

INTRODUCTION

The Tosingermann Lakes map area straddles the northwestern extent of the Ruby and Nisling Ranges. Broad treeless uplands dominate the landscape and summits reach more than 1800 m a.s.l. Upland surfaces are covered with weathered bedrock colluvium and sporadically distributed till deposits from former alpine glaciers and ice sheets. Upland surfaces are affected by active periglacial processes such as cryoturbation, inversion and solifluction. Cryoturbation terraces were documented on unglaciated upland surfaces in the Nisling Range (Fig. 1). The valleys occupied by the Klauene River and Tosingermann and Tincup lakes are steep sided and narrow. Many of these slopes are over-steepened and susceptible to rockslides, such as is documented at the south end of Tincup Lake and near Dog Park Lake just south of the map area. The valleys within the Nisling Range have a more mature or broad appearance. The Grace Lake valley that cuts across the northern part of the map area, may represent the paleo-drainage course of the Donjek River. The valleys are generally poorly drained due to the thick deposits of glacial sediment and widespread permafrost.

GLACIAL HISTORY

At least three ice sheets originating from the St. Elias Mountains have advanced into the Ruby and Nisling ranges during the Quaternary period. The oldest advance predates marine isotope stage (MIS) 4 (50ka BP to 80ka BP) and may correlate with the MIS 6 (130ka BP to 230ka BP) Reid glaciation documented for the Selwyn Lake (Ward et al., 2008). Mapping the Reid glacial limit in the Tosingermann Lakes map area was not possible since the younger Gladstone glaciation reached a similar extent up to 1200 - 1400 m and therefore masks the older features. The early Wisconsin Gladstone glaciation reached its maximum extent and had begun to retreat by 50,000 years ago (Ward et al., 2007). Deposits from this glaciation are preserved above and beyond the late Wisconsin McConnell limit, which reached up to 1180 m. Evidence of the Gladstone glaciation consists of moraine deposits and ice proximal features such as meltwater channels.

Most of the glacial deposits in the map area are attributed to the late Wisconsin McConnell glaciation. According to cosmogenic dates on erratics from the glacial limit immediately to the southwest of the map area, the McConnell glaciation was positioned at its maximum extent between 13,740 ± 500 years B.P. and 14,200 ± 800 years B.P. The style of glaciation in the map area could be characterized as a network of coalescing valley glaciers. The thickest ice penetrated the map area from the west by the Donjek and Klauene river valleys. This "Donjek glacier" followed the Grace Lake valley to the northeast and advanced southeastward up Tincup Creek valley. Here the Donjek glacier converged with a westward-flowing glacier occupying upper Orion Creek valley. Abundant alluvial moraine and flimsy of moraine ridges in the Orion and Grace lake valleys suggest the McConnell glaciers in the map area were active during deglaciation (Fig. 2). Thick glaciofluvial deposits in the Tincup Creek and Grace Lake valleys indicate that glacial lakes occupied the valleys as the ice retreated (Fig. 3). Small McConnell and Gladstone age glaciers formed on the upland separating Tosingermann and Tincup lakes and in one cirque to the east of Tincup Lake. These glaciers likely did not coalesce with the large valley glaciers originating from the St. Elias Mountains.

PERMAFROST

Most of the map area is underlain by permafrost. The nature of the ice within the sediments is generally a function of drainage. Thermokarst ponds are present within glaciofluvial sediments in the Grace Lake and Orion Creek valleys, which indicates that massive ice bodies are present (Fig. 3). Well-drained glaciofluvial deposits have a thicker active layer, but likely still contain ground ice at depth. The large glaciofluvial plain north of Tincup Lake supports sparse vegetation and appears to have a high water table. This may be attributed to a shallow permafrost table and poorly drained glaciofluvial sediments at depth.

WHITE RIVER TEPHRA

Significant thicknesses of White River tephra can be found in the Tosingermann Lakes map area (Fig. 4). The thickest deposits occur near Tosingermann Lakes and into the Grace Lake valley. The primary tephra accumulation, which has a medium sand-sized texture, is estimated to be 10-20 cm thick on average throughout the map area. Much greater thicknesses accumulated in valley bottoms when the tephra was eroded and re-accumulated off the hill tops. Thicknesses exceeding 20 cm are common on alluvial fans and deltas (Fig. 4). At Tosingermann Lakes, the cased of the northern lake became plugged with ash accumulating on a fluvial fan. As a result, the level of the northern lake rose by over 2 m until it was able to overflow the fan. Since that time,

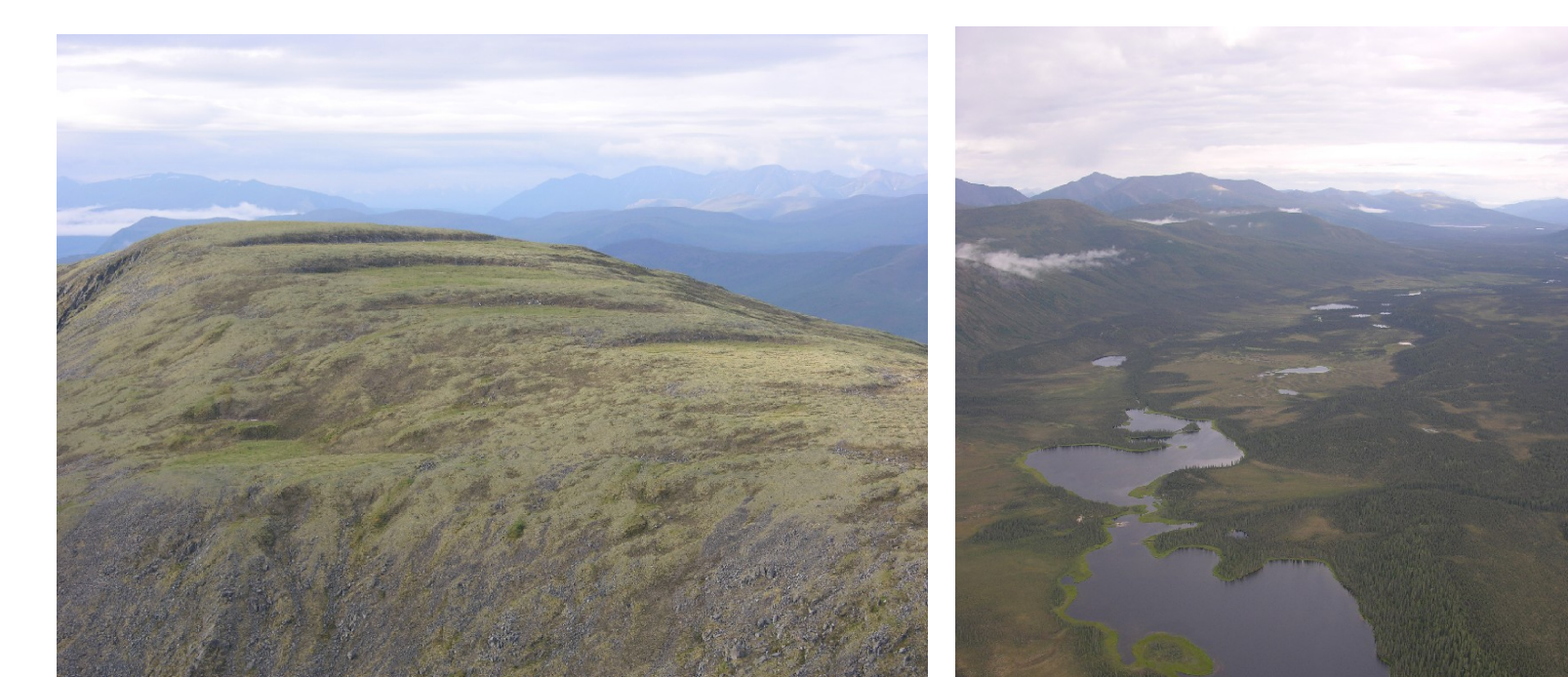


Figure 1. A view to the south of the cryoplanation terraces in the northeast part of the map area. A muddy colluvial dambank crosses the upland and bedrock outcrops in the scene. These landforms develop over hundreds of thousands of years through periglacial weathering and sediment creep. The terrace slopes are 8 to 10 m in height (06B035).

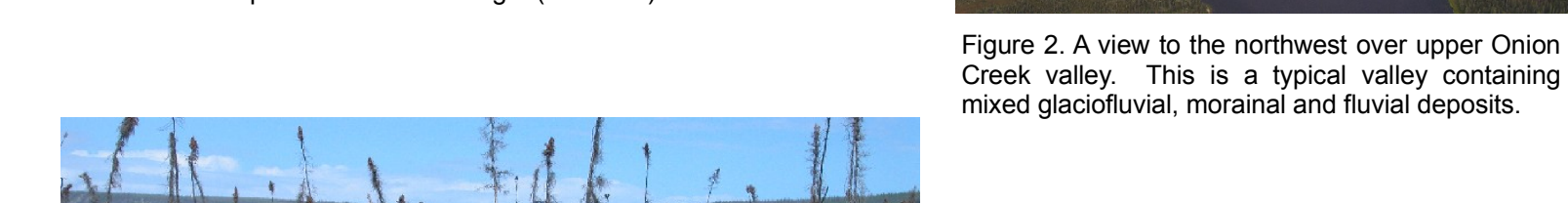


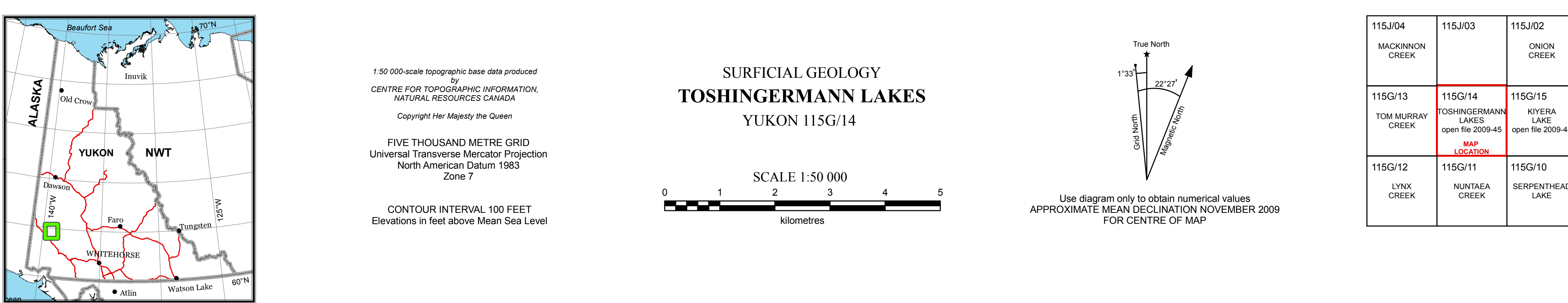
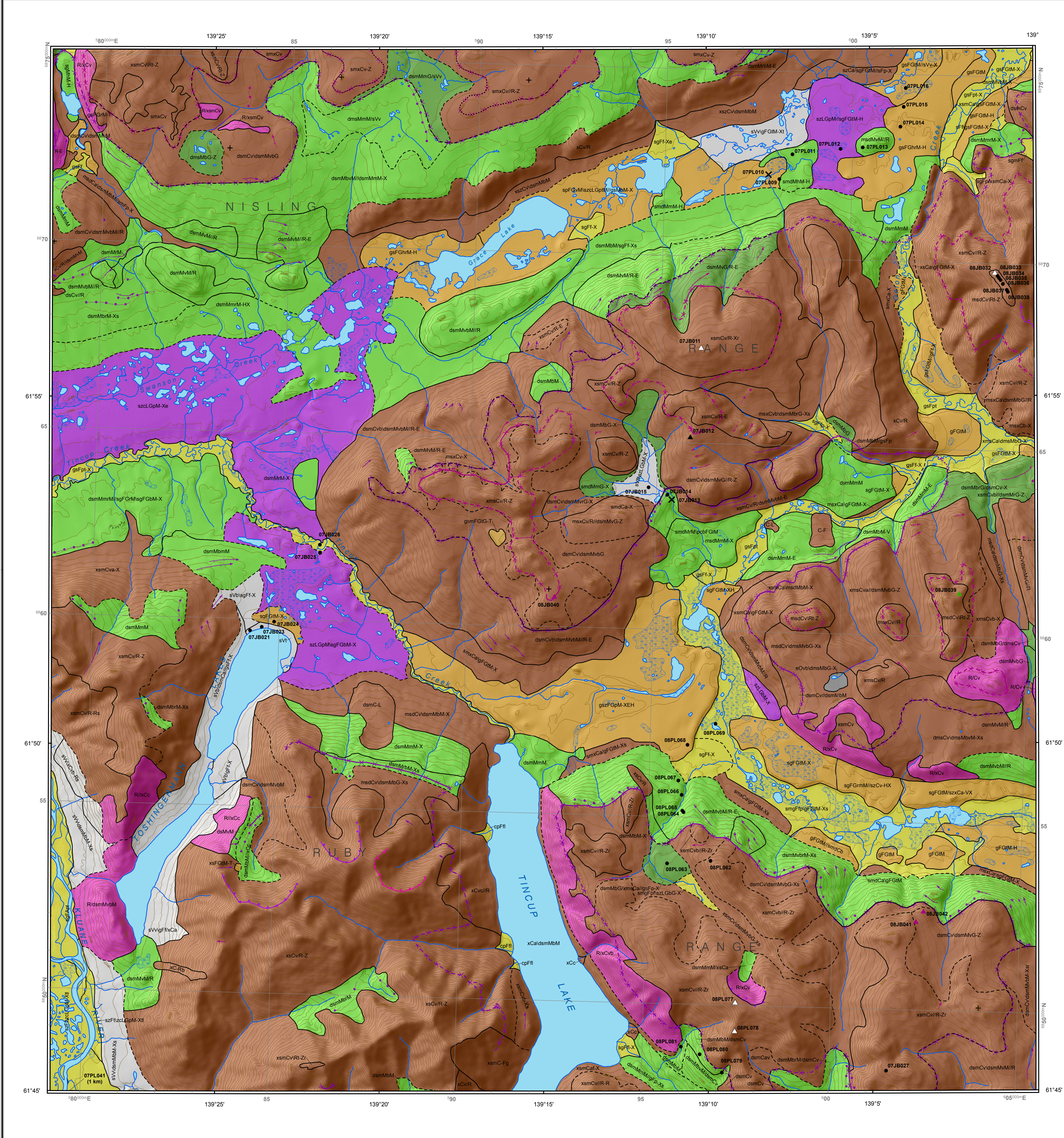
Figure 2. A view to the northwest over upper Orion Creek valley. This is a typical valley containing mixed glaciofluvial, moraine and fluvial deposits.



Figure 4. Ice-rich glaciofluvial sediments exposed in a thermokarst pond in Grace Lake valley (07PLD10).



Figure 3. White River tephra accumulation at the north end of Tosingermann Lakes (07B021). The grey band separates primary tephra accumulations (below band) from secondary reworked tephra accumulations.



SURFICIAL MATERIALS

Surficial materials are non-lithified, unconsolidated sediments. They are produced by weathering, sediment deposition, biological accumulation, human and volcanic activity. In general, surficial materials are of relatively young glacial age and they constitute the parent material of most (pedological) soils. On the map, surficial materials from the core of the polygon label, they are symbolized with a single upper case letter, with texture expression or glacial qualifier to the right. The glacial qualifier "G" was used to describe glacially modified materials. If actual activity state is different than the assumed activity state indicated in brackets next to the surficial material name (below), a qualifier A (active) or I (inactive) must be used as a superscript following the surficial material designator. Note that a single polygon will be coloured only by the dominant surficial material, but other materials may exist in that unit.

- D** - **HOLOCENE**
Organic: Organic deposits are accumulations of vegetative matter thicker than 1 m. They are commonly found in floodplains; areas of near-surface permafrost such as north-facing slopes, and locations where there is poor drainage. Thin veneers of organic material are widespread and often unmapped.
- V** - **Volcanic**: Volcanic tephra deposits found in the map area are from the 1140-year-old White River eruption. Primary tephra deposition across the map area was between 10 cm and 20cm and consisted of a grain size of medium-sand. Local reworking of the tephra into lake basins, onto fluvial fans or into cirque basins can result in accumulations exceeding 20cm.
- E** - **Eolian**: Material that was transported and directly deposited by wind. The dominant eolian sediment in the map area is loess. A thin veneer of loess (10-20 cm) was deposited over the landscape during the last glaciation. On stable sites, the loess is intact, whereas at cryoturbated or colluviated areas, the loess is revealed into the soil profile. Since loess deposits represent only a thin veneer, they were not mapped; however, loess is a widespread material in the map area.
- C** - **Colluvium**: Material that was transported and directly deposited by down-slope gravity-driven processes such as creep, landslides and snow avalanches. Due to the active periglacial processes that have occurred in the map area, colluvium is widespread across the map area. The texture and composition of colluvium vary more than any other material in the map area, depending on the parent material, and the mechanism and distance transported. For example, materials derived from till will likely resemble the in situ till, however they may contain slope-parallel reorientation structures. Comparatively, colluvium derived from physically weathered bedrock will be an angular, poorly sorted diamicton and will likely contain less matrix. Some materials formed by rapid processes, such as rock falls, debris flows and avalanches, are cryoturbated and colluvially buried, and will likely contain less matrix. Conversely, slower processes occur on gentle slopes and are commonly associated with permafrost, solifluction and creep. Colluvium is commonly derived from weathered bedrock and till, resulting in a silt-rich diamicton with angular, local bedrock and sub-rounded erratic clasts. Beyond the glacial limits and in alpine settings, colluvium is dominantly derived from weathered bedrock fragments and loess.
- F** - **Fluvial**: Fluvial materials are transported and deposited by modern streams and rivers. They typically consist of stratified sand and gravel that is well sorted and contains sub-angular to rounded clasts. These deposits result in floodplain, terrace and fan surface expressions within the map area. Due to scale limitations, fluvial deposits in most of the smaller valleys are not mapped.
- L** - **Lacustrine**: Sediment that has been deposited into a modern lake. Includes biologically produced material such as mat and gyttja. Lacustrine deposits are only mapped where a lake has drained exposing the lake bottom material.

Glaciofluvial: Glaciofluvial materials have been deposited directly by glacial meltwater. These deposits can form above, in, below or adjacent to a glacier. They are deposited in meltwater channels, eskers, terraces, fans and deltas. Glaciofluvial deposits consist of moderately to well-sorted, rounded, stratified sand and gravel, but can vary locally depending on transport distance. Near surface ground ice is generally absent in glaciofluvial deposits unless there is a poorly drained underlying unit present.

- FG** - **LATE WISCONSIN - MCCONNELL (M)**
Large glaciofluvial plains and terraces are found north of Tin Cup Lake.
- FG** - **EARLY WISCONSIN - GLADSTONE (G)**
Isolated pockets of Gladstone glaciofluvial deposits are found above the McConnell limit in the map area.
- IL** - **ILLINOIAN - REID (R)**
Isolated pockets of Reid glaciofluvial deposits may be found above the Gladstone limit in the map area.

Moraine: Moraine materials are diamictons deposited by either primary glacial processes such as lodgement, deformation and melt-out, or secondary glacial processes caused by gravity and water. Therefore, this term applies to all types of till including flow tills, which are not directly deposited by glacial ice. Abandon tills tend to have a hummocky or rolling surface morphology with a sandy matrix comprising 30-40% of the material. Lodgement tills have an even surface morphology with a silt and clay matrix comprising 40-50% of the material. Due to the uneven topography of the map area, tills are often colluviated. Permafrost is generally widespread within moraine deposits.

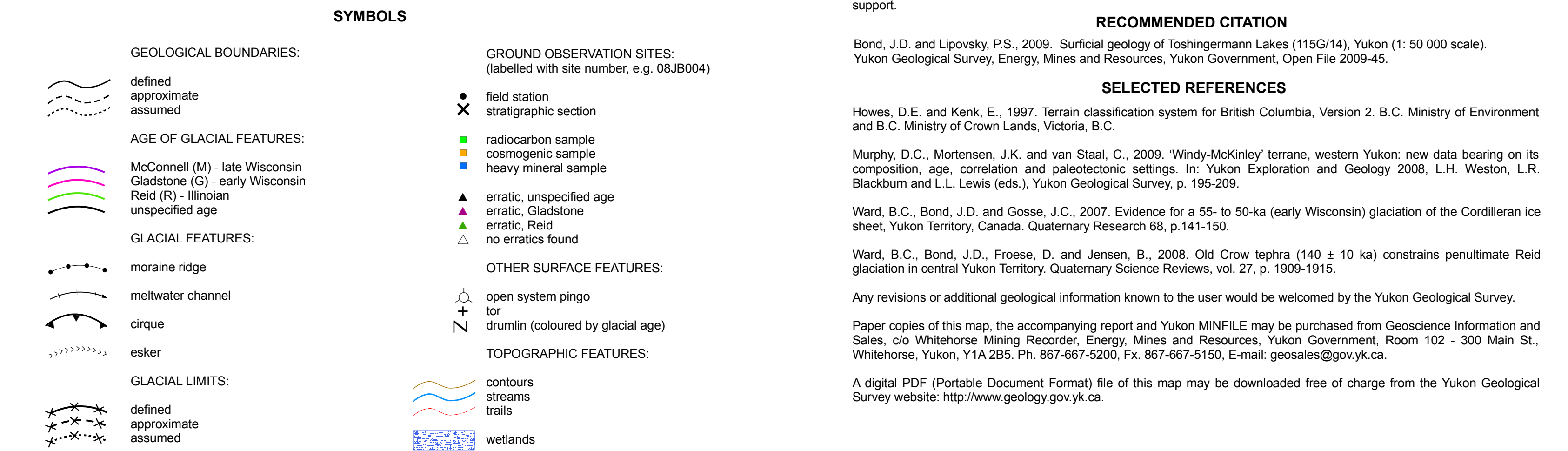
- M** - **LATE WISCONSIN - MCCONNELL (M)**
Both ablation and lodgement tills were observed and are common on the valley flanks throughout the map area.
- G** - **EARLY WISCONSIN - GLADSTONE (G)**
Both ablation and lodgement tills were observed and occur in isolated deposits immediately above, or beyond the McConnell limit.
- R** - **ILLINOIAN - REID (R)**
No Reid moraine deposits were mapped; however, isolated deposits may be found above the Gladstone limit.

Glacioclastic: Glacioclastic materials were deposited in a lake that formed on, in, under or beside a glacier. Glacioclastic sediments consist of stratified sand, silt and clay, loess-rich permafrost and thermokarst erosion is widespread in these deposits. Their poor drainage and high in situ moisture content can result in massive ice lenses.

- LG** - **LATE WISCONSIN - MCCONNELL (M)**
McConnell glacioclastic deposits are common in Tin Cup Creek, Grace Lake and the Klauene River.
- LG** - **EARLY WISCONSIN - GLADSTONE (G)**
Glacioclastic sediment of this age are rare and restricted to small tributary valleys above the McConnell limit. As a result, all of these deposits are now buried under colluvium or fluvial deposits.
- LR** - **ILLINOIAN - REID (R)**
Glacioclastic sediments of this age have not been documented in the map area.

Bedrock: Rocks in the Klauene Lake area are composed of three metamorphic assemblages, mid- and Late Cretaceous to Eocene granitoids and Upper Cretaceous to Eocene volcanic rocks. Late Cretaceous to the Ruby Range batholith, the largest pluton in the area. Metamorphic rocks to the northwest belong to Yukon-Tanana terrane and the structurally overlying Windy-McKinley terrane. Yukon-Tanana terrane consists of two assemblages, quartzite, amphibole, pelite, marble and amphibolite of the Proterozoic to Lower Paleozoic; Snowcap assemblage, and carbonaceous phyllite and quartzite and lesser felsic and mafic metavolcanic rocks of the mid-Paleozoic to Late Permian Finlayson assemblage. Foliated granitic rocks occur in both assemblages. Windy-McKinley terrane also comprises two assemblages: the schist-gabbro sub-division and the Harzburgtage Peak-Eklund Mountain ophiolite. The former assemblage is lithologically and stratigraphically similar to Yukon-Tanana terrane, differing only in the presence of voluminous bodies of Triassic metabasite. The Harzburgtage Peak-Eklund Mountain ophiolite in Klauene Lake area consists primarily of harzburgite, with lesser amounts of gabbro, dunite and pyroxenite. The third metamorphic assemblage, the Klauene schist, is southwest of the Ruby Range batholith. Klauene schist consists primarily of highly deformed, variably carbonaceous, porphyroblastic biotite schist and rare bodies of gabbro and harzburgite (Murphy et al., 2000).

- R** - **PRE-QUATERNARY**
Bedrock: Rocks in the Klauene Lake area are composed of three metamorphic assemblages, mid- and Late Cretaceous to Eocene granitoids and Upper Cretaceous to Eocene volcanic rocks. Late Cretaceous to the Ruby Range batholith, the largest pluton in the area. Metamorphic rocks to the northwest belong to Yukon-Tanana terrane and the structurally overlying Windy-McKinley terrane. Yukon-Tanana terrane consists of two assemblages, quartzite, amphibole, pelite, marble and amphibolite of the Proterozoic to Lower Paleozoic; Snowcap assemblage, and carbonaceous phyllite and quartzite and lesser felsic and mafic metavolcanic rocks of the mid-Paleozoic to Late Permian Finlayson assemblage. Foliated granitic rocks occur in both assemblages. Windy-McKinley terrane also comprises two assemblages: the schist-gabbro sub-division and the Harzburgtage Peak-Eklund Mountain ophiolite. The former assemblage is lithologically and stratigraphically similar to Yukon-Tanana terrane, differing only in the presence of voluminous bodies of Triassic metabasite. The Harzburgtage Peak-Eklund Mountain ophiolite in Klauene Lake area consists primarily of harzburgite, with lesser amounts of gabbro, dunite and pyroxenite. The third metamorphic assemblage, the Klauene schist, is southwest of the Ruby Range batholith. Klauene schist consists primarily of highly deformed, variably carbonaceous, porphyroblastic biotite schist and rare bodies of gabbro and harzburgite (Murphy et al., 2000).



SURFICIAL GEOLOGY TOSHINGERMANN LAKES YUKON 115G/14

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FIVE THOUSAND METRE GRID Universal Transverse Mercator Projection North American Datum 1983 Zone 7

CONTOUR INTERVAL: 100 FEET Elevations in feet above Mean Sea Level

SCALE 1:50,000

Use diagram only to obtain numerical values APPROXIMATE MEAN DECLINATION NOVEMBER 2009 FOR CENTRE OF MAP

115J04	115J03	115J02
MACKONNON CREEK	ONION CREEK	
115G13	115G14	115G15
TOM MURRAY CREEK	TOSHINGERMANN LAKES (open to 2004/15) WAP	WINDY LAKE (open to 2009/46)
115G12	115G11	115G10
LYNX CREEK	MUNTEA CREEK	SERPENTHEAD LAKE

RECOMMENDED CITATION
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SELECTED REFERENCES
Howes, D.E. and Kenk, E., 1997. Terrain classification system for British Columbia, Version 2. B.C. Ministry of Environment and B.C. Ministry of Crown Lands, Victoria, B.C.
Murphy, D.C., Mortenson, J.K. and van Staal, C., 2009. Windy-McKinley terrane, western Yukon: new data bearing on its composition, age, correlation and paleogeographic settings. In: Yukon Exploration and Geology 2008, L.H. Weston, L.R. Blackburn and L.L. Lewis (eds.), Yukon Geological Survey, p. 195-209.
Ward, B.C., Bond, J.D. and Gosse, J.C., 2007. Evidence for a 55- to 50-ka (early Wisconsin) glaciation of the Cordilleran ice sheet, Yukon Territory, Canada. Quaternary Research 68, p. 141-150.
Ward, B.C., Bond, J.D., Fosse, D. and Jensen, B., 2008. Old Crow tephra (140 ± 10 ka) constrains penultimate Reid glaciation in central Yukon Territory. Quaternary Science Reviews, vol. 27, p. 1909-1915.

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Yukon Geological Survey
Energy, Mines and Resources
Government of Yukon

Open File 2009-45
Surficial Geology of Tosingermann Lakes (NTS 115G/14)
Yukon
(1:50,000 scale)

by
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