

Fine- to medium-grained hornblende diorite sills and dykes. Medium-to-coarse grained hornblende quartz diorite.

(Carbonate rocks altered by intrusions are white, de-dolomitized)

Fine-grained, dark grey to black quartzite and interbedded black, carbonaceous phyllite

GOSSAGE FORMATION: alternating dark and light grey banded

Light grey weathering, thick-bedded and massive dolostone

Dark grey thin-bedded dolostone, with abundant siderite replacement along laminae and in fractures; in northeast includes tan- and green-weathering dolosiltstone: Possibly equivalent to

Dark-weathering, medium-grained white sandstone, thick bedded and locally hematized, with maroon siltstone interbeds: Probably equivalent to SLATS CREEK FORMATION of Norris (1982)

Age uncertain: Probably Late Proterozoic or Cambrian

Light grey, medium-bedded limestone

Stratigraphically equivalent (but no lithologic similarities) to PINGUICULA GROUP of Eisbacher (1981). Several rock units resemble Units 4,5 of Thompson and Roots (1982; now called Lower Fifteenmile

Undivided (includes one or more rock types below)

Grey and brown siltstone, dark grey fissile mudstone; occood conglomerate with quartz, plutonic and shale pebbles, arkose at

Volcanic breccia and conglomerate, mafic tuff

Pink-weathering, poorly bedded dolostone; 'tungussoid', finelaminated stromatolites (Mustard et al., 1990) diagnostic

Brown, fine-grained sandstone, thin-bedded

Dark grey siliceous siltstone at base, dark brown, fissile mudstone; minor thin beds of dark brown-weathering dolostone

GILLESPIE LAKE GROUP (Wernecke Supergroup)

Undifferentiated divisions (not visited)

Upper division: Orange-brown weathering, thick-bedded dolostone with columnar stromatolites; interbedded fine sandstone and dark grey cherty mudstone; orange-weathering dolomitic limestone in

Chloritized schist, formerly mafic flow and tuff deposit

Lower division: Dark- and light brown-weathering, thin- to medium-bedded dolostone, locally cyclic from dolo-siltstone to

QUARTET GROUP (Wernecke Supergroup)

Dark grey- or green-grey weathering quartz siltstone, typically with white alteration along bedding planes; slaty cleavage

.... Limit of outcrop and geologic mapping

Geologic contact (defined, approximate, assumed)

Fault: sense of movement unknown (defined, approximate, assumed)

Normal fault; ball on downthrown side

Thrust fault: teeth on over-riding plate

Cleavage (inclined, vertical)

Jointing (inclined, vertical)

Sulphide occurrence (Zn: sphalerite; Pb: galena; Cu: chalcopyrite)

# **Economic Mineralization**

There are no mines in the map-area, although several prospects have been tested intermittently since the early fifties. Of the visited showings, only the Blende is currently under investigation.

1) 64°24'N 134°40'W. Blende: Sphalerite, galena, pyrite, chalcopyrite and tetrahedrite occur in gash veinlets and the matrix of brecciated dolostone within the upper division of the Gillespie Lake Group (GLG). Unmineralized dark siltstone and pinkish weathering dolostone that unconformably overlie the GLG are interpreted Pinguicula strata. The mineralization is found as much as 300 m below the ridge crest discovery zone and 800 m to the southeast. This trend is coincident with a faulted anticline, and mineralization is therefore considered both structurally and

stratigraphically controlled. References: D.I.N.A. 1982, P. 195

(2) 64015'N 134015'W Kathleen: Fine-grained galena was observed in limestone with abundant goethite, scorodite(?) and blue-black hematite; chalcopyrite is reported in coarse crystalline dolostone. Most trenches and drilling occur at the unconformity between deformed Gillespie Lake dolostone and outliers of white limestone of probable early Paleozoic age.

(3) 64017'N 134002'W Zap: Sphalerite and galena in dolostone breccia and tetrahedrite and galena in black baritic dolostone are reported. The latter rock type occurs along fault splays between quartzite (to the east) and blue-grey weathering dolostone (probably Cambrian) in a south-facing cirque. Pods of translucent crystalline barite up to 1 m wide are prominent in the carbonate cliffs. The quartzite is medium-grained, mature, and could be a Pinguicula unit (Corn Creek member of Eisbacher, 1982), or might be an unrecognized early Paleozoic sandstone, but is unlikely the Mississippian quartzite. This lithology appears to extend eastward into Rusty Mountain map

(4) 64018'N 134005'W. Barite float was found around an old pit at the unexposed) contact of oxidized Gillespie Lake dolostone with overlying

(5) 64017'N 134001'W Malachite and chalcopyrite flecks occur at contact of Gillespie Lake dolostone with overlying grey limestone (Paleozoic?). (6) 64019, 134006, W Chalcopyrite in 1 m wide white quartz vein cutting thin-bedded oxidized siltstone within the 'Pinguicula' Group.

## Notes

The southern Wernecke Mountains are accessible by helicopter from Mayo (100 km southwest), or from the end of the road at the abandoned settlement of Wernecke (60 km south-southwest). A winter road follows the south side of the Beaver River, and an overgrown branch road passes north of the Kathleen

┌2 km

The central part of the area, some 100 km<sup>2</sup> and including the peaks about 10 km north of Mount Good, is included in native land claims currently under negotiation.

### Regional Setting and Structure

The map area lies within the Foreland Fold and Thrust belt. In the southern Wernecke Mountains the southeast structural trend of Proterozoic strata bends northward around Misty Creek Embayment, a re-entrant of the Paleozoic Selwyn Basin (Cecile, 1981). Paleozoic carbonates in the top left corner of the map area are part of a northwest-plunging synclinorium that extends into the northern Ogilvie Mountains. Remnants of younger strata on ridge crests are recumbantly folded, suggesting that a higher structural level has been almost completely eroded.

The Proterozoic units are telescoped along contraction faults and most have been thickened by structural imbrication. Thrusts are visible in only a few places; mostly they are inferred from older-over-younger stratigraphic relationships where units can be positively identified.

Wide northeast-trending valleys disrupt structural trends and thrusts. Underlying post-Early Cretaceous, near-vertical faults probably had a combination of dextral and vertical slip. The northwest trend of the Wind River valley, and a parallel drainage 12 km south of it, truncate the northeast valleys. They appear to be distant splays of the Late Cretaceous and Early Tertiary Knorr-Richardson fault system (Norris and Hopkins, 1977) best displayed 50 km to the northeast. The Kathleen Lakes fault zone (Abbott, 1990) post-dates the diorite (as young as Triassic?) as well as Mississippian(?) quartzite. On a regional scale, it is a splay of the Dawson fault, and probably directly connected to the basal detachment.

These notes supplement discussion in Roots (1990). Some 3000 m of siliciclastic and dolomitic strata comprise the Quartet Group and Gillespie Lake Group in the map area. Both the Quartet and Pinguicula groups contain dark, fine-grained sediments that are difficult to distinguish (Table 1). Because thrusts are preferentially located within this incompetant lithology, the groups may be interleaved locally.

The Gillespie Lake Group is divided into a lower sub-unit (deep-water depositional environment) grading into an upper sub-unit (shallow and tide-dominated environment). The lower division includes beds that coarsen and thicken upward, trough cross-bedded arenite, and tool marked and slump-folded mudstone layers. The upper Gillespie Lake division is characterized by columnar and domal stromatolites in shallowing-upward sequences, beach rock, beach rosettes and dessication cracks (Mustard et al., 1990). Molar-tooth structure and oolite-pisolite beds are visible where preserved by early silicification. Chloritized volcanic tuff (locally mafic flow) less than 2 m thick was noted near the gradational boundary between these divisions in the Mount Good area.

Pinguicula is the group name applied to a complex assemblage of clastic and carbonate rocks deposited in the 600 Ma interval between the Gillespie Lake and Cambrian units. Internal stratigraphy is not understood, requiring integration with sequences northeast (type area) and west of the map area. The fine-laminated stromatolites in the pink dolostone member resemble those of the Lower Fifteenmile Group (proposed name, also Unit 5 of Thompson and Roots, 1982) in the Ogilvie Mountains. There a lateral facies change from carbonate shelf to deep-water clastics spans 30 km. The predominantly clastic succession in this map area contrasts with the shallow-water succession described by Eisbacher (1981) near Pinguicula Lake.

Lower Paleozoic, white-weathering limestone is separated from dolostone of similar age because composition reflects contrasting primary features. The dolostone unit ('CDb' in areas north and east; Norris, 1982) includes coarse syn-sedimentary breccia and arenaceous lenses. Pb-Zn-Cu and Ba showings occur in the southeast where the limestone unit unconformably overlies Gillespie Lake dolostone.

Tan-weathering sandstone may be Slats Creek Formation, which lies conformably beneath Cambrian dolostone 25 km north of the map area.

One of the most distinctive rock types is orange to grey-weathering, thin bedded limestone and dolostone laced with siderite veinlets and replacements along laminae. It is unclear whether this 'zebra rock' is a secondary characteristic, but this unit (tentatively correlated with Taiga Formation, Norris, 1982) occurs over significant areas near the Zap and north of the Blende showings.

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### Table 1. Lithological characteristics to distinguish dark clastic rocks of the Quartet Group from those in the Pinguicula

weathered surfaces diag-

nostic, limonitized

- monotonous succession of - thick sittstone successions but colour varies; olive and - white 'stringer' alteration - contains thin sandstone. along bedding planes, fine carbonate interbeds - pervasive micaceous cleav- - fissile and soft; breaks in age; breaks in long indu- small chips; disturbed pieces 'tinkle'

GEOLOGY OF 106D/8 AND 106D/7 (East half) MAP AREAS

Charlie Roots

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