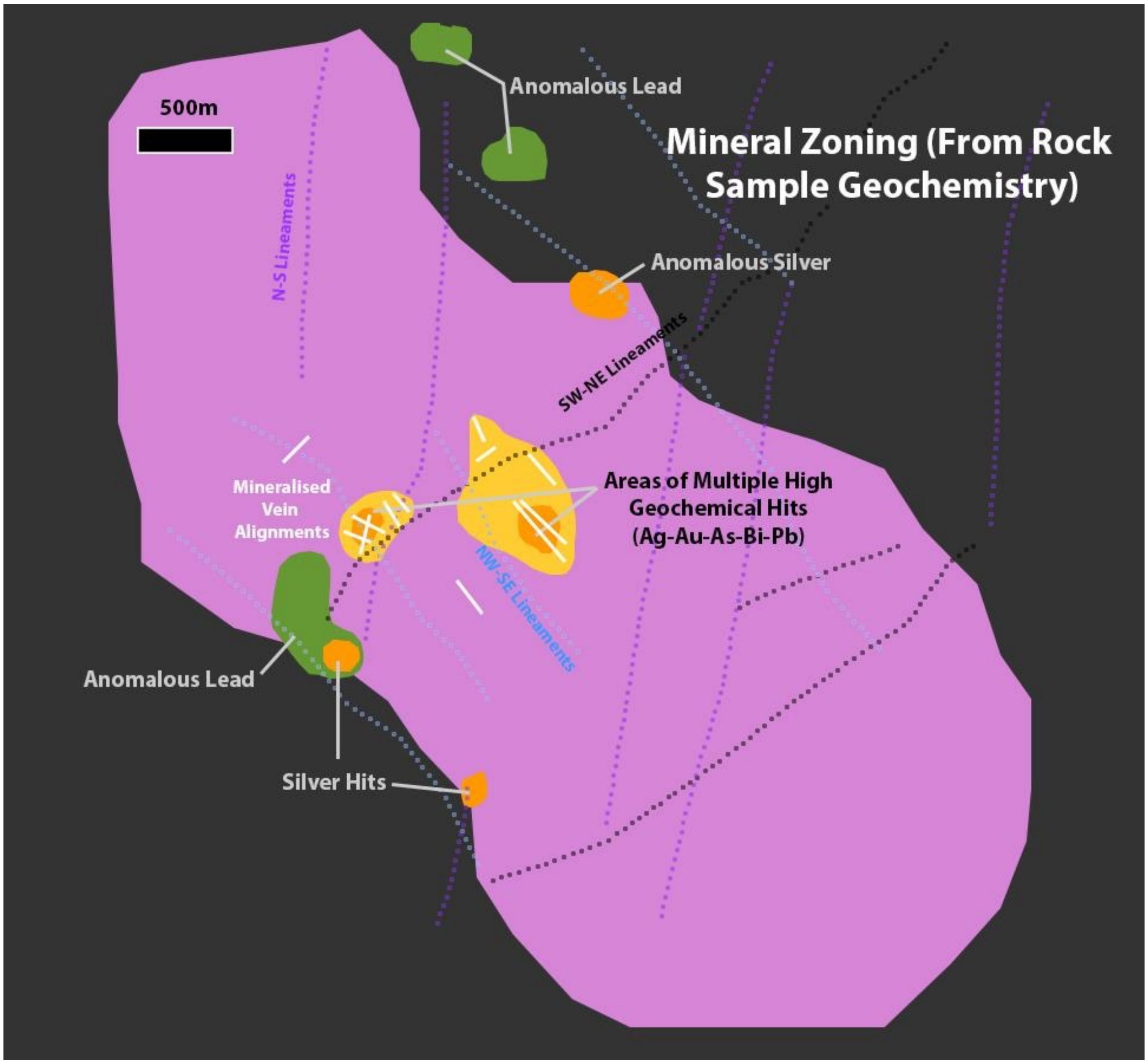

Figure 10.1 *Preliminary element zoning from rock sample geochemistry. Green = Anomalous Lead, Orange = Anomalous Silver, Yellow = Multi-element anomalies (Ag-Au-As-Bi-Pb).*

10.2 Structural Elements

The Energy, Mines and Resources Library in Whitehorse, Yukon Territory, has a considerable library of aerial photos that are publicly accessible. Using their online service, Skyline, the flight lines and plates that intersected McConnells Jest were noted and scans of the photos obtained.

The flight line for this property is ‘A28301, and plates 185 - 188 inclusive cover the ground. The photos were flown in 1996, and have a scale of 1:30,000 (Fig. 10.2).

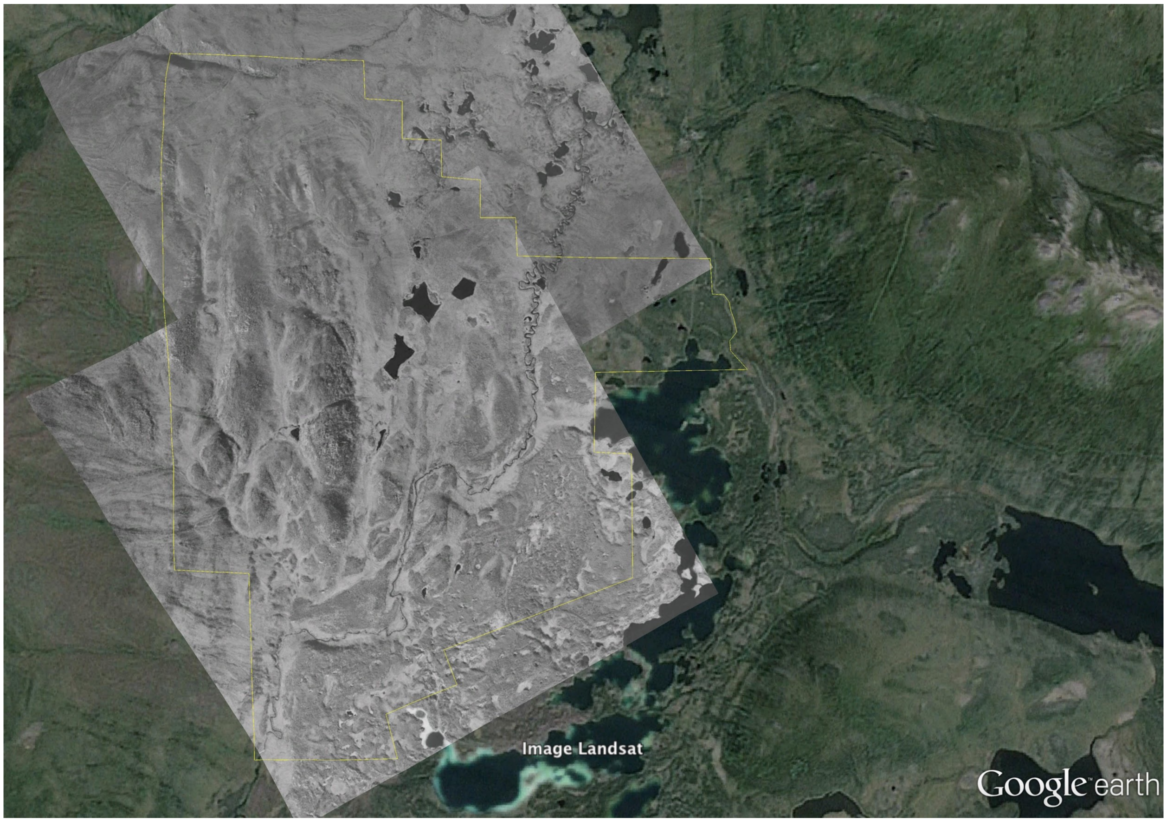

Figure 10.2 *Combined aerial photographs for the McConnells Jest property. Flight line A28301 plates 185 to 188 were combined to form the image. Scale is 1:30,000.*

These images reveal a wealth of information, and the high contrast black and white photos reveal lineations and structures that can be related to ground based observations.

Interpretation of the photos seems to show three sets of lineations (Fig. 10.3);

- North - South trends
- Northwest - Southeast trends
- Southwest - Northeast trends

These have been sketched out on the following map. Initially it seems that the southwest to northeasterly trending lineations are regional, and expand well beyond the boundaries of the property. The areas of intense mineralisation and geochemical highs are found along the area where these regional trends intersect with the other two sets, most notably the northwest to southeast lineaments. This also is similar to the mineralized vein orientations noted in the field, which could be directly related to this trend, or be propagated from smaller riedel shear systems. Further investigations in the field would be required to take more accurate measurements.

It should be noted again that these lineaments align well with the mineralising structures observed at Dublin Gulch, especially with the historic high-grade Olive, Shamrock and Catto veins.

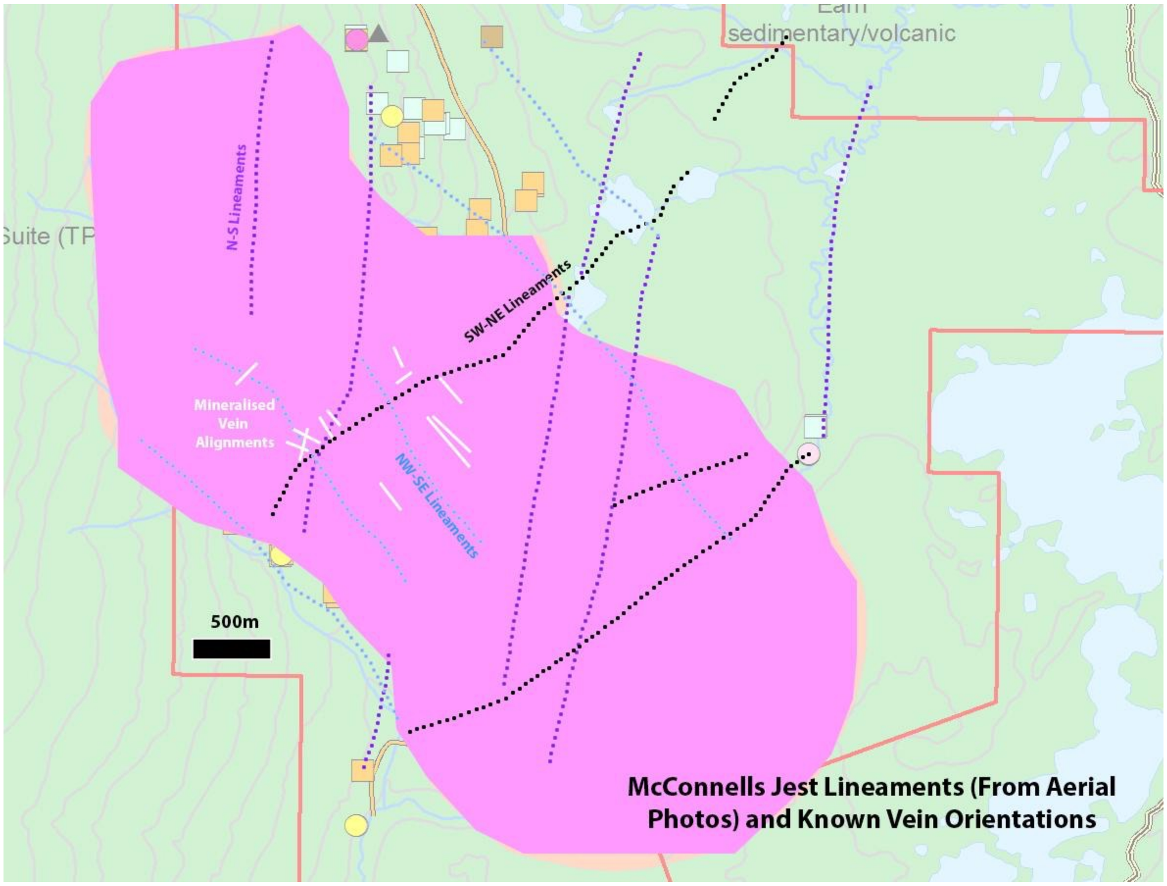

Figure 10.3 *Interpreted structural lineaments on the McConnells Jest property. Three main lineaments are present, N-S (Purple), SE-NW (Light Blue), SW-NE (Black). Known vein orientations are shown in white.*

10.3 Rock Types

No geology map was presented for this project, and historically only the Yukon Geological Survey maps had been used which are inaccurate due to their scale.

During the assimilation and correlation of the data, any information pertaining to rock types was teased from the data, and a database built. At the end of this, over two hundred data points were consolidated and placed in broad lithological categories. Plotting these points, as well as structural data, created an embryonic geological map for further scrutiny (Figs. 10.4, 10.5). It has shown that the pluton is more or less the right volume and orientation, although it is not homogeneous in texture. There are also several outlying igneous exposures, especially in the northeast, which could represent dyke swarms of a cupola of the main pluton. These also correlate with the Ag-Pb mineralisation in the area.

There also seems to be a significant inlier of Earn Group sedimentary rocks in the northwestern quadrant, that also aligns with one of the north-south lineaments. This could be representative of a roof pendant, or faulting that has created an uneven profile around the pluton.

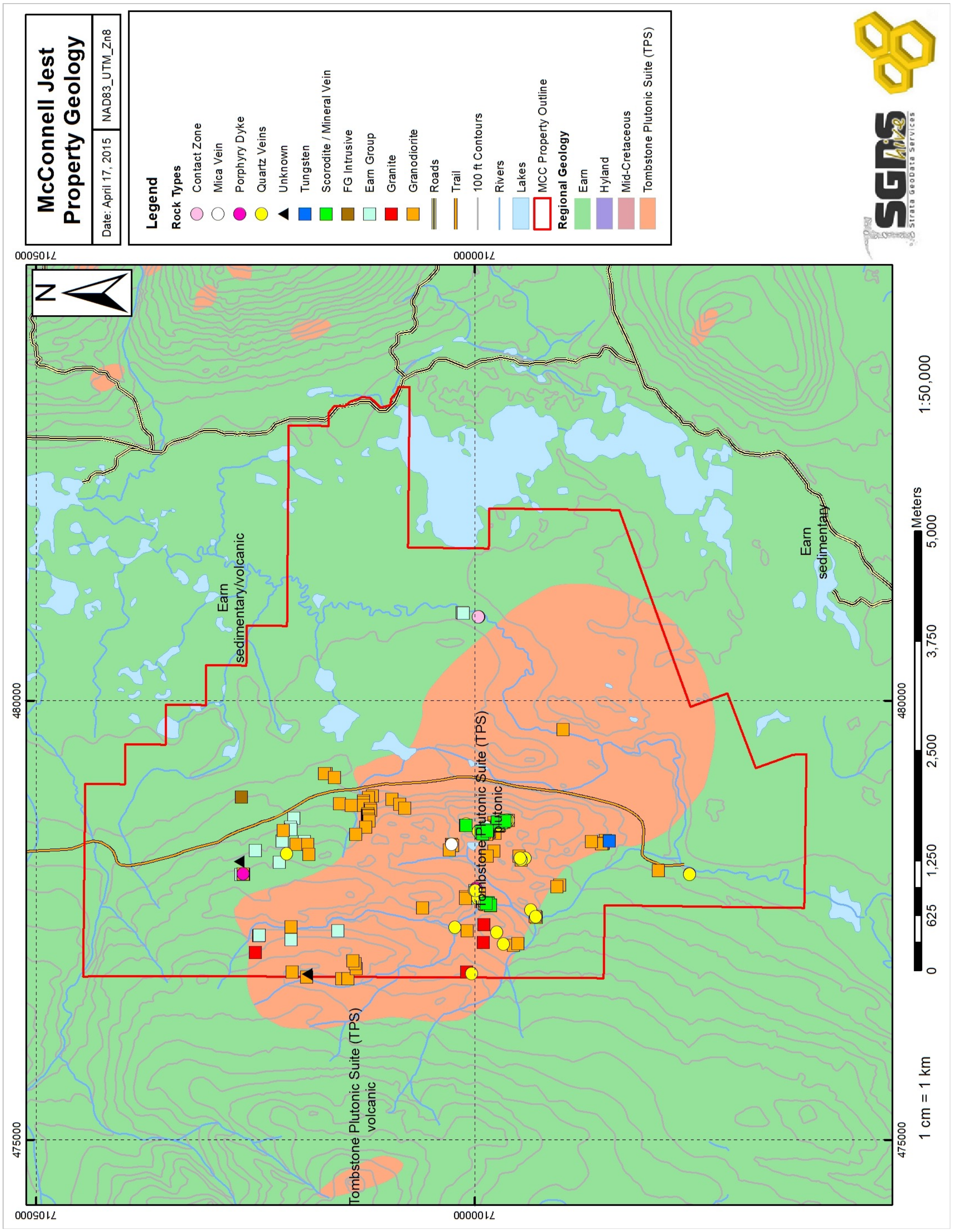

Figure 10.4 *Distribution of rock types at McConnells Jest interpreted from field sampling notes and sample photographs. Rock types are overlain on the current YGS map for the area.*

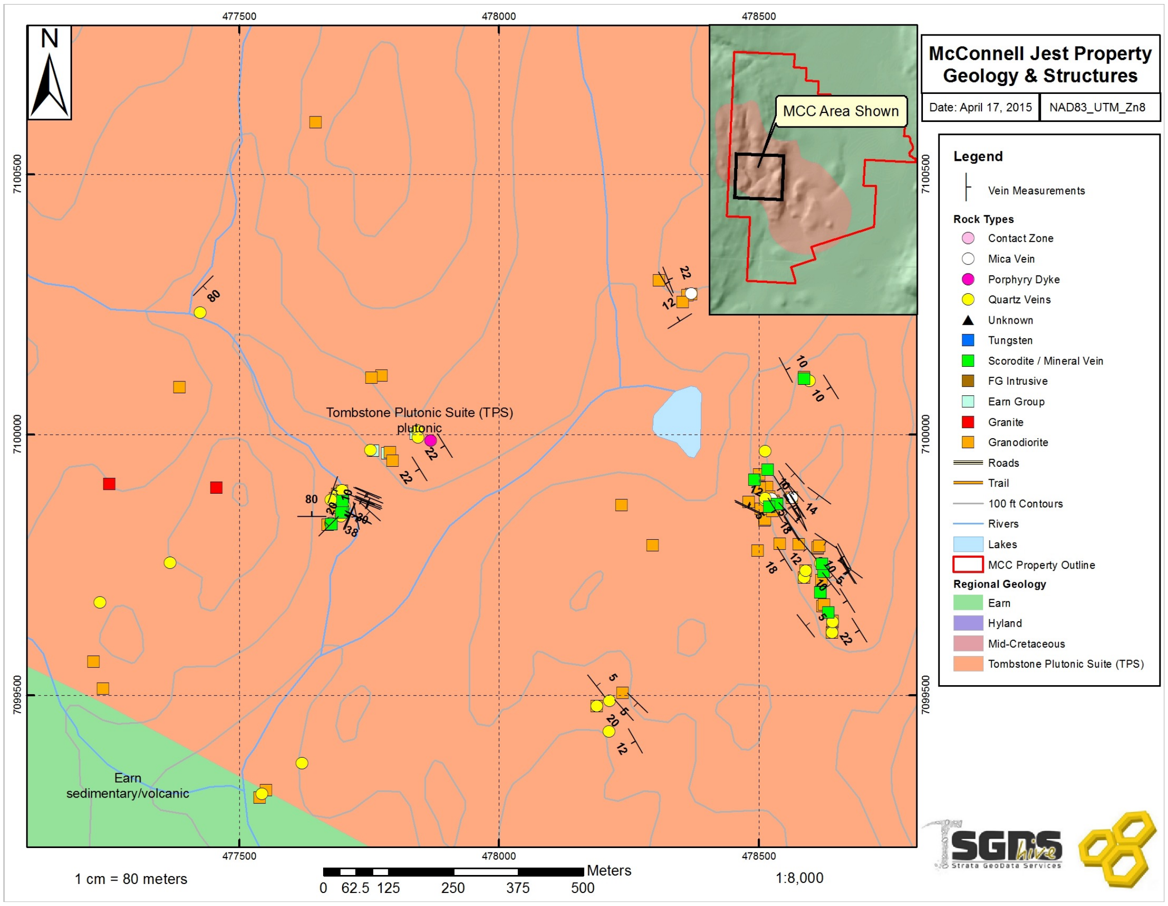

Figure 10.5 *Rock types and vein orientations on the southern shoulder of the central portion of the McConnell pluton.*

10.4 Revised Geology

Using all of this information, the first geological interpretation of McConnells Jest could be drawn (Fig. 10.6). Although this would need to be followed up in the field, the data used is reliable.

It would appear that the north-south lineations represent normal faults, which have displaced blocks of the pluton either 'up' or 'down' relative to one another. This could also account for the finger of Earn Group sedimentary rocks in the northwest quadrant.

The northwest-southeast trends seem to be the main mineralising systems, and align with the mineral veins observed in the field. It is not known if these are fault related, but they seem local to the pluton and do not extend far into the country rock, although further work would be required to test this.

The last trend, those that run southwest to northeast, could be regional shear zones that pass through the entire area.

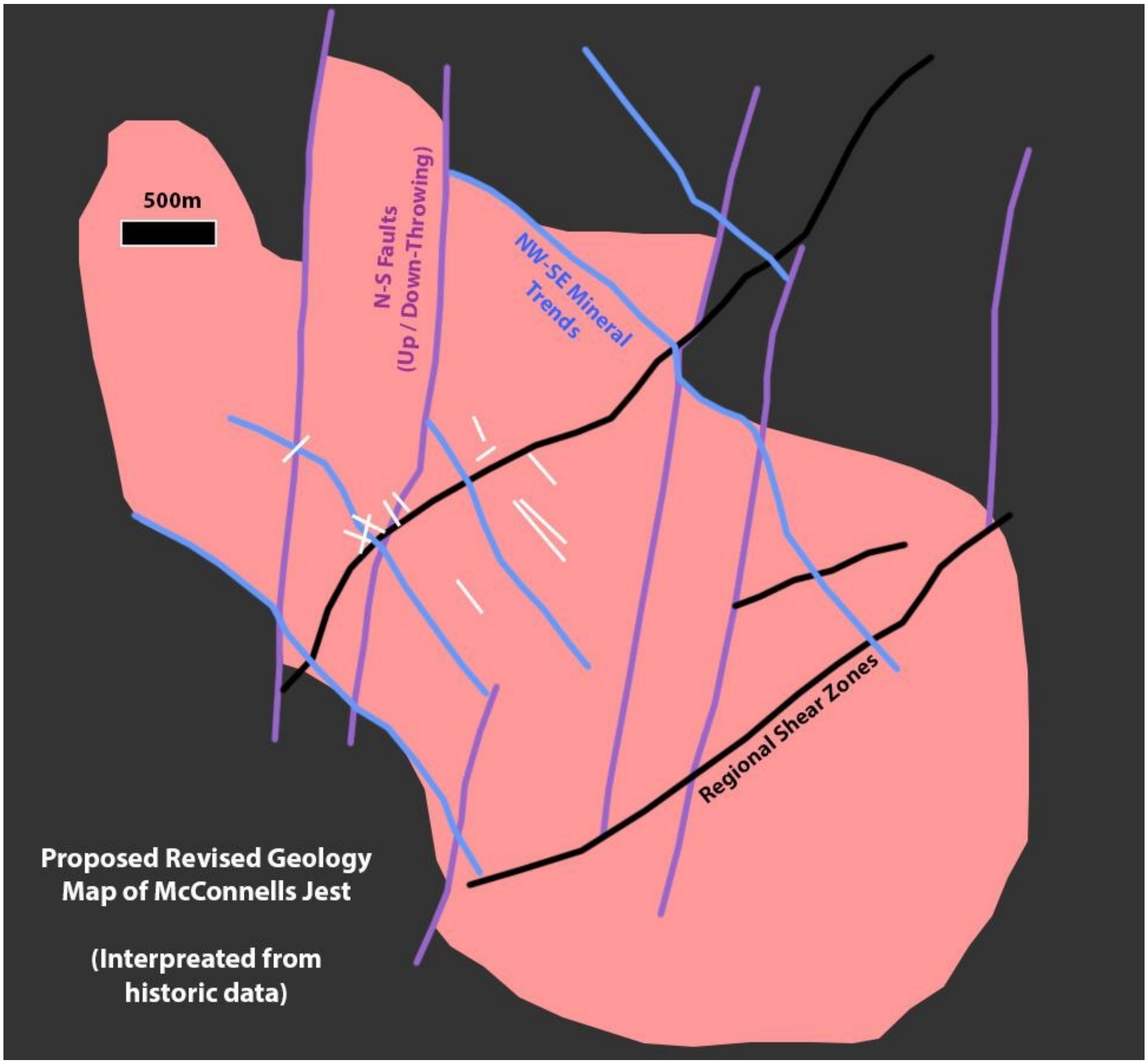

Figure 10.6 *The proposed revised geologic map of McConnells Jest based on structural interpretation of the surrounding geology.*

10.5 Interpretation

From this gathering and consolidation of data, there are several similarities structurally, geologically and geochemically to Dublin Gulch.

It has to be remembered that the Dublin Gulch discovery was overlooked for sometime due to the lack of a surficial gold anomaly. The lack of a strong gold-in-soil anomaly at McConnells Jest does therefore not preclude the existence of an ore body.

The proximity to Keno Hill may also mean that we see some silver - lead - zinc mineralisation on the property too. These can also be cooler, more distal systems to the pluton, but we cannot rule out some overprinting from the Keno-Elsa corridor.

When a map of McConnells Jest and Dublin Gulch are placed side by side (Fig. 10.7), there is a clear similarity in structural regime, although the McConnells Jest pluton has around double the surface area.

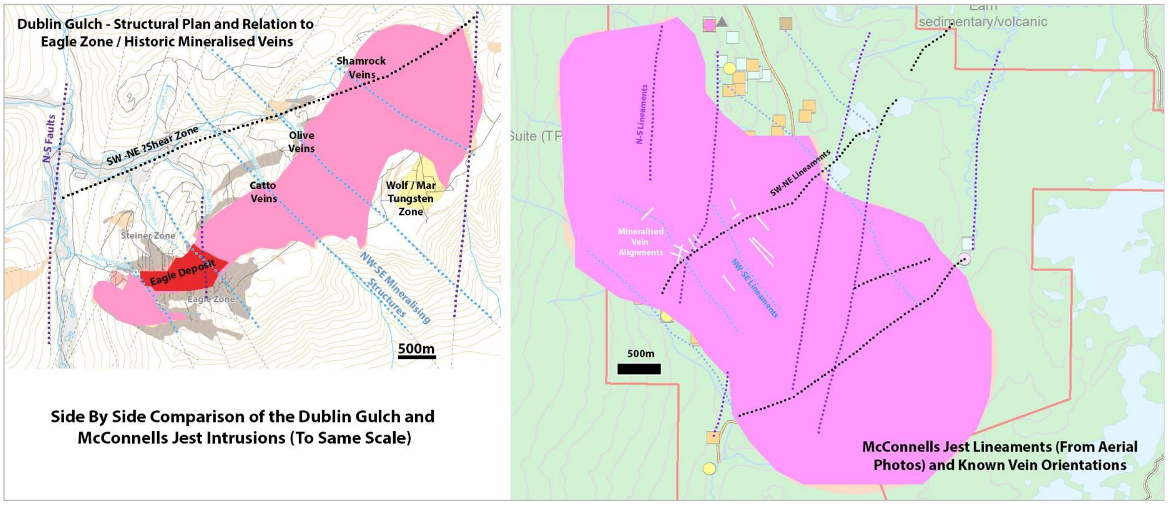

Figure 10.7 *Comparison of structural lineaments at Dublin Gulch (left) and McConnells Jest (right). Both plutons share the same three groups of structural alignments N-S (Purple), SE-NW (Light Blue) and SW-NE (Black). The images are drawn to the same scale and show that the exposed surface of the McConnell pluton is significantly larger than that at Dublin Gulch.*

Such Ag-Pb-Zn veins also occur at Dublin Gulch, in particular their Olive and Shamrock Zones, located about 1km east from the main deposit. It is hypothesized that these veins formed in the crystalline carapace of the pluton, where late stage gases and fluids deposited low-temperature minerals into shrinkage fracturing. On the surface, these are expressed as outcroppings of scorodite, a mineral that occurs from the weathering of arsenic-rich minerals. Historically, these veins were chased underground in small mining operations and were quite productive, if not short lived.

McConnells Jest has both the sheeted vein systems (Pink Mountain) and scorodite outcrops (Bullion Blister). The presence of till cover has somewhat impeded more extensive mapping and collection methods, which is something that should be addressed in the future.

11.0 Social Licence

11.1 First Nations

The property is located within the traditional territory of the Nacho Nyak Dun First Nations. The nearest settlement land or R-block is R-05A on Davidson Range, on the eastern edge of the property and east of McQuesten Lake. Two other nearby R-blocks are R-09B and A-07A, which are located northeast and northwest of the property, respectively.

11.2 Environmental Issues

The surrounding region of the property, which itself is located on the border between the administrative boundaries of 105M and 106D (Fig 11.1), is home to wildlife that range from being very common throughout the Yukon to being only endemic within the territory's central eastern region. The administrative boundaries of 105M and 106D are home to about 43 species of mammals, over 100 species of birds, 38 species of butterflies, 14 species of fish, the common wood frog and the less common boreal snaketail dragonfly.

The nearest significant wildlife key area is located just north of the property and is a summer nesting area shared by four species of raptors: peregrine falcon (*Falco peregrinus*), osprey (*Pandion haliaetus*), golden eagle (*Aquila chrysaetos*) and bald eagle (*Haliaeetus leucocephalus*). This area in turn overlaps a waterfowl breeding area to the east (Fig. 11.1).

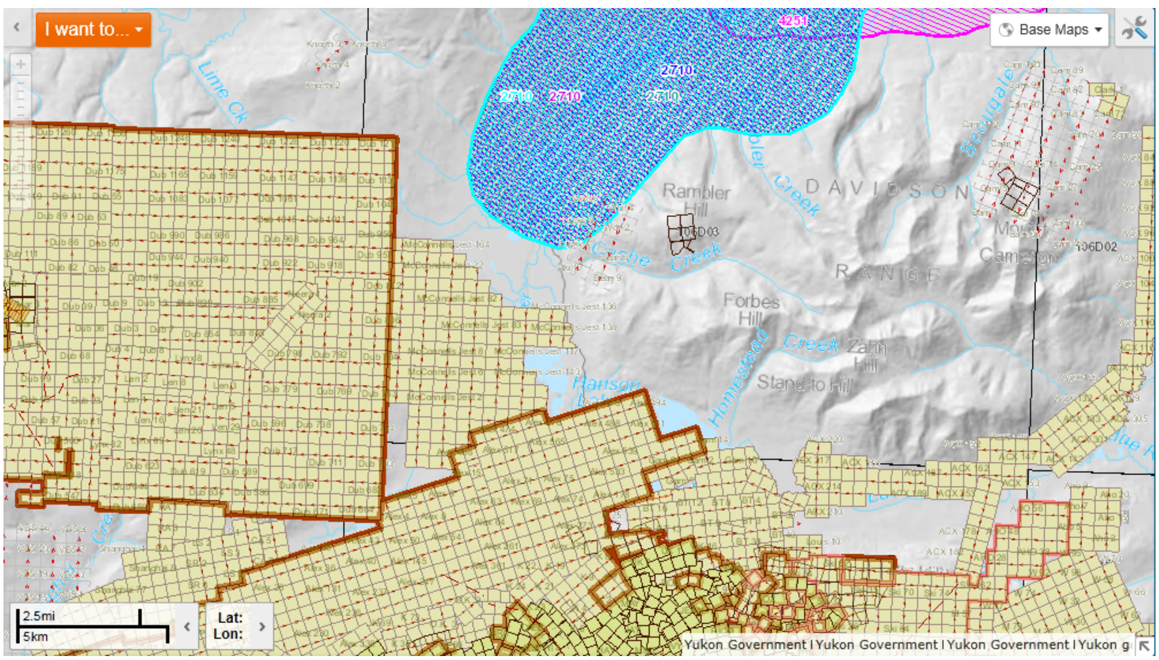

Figure 11.1 *Map showing nesting areas of local birds of prey (blue) and waterfowl (pink) north of the property.*

Aside from animals, the area is also home to rare vascular plant species endemic to the Yukon's Central Territory. A 2014 edition guide with descriptions of 39 of the rarest plants in the Central Territory of the Yukon can be found on the Yukon Conservation Data Centre website (Yukon Conservation Data Centre, 2014).

Given the concentration of mines and related development within the vicinity of the property, there does not appear to be any urgent conservation issues related to this area. There is however hunting and trapping activities present in the vicinity, as there is an abundance of small game such as weasels, waterfowl and grouse.

11.3 Local Populations

The capital city of Yukon Territory, Whitehorse is located ~350km south of the McConnells Jest property. According to the 2011 National Household Survey (NHS) the population of Whitehorse is 22,815 where 19,040 have a non-Aboriginal identity. Amongst the 19,040 people 17,130 have European origins and 1,905 have Asian origins. Filipino (705) and Chinese (535) make up the largest portions of the Asian visible minority. Other notable visible minorities include African (145) and Latin American (125) (Statistics Canada, 2013).

Local communities within ~100km of the McConnells Jest Property include: Elsa (~8km SW), Keno City (~10km SE), Mayo (~45km SW) and Stewarts Crossing (~90km SW).

According to the 2011 National Census, amongst the local communities, Mayo is largest with a population of 226, followed by Keno Hill (Keno City) with a population of 28 and Stewarts Crossing with a population of 25 (Statistics Canada, 2012c).

Elsa is considered a ghost town as its population moved out following the closure of the United Keno Hill mine in 1989.

According to Statistics Canada, 2011 NHS data for Mayo, Keno City and Stewarts Crossing has been suppressed for data quality or confidentiality reasons (Statistics Canada 2012a,b and d). Due to this, ethnicities for these communities is not public data.

12.0 Recommendations

Essentially this project is at a crucial point. The work so far has proven several similarities to Dublin Gulch, but the economic downturn has detracted potential investors from a property that shows sporadic anomalous assays. Two main factors have to be remembered here:

- The property is surrounded by mines that have successfully operated for many years, with new sites coming online in the next decade. These are not distal properties, but border the property and have geological continuity.

- The Eagle Zone at Dublin Gulch contains a global resource of 6.3 MOz of gold, yet has little to no geochemical expression on surface.

The next steps at McConnells Jest would follow these points:

- **Basic Mapping:** The property needs to have a geological map produced, showing the rock types, alteration, structures and mineralisation. Although this is somewhat difficult due to the glacial till cover, there is enough exposure on the high points to get a large portion of the property mapped.

- **Scorodite Prospecting:** If the property follows the Intrusion Related Gold model, then gold was delivered in late stage quartz or arsenopyrite veins which will be spaced 20-50m apart. As Scorodite is easy to identify on surface, prospecting this will reveal not only the grades, but also structural alignments. These can then be traced out over the entire property to identify other areas of interest.

- **Sheeted Vein Prospecting:** Typically, a series of aligned quartz veins host a significant portion of the gold in these systems. Mapping of the area, particularly in the northern and southern shoulders to locate areas of the highest density of veining. Once these are identified, these would produce the first exploration targets for future drill programs.

- **Identify Calcareous Areas of the Earn Group:** Looking for these horizons in the country rock could show evidence of skarn mineralisation. On the south-eastern side of the Dublin Gulch intrusion, contact with such layers produced a large tungsten deposit. A National Instrument 43-101 report authored by SRK Consulting (2008) stated that the Mar-Tungsten deposit contains 65.7 million pounds of tungsten in the Indicated category, and an additional 8.5 million pounds Inferred. This has not been studied at McConnells Jest, and doing so would complete the picture according the regional mineralisation models.

- The following map (Fig. 12.1) shows **several target areas** that are favourable for a variety of reasons. The most compelling are those that are in intersectional areas of the lineaments.

Much of this work would be undertaken by geologists in the field, collecting more information and samples from specific outcrops. Knowledge of the Dublin Gulch system will be an asset when looking at this ground.

Methods for sampling till have come a long way in recent years, especially on the geochemical front, with several laboratories offering specialized services. It could be proposed that additional soil work be undertaken in the southern and eastern portions, but only when regional trends had been identified from mapping, thus giving confidence in these potential extensions.

Access to the bedrock could also be achieved either by trenching or drilling. In both instances, it would be recommended to look at mobile, heliportable apparatus as the property is not directly accessible by ground at this stage. Low impact excavations also fit within existing exploration permits and demonstrate environmental stewardship.

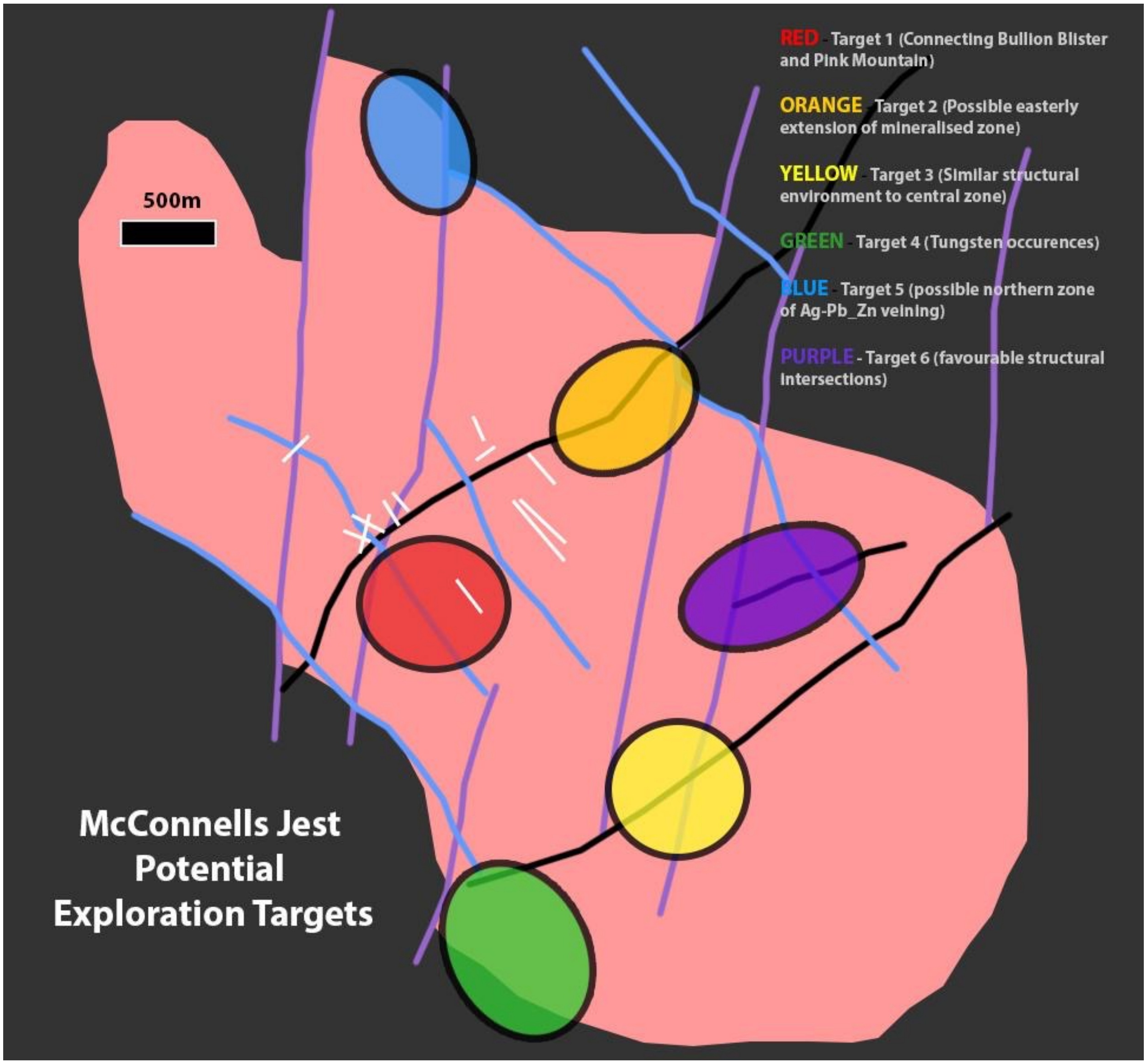

Figure 12.1 *Exploration targets generated from the geologic interpretations of the McConnells Jest property above. A total of 6 targets are proposed based on geochemical and structural data.*

In addition to geological work, it would be prudent at this stage to include some level of environmental assessment projects. Before any ground disturbance occurs through drilling or trenching, it is recommended to take water samples from the streams to analyze for basic composition and dissolved mineral content. This should be prioritized in the areas directly downstream from the scorodite veins that will likely be loading the water with arsenic, lead and antimony. Having these results acts almost as an 'insurance' policy to prove that any elevated levels measured in the future are not a direct consequence of exploratory work. Water sampling should be planned out in advance, as the samples need to reach a lab within 24 hours of being collected.

Other general environmental studies should be considered also: vegetation mapping, stream delineation (permanent waterways versus ephemeral streams as these will define protective buffers for exploration activity), and also some basic wildlife / plant surveys.

All of this work could be accomplished in a single field season with an experienced team of geologists / geotechnicians. Depending on funding available, the team could either fly camp on the site, or fly in from accommodations in Keno, Mayo or even the camp at Dublin Gulch.

The product of this work program would be to define structures and confirm the geological model. This would generate drill targets for future years and generate further investment interest.

13.0 References

Abbott, J.G., Gordey, S.P., and Templeman-Kluit, D.J., 1986, Setting of stratiform, sediment hosted lead-zinc deposits in Yukon and northeastern British Columbia, In: Morin, J.A., (Ed.), Mineral deposits of the northern Cordillera: Canadian Institute of Mining and Metallurgy Special Volume 37, p. 1–18.

Alexco Resources Corp., 2012, Preliminary Reclamation and Closure Plan, Keno District Mine Operations, Keno Hill Silver District, Revision 2 (internal report for Alexco Resources Corp. September 30, 2012), 306 p.

Bond, J.D, 1999, Glacial Limits and Ice Flow Patterns, Mayo Area, Central Yukon, Geoscience Map 1999-13: Exploration and Geological Services Division, Yukon Region, Indian and Northern Affairs Canada, 1:250,000 scale.

Cathro, R.J., 2006, The history and geology of the Keno Hill Silver Camp, Yukon Territory: Geoscience Canada, v. 33, p. 103–134.

Golden Predator Canada Corp., 2011, Assessment Report, 2010 Sampling Program, McConnell Property, Mayo Mining Division, Yukon Canada, internal report prepared for Bill Koe-Carson dated May 3rd 2011.

Golden Predator Canada Corp., 2013, Assessment Report, 2012 Sampling Program, McConnell Property, Mayo Mining Division, Yukon, Canada, internal report prepared for Bill Koe-Carson dated April 4th 2013.

Gordey, S.P., and Makepeace, A.J., 2003, Yukon Digital Geology, version 2.0, S.P., Gordey, and A.J., Makepeace (comp); Geological Survey of Canada, Open File 1749 and Yukon Geological Survey, Open File 2003-9 (D)

Gordey, S.P., and Anderson, R.G., 1993, Evolution of the northern Cordilleran miogeocline, Nahanni map area (105I), Yukon and Northwest Territories: Geological Survey of Canada, Memoir 248, p. 214.

Green, L.H., 1972, Geology of the Nash Creek, Larsen Creek and Dawson Map areas, Yukon Territory: Geological Survey of Canada, Memoir 364, 157 p.

Hammer, Ø, Harper, D.A.T., and Ryan, P.D., 2001, Past: Paleontological Statistics Software Package for Education and Data Analysis. Palaeontologia Electronica, vol. 4, issue 1, p. 4-9, 178kb. http://palaeo-electronica.org/2001_1/past/issue1_01.htm.

Hantelmann, J.J., 2013, The paragenesis and geochemistry of the Bellekeno Ag-Pb-Zn vein, Keno Hill district, Yukon, Canada: Unpublished M.Sc. thesis, University of Alberta, Edmonton, Canada, 292 p.

Hart, C.J.R, 2004, Mid-Cretaceous magmatic evolution and intrusion-related metallogeny of the Tintina Gold Province, Yukon and Alaska: Unpublished PhD. thesis, University of Western Australia, Perth, 188p.

Hart, C.J.R., 2005, Classifying, distinguishing and exploring for intrusion-related gold systems: The Gangue: Newsletter of the Geological Association of Canada Mineral Deposits Division, v. 87, p. 1, 4–9.

Hart, C.J.R., 2007, Reduced intrusion-related gold systems, In: Goodfellow, W.D. (Ed.), Mineral deposits of Canada: A synthesis of major deposit types, district metallogeny, the evolution of geological provinces, and exploration methods: Geological Association of Canada, Mineral Deposits Division, Special Publication No. 5, p. 95-112.

Hart, C.J.R., Mair, J.L., Goldfarb, R.J., and Groves, D.I., 2004, Source and redox controls on metallogenic variations in intrusion-related ore systems, Tombstone-Tungsten belt, Yukon Territory, Canada: Transactions of the Royal Society of Edinburgh, Earth Science, v. 95, p. 319−337.

Hitchins, A.C., and Orssich, C.N., 1995, The Eagle zone gold-tungsten sheeted vein porphyry deposit and related mineralization, Dublin Gulch, Yukon Territory: Canadian Institute of Mining and Metallurgy, Special Volume 46, p. 803–810.

Koe-Carson, B., 2010, McConnells Jest, 2010 Field Journal Notes, Final Report, Yukon Mining Incentives Program #10-001, 9 p.

Lang, J.R. and Baker, T., 2001, Intrusion-related gold systems: the present level of understanding: Mineralium Deposita, v. 36, p. 477-489.

Mair, J.L., Hart, C.J.R., and Stephens, J.R., 2006, Deformation history of the western Selwyn Basin, Canada: Implications for orogen evolution and the setting of mid-Cretaceous magmatism: Geological Society of America Bulletin, v. 118, p. 304−323.

Mair, J.L., Farmer, G.L., Groves, D.I., Hart, C.J.R and Goldfarb, R.J., 2011, Petrogenesis of Postcollisional Magmatism at Scheelite Dome, Yukon, Canada: Evidence for a Lithospheric Mantle Source for Magmas Associated with Intrusion-Related Gold Systems, Economic Geology, v. 106, p. 451–480.

MinFile#: 106D 055, 2008, Yukon MinFile for the "Zed" occurrence, Yukon Geological Survey.

Murphy, D.C., 1997, Geology of the McQuesten River region, northern McQuesten and Mayo Map Area, Yukon Territory (115P/14, 15, 16; 105M/13, 14): Exploration and Geological Services Division, Yukon, Indian and Northern Affairs Canada Bulletin 6.

Pearson, K., 1896, Mathematical contributions to the theory of evolution. III. Regression, heredity, and panmixia: Philosophical Transactions of the Royal Society A, v. 187, p. 253–318.

Rasmussen, K.L., 2013, The timing, composition, and petrogenesis of syn- to postaccretionary magmatism in the northern Cordilleran miogeocline, eastern Yukon and southwestern Northwest Territories: Unpublished Ph.D. Thesis, University of British Columbia, Vancouver, Canada, 810 p.

Roots, C.F., 1997, Geology of the Mayo Map area, Yukon Territory (105M): Exploration and Geological Services Division, Yukon, Indian and Northern Affairs Canada, Bulletin 7, 82 p.

Scott Wilson Mining, 2010, Technical Report on the Dublin Gulch Property, Yukon Territory, Canada: NI 43-101 technical report (internal report for Victoria Gold Corp., April 23, 2010), 242 p.

Selby, D., Creaser, R.A., Heaman, L.M., and Hart, C.J.R., 2003, Re–Os and U–Pb geochronology of the Clear Creek, Dublin Gulch, and Mactung deposits, Tombstone Gold Belt, Yukon, Canada: absolute timing relationships between plutonism and mineralization, Canadian Journal of Earth Science, v. 40, p. 1839–1852.

Spearman, C.E., 1904, The proof and measurement of association between two things: American Journal of Psychology, v. 15, p. 72-101.

SRK Consulting, 2008, Dublin Gulch Property – Mar-Tungsten zone Mayo District, Yukon Territory, Canada: NI 43-101 technical report, (internal report for StrataGold Corp., Dec. 1 2008), 161 p.

SRK Consulting, 2014, Updated Preliminary Economic Assessment for the Keno Hill Silver District Project – Phase 2, Yukon, Canada: NI 43-101 technical report (internal report for Alexco Resources Corp., December 10, 2014), 381 p.

Statistics Canada, 2012a, Keno Hill, Yukon (Code 6001052) and Yukon (Code 60) (table), Census Profile, 2011 Census: Statistics Canada Catalogue no. 98-316-XWE, Ottawa, Released October 24, 2012.

Statistics Canada, 2012b, Mayo, Yukon (Code 6001022) and Yukon (Code 60) (table), Census Profile, 2011 Census: Statistics Canada Catalogue no. 98-316-XWE, Ottawa, Released October 24, 2012.

Statistics Canada, 2012c, Population and dwelling counts, for Canada, provinces and territories, and census subdivisions (municipalities), 2011 and 2006 censuses (table), Population and Dwelling Count Highlight Tables. 2011 Census: Statistics Canada Catalogue no. 98-310-XWE2011002, released February 8, 2012.

Statistics Canada, 2012d, Stewart Crossing, Yukon (Code 6001050) and Yukon (Code 60) (table), Census Profile, 2011 Census: Statistics Canada Catalogue no. 98-316-XWE, Ottawa, Released October 24, 2012.

Statistics Canada, 2013, Whitehorse, CY, Yukon (Code 6001009) (table), National Household Survey (NHS) Profile, 2011 National Household Survey: Statistics Canada Catalogue no. 99-004-XWE, Ottawa, Released September 11, 2013.

Wardrop Engineering Inc., 2011, Technical report on the Eagle zone, Dublin Gulch property, Yukon Territory, Canada: NI 43-101 technical report, (internal report for Victoria Gold Corp., April 15 2011), 48 p.

Wardrop Engineering Inc., 2012, Technical report on the Eagle zone, Dublin Gulch property, Yukon Territory, Canada: NI 43-101 technical report, (internal report for Victoria Gold Corp., April 05 2012), 419 p.

Yukon Conservation Data Centre, 2014, Rare species database: Yukon Department of Environment, Whitehorse, Yukon, www.env.gov.yk.ca/cdc, accessed April 8, 2015.

Appendix 1: Team Details and Qualifications

Andrew Randell (PGeo)

Andy is a Professional Geoscientist (PGeo) registered with the Association of Professional Engineers and Geoscientists in British Columbia (APEGBC).

He has over a decade of experience across a multitude of mineral exploration and mining project types and commodities, specializing in geological mapping, design of exploration projects and drilling programs, geochemical targets, soil sampling, mineralogy and metallurgy.

Fraser Kirk

Fraser Kirk is a geologist and research scientist. He holds a B.Sc. (Hons) in Geoscience from the University of St Andrews, Scotland and is currently completing a M.Sc. in Geochemistry and Economic Geology at Memorial University, NL, Canada. His research focuses on gold mineralisation at the Dublin Gulch property, YT, Canada; using mineralogical studies, geochemistry, stable and radiogenic isotopes, and U-Pb dating to determine the genesis and distribution of gold there. Fraser seeks to integrate a traditional 'boots on the ground' approach with modern geochemical techniques to accurately and efficiently solve geologic problems. In addition, Fraser has a keen interest in quality systems, particularly the TQM model and its application to mineral exploration.

Hanson Wong

Hanson is an avid volunteer at the University of British Columbia's Pacific Museum of the Earth, where he regularly conducts earth science related workshops and museum tours for school groups visiting the Museum. He holds a B.Sc. in Earth Sciences from Simon Fraser University and has a keen interest in paleontology. He has experience doing fieldwork in northeastern British Columbia for a Junior coal exploration company and a paleontological consultant company.

Carrie Wong

Carrie Wong is a GIT with three years good standing at APEGBC. She graduated with honors from the University of British Columbia in 2011. She has worked in the gold exploration industry and spent a few summer seasons in the Yukon. She has created hundreds of maps over the years. From acQuire database management to logging core with world experts (like Dr. Richard Sillitoe within a porphyry system), Carrie has had a number of diverse roles within the mining industry. In response to the mining recession, she has launched her own cartography business and is looking for clients. She has also written articles for www.geologyforinvestors.com. For more information: www.wongmanagement.ca

Daniela Marcoux

Daniela holds an Honours BSc in Geology from the University of Ottawa. A registered G.I.T. and a true lover of her homeland, her experience is mainly based in the foundation of it all, the Canadian shield. Amongst her experience she has worked on the Borden Gold Project, recently acquired by Goldcorp. Her geological interests include: any project that can stimulate the country's economy, mineral deposits in greenstone provinces and cold beer.