# $\mathrm{U}-\mathrm{Pb}$ zircon age and Pb isotopic constraints on the age and origin of porphyry and epithermal vein mineralization in the eastern Dawson Range, Yukon 

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#### Abstract

Six dikes that are closely associated with the Klaza epithermal vein system in the Mt. Nansen district yield early Late Cretaceous U-Pb zircon ages (78.2-76.3 Ma); this age is similar to that obtained from the porphyry stock that hosts the Cyprus Cu-Mo-Au porphyry occurrence immediately to the southeast. These results support the interpretation that epithermal veins in the Mt. Nansen district are likely genetically related to subvolcanic magmatism. Granodiorite of the Dawson Range batholith that underlies most of the Klaza property gives a U-Pb zircon age of $107.9 \pm 0.3 \mathrm{Ma}$. These dates overlap with previously reported mid-Cretaceous U-Pb zircon ages for felsic dikes associated with the Brown-McDade and related vein and breccia deposits in the Mt. Nansen mine. The new results, together with regional dating and Pb isotopic data from western Yukon, emphasize the metallogenic importance of the "early Late Cretaceous" magmatic-hydrothermal event in this region.


## INTRODUCTION

The eastern Dawson Range in west-central Yukon hosts a wide variety of porphyry, epithermal and related styles of mineralization, including those in the Sonora Gulch area as well as the Freegold and Mt. Nansen camps (Fig. 1). Much of the mineralization in this region is known to be "early Late" Cretaceous in age and associated with the Casino plutonic suite (~78-72 Ma; Sonora Gulch, Nucleus, Revenue, Casino); however, at least some intrusionrelated deposits, such as Antoniuk (Fig. 1), have yielded unambiguous mid-Cretaceous formation ages (Allan et al., 2013). In addition, some occurrences, such as the Bonanza vein swarm in the Prospector Mountain area (Fig. 1) are spatially and temporally associated with the "late Late" Cretaceous (71-67 Ma) Prospector Mountain suite magmatism and deposition of Carmacks Group volcanic rocks (Allan et al., 2013).

The timing of formation of epithermal vein systems in the Mt. Nansen camp, including the main "Brown-McDade cluster" (Brown-McDade, Huestis, Weber, Flex and Dickson veins; Fig. 2), as well as the extensive area of veining in the Klaza deposit area to the north ("Klaza cluster"; Fig. 2), has been problematic. Mortensen et al. (2003) report U-Pb zircon ages ranging from $107.9 \pm 0.9 \mathrm{Ma}$ to $109.0 \pm 0.7 \mathrm{Ma}$ for northwest-trending quartz-feldspar porphyry dikes that are closely associated with the veins in the main BrownMcDade cluster, and, in the absence of any other direct constraints on the timing of mineralization, concluded that the veins formed at approximately the same time. Hart and Langdon (1998) had previously argued that epithermal veins in the Mt. Nansen camp (including both those in the main camp and the Klaza system) could represent parts of a "porphyry-to-epithermal transition" that was centered on the Cyprus porphyry and contained Cu-Mo-Au porphyrystyle mineralization (Fig. 2). However, a Re-Os age of $71.1 \pm 0.3$ Ma reported by Selby and Creaser (2001) for molybdenite from the Cyprus porphyry deposit was incompatible with such a model if the Mt. Nansen veins were indeed middle Cretaceous in age.

A U-Pb dating study of intrusive rock units that either host or are spatially closely associated with mineralization in the Mt. Nansen area was undertaken. This study also evaluates existing and new measured Pb isotopic compositions of igneous feldspar minerals from many of the intrusions as well as sulphides (mainly galena) from the
various deposits and occurrences. The aim of the work is to better constrain the timing of mineralization throughout the eastern Dawson Range, and to test models such as the "porphyry-to-epithermal transition" that was proposed for the Mt. Nansen area. New results resolve some of the questions concerning the age(s) of mineralization in much of the area, although some areas of uncertainty still exist.

## GEOLOGY OF THE EASTERN DAWSON RANGE

The eastern Dawson Range is mainly underlain by intermediate to felsic intrusive rocks of the Early Cretaceous Dawson Range batholith (Whitehorse plutonic suite; Fig. 1). Eruptive equivalents of the Whitehorse suite are locally preserved as the Mt. Nansen Group volcanic rocks in the vicinity of Mt. Nansen (Figs. 1 and 2). The batholith was intruded into metamorphic rocks of the Yukon-Tanana terrane, as well as large bodies of Late Triassic and Early Jurassic plutonic rock of the Minto plutonic suite that underlie much of the easternmost end of the Dawson Range (Fig. 1). Small bodies of hypabyssal felsic porphyry of early Late Cretaceous age are widespread in the region; some of these host, or are spatially closely associated with, some of the main mineral deposits and occurrences in the area (e.g., Casino, Nucleus, Revenue). The slightly younger late Late Cretaceous Carmacks Group mafic volcanic rocks cover extensive areas in the eastern Dawson Range (Fig. 1). Paleocene and Eocene felsic intrusive and extrusive rocks are increasingly abundant along the southwestern flank of the Dawson Range.

## MINERALIZATION IN THE MT. NANSEN DISTRICT

The Mt. Nansen district (Fig. 2) includes two main concentrations of mineralized veins: the "Brown-McDade cluster", centered approximately 11 km southeast of the top of Mt. Nansen and includes the Brown-McDade vein system that was mined by BYG Resources in 1995-98; and the "Klaza cluster", centered approximately 3 km eastnortheast of Mt. Nansen and based on recent exploration work by Rockhaven Resources, contains an inferred resource of 948 koz of gold and 21.8 Moz of silver (Wegnzynowski et al., 2015; Fig. 2).


Figure 1. Regional geology of the central and eastern Dawson Range, showing locations of some of the main mineral deposits and occurrences referred to in the text. Box shows the location of Figure 2.


Figure 2. Geology of the Mt. Nansen District, showing locations of the Klaza and Brown-McDade vein clusters and the Cyprus porphyry. Ages shown in the Brown-McDade cluster are ID-TIMS U-Pb zircon ages from Mortensen et al. (2003); ages for the Mt. Nansen Group are from Mortensen et al. (unpublished); Re-Os age from the Cyprus porphyry is from Selby and Creaser (2001).

The Brown-McDade cluster consists of veins and deposits south of the Dickson fault, comprising the Webber, Flex, Heustis, Dickson, Vince and Brown-McDade occurrences (Fig. 3; Hart and Langdon, 1998; Anderson and Stroshein,

1998; Stroshein, 1999). These are historically amongst the most economic in the district and cumulatively host a resource of about 1 Mt with grades ranging from 5 to $14 \mathrm{~g} / \mathrm{t} \mathrm{Au}$ and 50 to $300 \mathrm{~g} / \mathrm{t} \mathrm{Ag}$.


Figure 3. Geology the Mt. Nansen District including the Brown-McDade cluster in the southeast, the Cyprus porphyry and related occurrences in the central area, and the Klaza cluster in the northwest (modified from Hart and Langdon, 1998).

Most occurrences are hosted in Yukon-Tanana terrane schist and gneiss, but the Brown-McDade deposit, which is the largest, is mostly hosted in granodiorite. Each occurrence typically includes a 'swarm' of several veins, but may be dominated by a larger, single vein. Most veins are at least partly hosted by, northwest-trending quartz-feldspar porphyry dikes and small quartz-porphyry plugs. Most veins strike northwesterly between 320 and $340^{\circ}$ and dip moderate to steeply ( 65 to $80^{\circ}$ ) to the southwest. Many of these veins have strike lengths of 500 m or greater, but their extents are variably displaced along north-northeast-striking sinistral strike slip faults. They are typically 0.2 to 0.8 m wide, but locally up to 2.5 m thick. The thicker parts result from intersections with $020^{\circ}$-striking fractures.

The veins are typically crudely banded, but are locally, variably brecciated. The Brown-McDade deposit occurs mostly as a breccia pipe (Stroshein, 1999). All veins are generally similar in mineralogy consisting of 'cherty' finegrained grey and white quartz, with varying abundances of sulphide minerals dominated by pyrite and arsenopyrite, with essential galena and sphalerite typically to $10 \%$, and lesser stibnite, tetrahedrite and various sulphosalt minerals such as boulangerite and jamesonite.

The Klaza area (Fig. 4) is mainly underlain by massive and unfoliated, medium to coarse-grained, equigranular hornblende-biotite granodiorite and quartz monzonite of the Dawson Range batholith, which is part of the regionally developed Whitehorse plutonic suite. A stock of quartz monzonite to granite, termed the Cyprus porphyry, intrudes the Dawson Range batholith on the southeastern corner of the Klaza property (Figs. 2 and 3). The Dawson Range batholith is cut by a set of northwest-trending, steeply to moderately southwest-dipping fault zones, along which dikes of vari-coloured (grey, green, tan, red and purple) feldspar and quartz-feldspar porphyry have been emplaced (Fig. 4). Individual porphyry dikes commonly pinch and swell along strike, and are up to 30 m wide. They consist of an aphanitic groundmass with up to $15 \%$ K-feldspar phenocrysts ( 1 to 2 mm ) and minor biotite and rare quartz phenocrysts. Contacts with the granodiorite range from sharp and undulating to brecciated with locally abundant gouge development. A second, much less common set of northwest-trending dikes is also locally present, particularly in the southeastern part of the Klaza property, near the Cyprus porphyry stock. These dikes are typically fine grained to aphanitic, and appear to be intermediate to mafic in composition. They contain a
significant amount of magnetite, and are locally strongly magnetic. These mafic dikes are up to 100 m in width, and are larger and more abundant near the Cyprus porphyry.

The main mineralized zones at Klaza range from 1 to 100 m wide and are usually associated with feldspar porphyry dikes. Mineralization occurs as veins, sheeted veinlets and tabular breccia bodies. The feldspar porphyry dikes generally occupy the same northwest-trending structural zones as the mineralized veins and they are commonly strongly fractured. They are locally cut by the main mineralized veins as well as by late stage white carbonate veinlets. Mineralization at Klaza mainly consists of gold and silver bearing, sulphide-rich veins that comprise pyrite, arsenopyrite, galena, sphalerite, various sulphosalts and electrum in a gangue of quartz, rhodocrosite and barite. The mafic dikes are also locally mineralized, hosting small amounts of disseminated pyrite and pyrite stringers, together with minor chalcopyrite and rare molybdenite. The strongest copper mineralization on the property is always associated with argillic altered versions of the mafic dikes.

The feldspar porphyry dikes at the Klaza property have been interpreted to emanate from the Cyprus porphyry stock to the southeast. Two distinct centres of $\mathrm{Cu}-\mathrm{Mo} \pm \mathrm{Au}$ porphyry-style mineralization have been identified within the Cyprus porphyry stock, the Cyprus zone and the Kelly zone (Fig. 3). These zones appear to contain modest grades and be of limited extent (Wengzynowki et al., 2015).

The northwest-trending feldspar porphyry dikes and the mineralized structures at Klaza have been displaced by two northeast trending faults (Fig. 4). The exact timing of these faults is uncertain; however, they appear to be coeval with, or post-date, the mineralized zones on the property.

## U-PB ZIRCON GEOCHRONOLOGY

## METHODOLOGY

Zircons were separated from 2-4 kg samples using conventional crushing, grinding, wet shaking table, heavy liquids and magnetic separation methods. U-Pb dating of zircons was done using laser ablation (LA-) ICP-MS methods at the Pacific Centre for Isotopic and Geochemical Research (PCIGR) at the University of British Columbia. Analytical techniques employed are as described by Allan et al. (2013). Twenty individual


Figure 4. Map of the Klaza property showing the distribution of the various zones as well as drill holes and the specific locations of samples that were analyzed in this study (modified from Wengzynowki et al., 2015). UTM coordinates are in NAD83 datum.
zircon grains were analyzed from each sample. Complete analytical data are provided in Appendix 1 and the results are plotted as compilations of the calculated ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages for individual sample analyses in Figures 3 and 4. Errors quoted for final assigned ages are given at the $2 \sigma$ level.

## RESULTS

All zircons analyzed from the Klaza U-Pb samples yielded concordant analyses. Sample locations are shown on Figure 3, and interpretations of the results for each of the samples are presented briefly below.

## Granodiorite of the Dawson Range batholith

Twenty zircon analyses from a sample of massive mediumgrained granodiorite from the Pika Zone (Fig. 4) yield a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $107.9 \pm 0.3 \mathrm{Ma}(M S W D=0.74$; $\mathrm{POF}=0.78$; Fig. 5a).

## Feldspar porphyry dikes in the Western Klaza Zone

Twenty zircon grains from a sample of feldspar porphyry from DDH KL-12-75 (189-191 m; Fig. 4) yield a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $76.9 \pm 0.3 \mathrm{Ma}(\mathrm{MSWD}=1.00 ; \mathrm{POF}=0.45$;
Fig. 5b). Eighteen zircon grains from a second feldspar porphyry dike from the Western Klaza Zone in DDH KL-14-185 (0-11.5 m; Fig. 4) yield a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $76.8 \pm 0.4 \mathrm{Ma}(M S W D=1.4 ; \mathrm{POF}=0.14$; Fig. 5c). A single grain from this sample gives a much older age ${ }^{207} \mathrm{~Pb} /{ }^{206} \mathrm{~Pb}$ age of $1600.6 \pm 40.6 \mathrm{Ma}$ ) and represents an older xenocryst likely entrained from underlying metamorphic basement rocks. One of the analyses gave a significantly younger ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age, likely related to postcrystallization Pb-loss, and is not included in the calculated weighted average age.

## Feldspar porphyry dike from the Western BRX Zone

Eighteen zircon grains from a porphyry dike sample in DDH KL-14-154 (52.1-52.5 m; Fig. 4) yield a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $77.1 \pm 0.4 \mathrm{Ma}(\mathrm{MSWD}=0.78 ; \mathrm{POF}=0.72$; Fig. 5 d ). Two grains give ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages of 102.6 and 105.0 Ma , and are interpreted to have been xenocrystic zircons entrained from the underlying Dawson Range batholith.

## Feldspar porphyry dike from the Central BRX Zone

Sixteen zircon grains from a porphyry dike in DDH KL-14165 (85.7-87.5 m; Fig 4) give a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $76.5 \pm 0.4 \mathrm{Ma}(M S W D=0.68 ; ~ P O F=0.80 ; ~ F i g .5 e)$. Four grains give ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages of 105.0-105.7 Ma and a single grain gives an age of 369.7 Ma ; these are interpreted to have been xenocrysts from the Dawson Range batholith and the metamorphic basement rocks, respectively.

## Feldspar porphyry dike from the Central Klaza Zone

Seventeen zircon grains from a porphyry dike in DDH KL-11-12 (181-187 m; Fig. 4) give a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $71.6 \pm 0.3 \mathrm{Ma}(\mathrm{MSWD}=0.55$; $\mathrm{POF}=0.92$; Fig. 5f). Two grains give slightly younger ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages and are not included in the calculated weighted average age. A single grain gave a ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $77.1 \pm 1.7 \mathrm{Ma}$. The calculated age for this sample is significantly younger than that of most of the other Klaza dikes that were dated, and it appears to represent a distinctly younger intrusion. The presence of the single 77.1 Ma zircon indicates that the magma did interact to some extent with the "early Late" Cretaceous rocks.

## Mafic dike in the Eastern BRX Zone

Sixteen zircon grains from a mafic dike in DDH KL-12-134 (218-221.5 m) give a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $78.2 \pm 1.0 \mathrm{Ma}(M S W D=0.34 ; \mathrm{POF}=1.00$; Fig. 6a). Four grains give ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages of 107.4-109.8 Ma, and are interpreted to have been xenocrysts entrained from the Dawson Range batholith.

## Feldspar porphyry dike from the Far Eastern Klaza/ Pearl Zone

Fifteen zircon grains from a feldspar porphyry dike in DDH KL-14-162 (86.2-86.5 m; Fig. 4) give a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $76.3 \pm 0.3 \mathrm{Ma}(M S W D=0.49 ;$ POF=0.94; Fig. 6b). Two grains give slightly older ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages and are excluded from the weighted average age calculation. Two
other grains gave ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages of 101.4 and 102.6 Ma , and a single grain gave an age of 446.9 Ma ; these are interpreted to have been xenocrysts from the Dawson Range batholith and the underlying metamorphic bedrock, respectively.

## Cyprus porphyry

Seventeen zircon grains from sample 97-CH-33-3 in the Cyprus porphyry stock give a weighted ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ age of $76.0 \pm 0.4 \mathrm{Ma}(\mathrm{MSWD}=0.75 ; \mathrm{POF}=0.74$; Fig. 6c). Three grains give slightly younger ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages, likely reflecting the effects of post-crystallization Pb-loss and were excluded from the weighted average age calculation.

## SUMMARY OF KLAZA DATING RESULTS

Results from the dating study of the Klaza samples confirm an Early Cretaceous age of $107.9 \pm 0.6$ Ma for granodiorite of the Dawson Range batholith in the Mt. Nansen area.
All but one of the northwest-trending dikes associated with Klaza mineralization give ages ranging from 78.2 to 76.0 Ma , indicating that these intrusions are part of the $\sim 78-72 \mathrm{Ma}$ Casino plutonic suite. An intermediate composition dike that contains disseminated chalcopyrite and minor molybdenite (sample KL-12-134; Eastern BRX zone), and was initially speculated to be part of the younger Prospector Mountain suite, actually yielded the oldest age at $78.2 \pm 1.0 \mathrm{Ma}$. A sample of the felsic porphyry that hosts the Cyprus Cu-Mo-Au porphyry occurrence returned a crystallization age of $76.0 \pm 0.4 \mathrm{Ma}$, confirming that this body is coeval and likely co-magmatic with the felsic dikes in the Klaza area. One of the quartzfeldspar dike samples from Klaza (KL-11-12; Central Klaza zone) gave a significantly younger age of $71.6 \pm 0.3 \mathrm{Ma}$, suggesting that this body either represents a late stage Casino suite intrusion or (less likely) an early intrusion of the Prospector Mountain suite.

## RE-EXAMINATION OF DATING RESULTS FROM THE BROWN-MCDADE CLUSTER

The porphyry dikes at Klaza are very similar to those previously dated in the Brown-McDade cluster, in terms of composition, structural setting in northwest-trending faults, and association with sulphide-rich epithermal veins. However, Mortensen et al. (2003) report multi-grain TIMS zircon ages for five porphyry dikes and small stocks from the Brown-McDade cluster ranging from 109.0 to 107.9 Ma , suggesting that these intrusions are temporally related to the Dawson Range batholith. The TIMS data for


Figure 5. Plots of ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages for zircons from Klaza igneous samples, with calculated weighted average ages. Error bars are shown at the $2 \sigma$ level. Analyses shown by blue boxes were not used in calculation of the weighted average ages.


Figure 6. Plots of ${ }^{206} \mathrm{~Pb} /{ }^{238} \mathrm{U}$ ages for zircons from Klaza igneous samples, as well as the Cyprus porphyry and re-analysis of three porphyry samples from the Brown-McDade cluster. Error bars are shown at the $2 \sigma$ level. Analyses shown by blue boxes were not used in calculation of the weighted average ages.
some of the samples show considerable scatter and there is evidence for a substantial amount of older inherited zircon present in some cases. Although considered unlikely, it is conceivable that the Early Cretaceous zircons dated in 2003 might have been of xenocrytic origin, and entrained from underlying bodies of 108 Ma Dawson Range batholith; the Brown-McDade cluster dikes and plugs are mostly emplaced into Yukon-Tanana terrane metamorphic rocks at the present level of exposure. In order to evaluate this possibility, twenty zircon grains from each of three samples from the 2003 study, selected as representative of the entire range of grain size and morphology present in the concentrate, were analyzed using LA-ICP-MS methods (Appendix 1; Fig. 6d-f). The new results confirm the original age assignments: Flex zone porphyry, 108.4 $\pm 0.7$ Ma TIMS and $107.4 \pm 0.2$ Ma LA; Weber zone porphyry, $107.9 \pm 0.9 \mathrm{Ma}$ TIMS and 107.4 $\pm 0.4$ LA; Brown-McDade porphyry: $109.0 \pm 0.7 \mathrm{Ma}$ TIMS; 107.6 $\pm 0.3 \mathrm{Ma}$ LA. Therefore, the reanalysis of zircons from these samples strongly argues for an Early Cretaceous crystallization age for the intrusions in the Brown-McDade cluster. The most reasonable conclusion, therefore, is that two similar, but completely unrelated sets of felsic dikes and small stocks were emplaced into northwest-trending faults in the Mt. Nansen area.

## Pb ISOTOPES

## METHODOLOGY

In general the Pb isotopic compositions of sulphides precipitated from magmatic-hydrothermal fluids in intrusion-related vein systems are very similar to the Pb isotopic compositions of igneous feldspar within the genetically related intrusions (Tosdal et al., 1999). Lead isotopes therefore provide a powerful tool for determining the source(s) from which Pb (and by analogy, other metals) in a deposit is derived, and in the case of intrusionrelated mineralization, for identifying which intrusion (or intrusive suite) is genetically associated with the mineralization. In some cases, Pb compositions can also yield some indication of the age of formation of epigenetic mineralization, although this must be utilized with caution. Lead isotopic compositions of several samples of galena from veins from both the Brown-McDade and the Klaza cluster were determined at the PCIGR, using TIMS methods as described by Mortensen and Gabites (2002). Analytical data are given in Appendix 2. An extensive Pb isotopic database for both igneous feldspar minerals from intrusive rocks and galena and other sulphides from
various mineral deposits and occurrences from throughout west-central Yukon, including data from various published and unpublished sources, is also discussed.

## RESULTS

New galena Pb isotopic analyses from Klaza ( $\mathrm{n}=3$ ) and Brown-McDade ( $n=3$ ) from this study are plotted on a ${ }^{207} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ vs. ${ }^{206} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ diagram in Figure 7. Also shown for comparison on this plot are fields of Pb isotopic compositions from additional Klaza and Brown-McDade vein samples, as well as sulphides from a number of other epigenetic, polymetallic vein occurrences (Fig. 1) in the Dawson Range and elsewhere in western Yukon, including the Prospector Mountain area (Frog and Lilypad occurrences), the Tinta Hill deposit in the Freegold Mountain area, the Longline Au deposit in the Moosehorn Range (Yukon MINFILE 115N 024), the Bomber vein at Casino, and numerous other Late Cretaceous vein occurrences in the region (data from Godwin et al., 1988; Glasmacher, 1990; Smuk, 1999; Selby et al., 2001; Joyce, 2002; Selby et al., 2001; Bineli-Betsi et al., 2013; Mortensen, unpublished data). Also shown on the plot are fields for igneous K-feldspars from Early Cretaceous intrusions in the Moosehorn Range (from Joyce, 2002) and Late Cretaceous intrusions in the Casino area (from Selby et al., 1999), as well as other Early and Late Cretaceous intrusions from throughout west-central and western Yukon (Mortensen, unpublished data).

The Pb isotopic study tries to resolve the question of whether veins in the Brown-McDade cluster formed at the same time as those in the Klaza cluster (i.e., early Late Cretaceous), and coincidentally happen to be spatially associated with a lithologically similar but $\sim 35$ my older set of porphyry dikes, or represent a completely unrelated mineralizing event. The study also evaluates the relative importance of mid-Cretaceous vs. early Late and late Late Cretaceous mineralizing events in western Yukon.
The field for Early Cretaceous igneous feldspar from the Moosehorn Range (dotted green line in Fig. 7) is completely separate from that for sulphides in gold-bearing veins from the Longline occurrence that is hosted by those intrusions (solid green line). This observation led Joyce (2002) to conclude that the metals in the Longline veins were not derived from the host intrusions, or indeed from any other intrusions in western Yukon or eastern Alaska, but were actually orogenic rather than intrusion-related.
The field for igneous feldspar from other Early Cretaceous intrusions in western Yukon (dotted red line in Fig. 7)
overlaps with the Moosehorn Range feldspar data but shows somewhat more scatter. Feldspar from early Late Cretaceous Casino suite intrusions and from other early and late Late Cretaceous intrusions in western Yukon (blue dotted lines on Fig. 7; including the Cyprus porphyry in the Mt. Nansen district, shown as the blue triangle) show considerable scatter; although there is substantial overlap with the fields for the Early Cretaceous feldspar, most of the Late Cretaceous feldspar show higher ${ }^{206} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ ratios and slightly lower ${ }^{207} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ ratios (Fig. 7). The field for galena from the Bomber veins, which are considered to be related to the early Late Cretaceous Casino porphyry deposit, overlap with both the more radiogenic (higher $\mathrm{Pb} / \mathrm{Pb}$ compositions) part of the Early Cretaceous feldspar Pb field and the less radiogenic (lower $\mathrm{Pb} / \mathrm{Pb}$ compositions) part of the Late Cretaceous feldspar field. However, galena from other early and late Late Cretaceous intrusion-related veins in western Yukon (solid blue line on Fig. 7), show a good overlap with Late Cretaceous feldspar but are completely separate from the Early Cretaceous feldspar field. Taken together, the data suggest: 1) Early and Late Cretaceous igneous feldspar minerals yield partly overlapping compositional fields, but most Late Cretaceous samples have higher ${ }^{206} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ and slightly lower ${ }^{207} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ ratios than the bulk of the Early Cretaceous samples; and 2) epigenetic vein occurrences that are known to be of Late Cretaceous age more closely match the compositional fields for the Late Cretaceous igneous feldspar than the Early Cretaceous samples, which is consistent with the bulk of the metals in these occurrences having been derived predominantly from the Late Cretaceous intrusions.
Sulphide Pb isotopic compositions from four separate areas (Klaza, Brown-McDade, Prospector Mountain, Tinta Hill), however, are more puzzling. Field relationships discussed above provide unequivocal evidence that the Klaza veins (pink field in Fig. 7) are spatially and temporally related with late Late Cretaceous porphyry intrusions, whereas very similar veins in the Brown-McDade cluster (blue field in Fig. 7) are most reasonably interpreted to be Early Cretaceous in age. Both of these compositional fields overlap with the least radiogenic part of the compositional ranges for the Early and Late Cretaceous igneous rocks. Veins in the Prospector Mountain area, whose galena Pb isotopic compositions overlap almost perfectly with the Brown-McDade galena field (Fig. 7), are interpreted to cut Late Cretaceous Carmacks Group volcanic rocks, and must therefore be Late Cretaceous in age. The Tinta Hill Au-Ag-Cu-Pb-Zn deposit in the Mt. Freegold area is a
polymetallic, epithermal vein that closely resembles the Klaza and Brown-McDade veins in most respects (Bennett and Bineli Betsi, 2010). The age of the Tinta Hill veins is uncertain; Bineli Betsi and Bennett (2010) stated that the vein is crosscut by a porphyry dike that yielded a 108 Ma U-Pb zircon (TIMS) age; however, there is somewhat less certainty regarding this crosscutting relationship in subsequent publications (Bineli Betsi, 2012; Bineli Betsi et al., 2013). These four vein systems were each emplaced into very different host rocks that would be expected to have quite distinct Pb isotopic characteristics. The BrownMcDade cluster of veins are mainly hosted by YukonTanana terrane metasedimentary rocks; whereas the Klaza veins are hosted by the Early Cretaceous Dawson Range batholith, the Prospector Mountain veins by Late Cretaceous Carmacks Group volcanic rocks, and the Tinta vein by the Early Jurassic Granite Mountain batholith. The fact that sulphides from these four vein systems show rather similar Pb isotopic compositions suggests that the Pb in each was likely derived mainly from the associated intrusions rather than the host rocks, and in general the Pb isotopic compositions of the sulphides is more consistent with the metals having been derived from Early, rather than Late, Cretaceous magma.

The implications of the Pb isotopic study are therefore uncertain at this time. It is very difficult to directly date epithermal veins such as those in the eastern Dawson Range, and vein ages typically must be inferred from contact relations with dated igneous rock units. More detailed geological investigations will be required to resolve the age relations between the various intrusive suites and vein systems in the area.

## IMPLICATIONS FOR THE METALLOGENY OF WEST-CENTRAL YUKON

At least three distinct ages of porphyry-style mineralization are known to exist in western Yukon: Early Cretaceous (e.g., Pattison, Idaho); early Late Cretaceous (e.g., Casino, Cash, Cyprus, Nucleus, Revenue, Sonora Gulch, Tad/ Toro and Bluff/Taurus in eastern Alaska); and late Late Cretaceous (e.g., Mt. Cockfield, Sixtymile; Allan et al., 2013). Similarly there are gold-bearing vein and breccia systems of the same three ages: Early Cretaceous (e.g., Antoniuk, possibly Brown-McDade, possibly Tinta Hill); early Late Cretaceous (e.g., Klaza); and late Late Cretaceous (e.g., Prospector Mountain, Connaught (Yukon

MINFILE 115N 040)). The largest and most prospective porphyry systems, however, appear to be related to the early Late Cretaceous Casino suite, and the most significant epithermal vein systems identified thus far appear to be the Klaza deposit and some zones in the Sonora Gulch area, also of early Late Cretaceous age. Our new age constraints strongly support the "porphyry-to-epithermal transition" model that was proposed by Hart and Langdon (1998), at least for veins of the Klaza cluster and the Cyprus and Kelly porphyry deposits. The age of the epithermal veins in the Brown-McDade cluster, however, is still not resolved, and it remains unclear whether the Brown-McDade veins have any relationship to any of the Late Cretaceous magma in the Mt. Nansen area.

Lead isotopic studies of many of the mineral deposits and occurrences in western Yukon, together with analyses of igneous feldspars from potentially related intrusive rocks, indicate that there is a substantial amount of scatter within each of the plutonic suites in the region. It appears that lead isotopes do not provide a simple tool for identifying which intrusive rocks are genetically related to specific mineralization in this region. It is unclear whether this is because of the isotopic heterogeneity of the igneous rocks themselves, or the fact that some amount of mixing has occurred between metal-bearing magmatic fluids and metals derived from compositionally variable host rocks.


Figure 7. Compilation plot showing Pb isotopic compositional fields for igneous feldspars from Early to Late Cretaceous intrusions from west-central Yukon, together with sulphide (mainly galena) compositions from a variety of intrusion-related orogenic gold deposits and occurrences. Fields for sulphide Pb analyses from the Mt. Nansen veins and from the Klaza deposit are shown as colored fields. The Cyprus porphyry feldspar Pb is shown by the blue triangle. Sources of data are discussed in the text.

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APPENDIX 1. U-Pb analytical data by laser ablation (LA)-ICP-MS methods.

| Fraction | Isotopic Ratios |  |  |  |  |  |  | Isotopic Ages |  |  |  |  |  | Background corrected mean counts per second at specified mass |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | ${ }^{207} \mathrm{~Pb} /{ }^{235} \mathrm{U}$ | \% 1\% | ${ }^{209} \mathrm{~Pb} /{ }^{219} \mathrm{U}$ | \% $1 \sigma$ | rho | ${ }^{20} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ | \% 1 $\sigma$ | ${ }^{207} \mathrm{~Pb} /{ }^{35} \mathrm{U}$ | $1{ }^{\circ}$ | ${ }^{206} \mathrm{~Pb} /{ }^{23} \mathrm{U}$ | $1 \sigma$ | ${ }^{20} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ | $1{ }^{\circ}$ | 202 | 204 | 206 | 207 | 208 | 232 | 235 | 238 |
| Dawson Range batholith granodiorite - Pika Zone |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 1 | 0.11355 | 0.00624 | 0.01703 | 0.00032 | 0.34 | 0.04933 | 0.00256 | 163.7 | 117.2 | 109.2 | 5.69 | 108.8 | 2.02 | 1 | 15 | 4014 | 194 | 378 | 21508 | 692 | 72551 |
| 2 | 0.11273 | 0.00341 | 0.01708 | 0.00018 | 0.35 | 0.04837 | 0.00138 | 117.5 | 65.74 | 108.5 | 3.11 | 109.2 | 1.11 | 24 | 32 | 3436 | 163 | 524 | 32456 | 585 | 61901 |
| 3 | 0.11333 | 0.00504 | 0.0165 | 0.00025 | 0.34 | 0.04987 | 0.0021 | 188.8 | 94.96 | 109 | 4.6 | 105.5 | 1.61 | 50 | 0 | 6206 | 304 | 605 | 31463 | 1084 | 115710 |
| 4 | 0.11395 | 0.00431 | 0.01678 | 0.00022 | 0.35 | 0.04926 | 0.00176 | 160.4 | 81.28 | 109.6 | 3.93 | 107.3 | 1.38 | 0 | 8 | 7453 | 361 | 630 | 36829 | 1279 | 136678 |
| 5 | 0.1122 | 0.0033 | 0.01684 | 0.00017 | 0.34 | 0.04839 | 0.00134 | 118.4 | 63.98 | 108 | 3.01 | 107.6 | 1.05 | 36 | 9 | 4504 | 214 | 456 | 24503 | 771 | 82318 |
| 6 | 0.11538 | 0.00401 | 0.01646 | 0.00019 | 0.33 | 0.04969 | 0.00163 | 180.7 | 74.67 | 110.9 | 3.65 | 105.2 | 1.22 | 0 | 0 | 5200 | 254 | 671 | 37560 | 889 | 97243 |
| 7 | 0.11335 | 0.00304 | 0.01693 | 0.00015 | 0.33 | 0.04901 | 0.00124 | 148.2 | 58.03 | 109 | 2.77 | 108.2 | 0.98 | 0 | 0 | 5375 | 259 | 743 | 42442 | 922 | 97722 |
| 8 | 0.10818 | 0.00282 | 0.01693 | 0.00015 | 0.34 | 0.04706 | 0.00116 | 51.7 | 57.93 | 104.3 | 2.59 | 108.2 | 0.94 | 27 | 2 | 12980 | 601 | 1285 | 76215 | 2241 | 235989 |
| 9 | 0.1135 | 0.00625 | 0.01706 | 0.00028 | 0.3 | 0.0491 | 0.00258 | 152.8 | 118.81 | 109.2 | 5.7 | 109.1 | 1.8 | 0 | 1 | 1309 | 63 | 175 | 8412 | 224 | 23616 |
| 10 | 0.11789 | 0.0034 | 0.01698 | 0.00017 | 0.35 | 0.05042 | 0.00137 | 214.3 | 61.62 | 113.2 | 3.09 | 108.5 | 1.06 | 16 | 3 | 4327 | 214 | 483 | 23909 | 734 | 78429 |
| 11 | 0.11212 | 0.0039 | 0.01691 | 0.0002 | 0.34 | 0.04837 | 0.00159 | 117.5 | 75.51 | 107.9 | 3.56 | 108.1 | 1.25 | 55 | 0 | 3592 | 171 | 429 | 24405 | 615 | 65367 |
| 12 | 0.11682 | 0.00349 | 0.01692 | 0.00017 | 0.34 | 0.05148 | 0.00144 | 262.3 | 63.16 | 112.2 | 3.17 | 108.2 | 1.1 | 0 | 21 | 3481 | 176 | 437 | 24168 | 609 | 63311 |
| 13 | 0.10328 | 0.00474 | 0.01671 | 0.00026 | 0.34 | 0.04829 | 0.00211 | 113.3 | 99.98 | 99.8 | 4.36 | 106.8 | 1.65 | 0 | 0 | 4156 | 197 | 595 | 33906 | 771 | 76564 |
| 14 | 0.10225 | 0.00436 | 0.01693 | 0.00024 | 0.33 | 0.04474 | 0.00181 | 0.1 | 26.11 | 98.9 | 4.02 | 108.2 | 1.52 | 26 | 7 | 4808 | 211 | 781 | 43916 | 835 | 87436 |
| 15 | 0.11135 | 0.00301 | 0.01696 | 0.00015 | 0.33 | 0.04807 | 0.00123 | 102.5 | 59.18 | 107.2 | 2.75 | 108.4 | 0.96 | 68 | 7 | 4579 | 216 | 491 | 25890 | 784 | 83092 |
| 16 | 0.11023 | 0.00562 | 0.01706 | 0.00028 | 0.32 | 0.04823 | 0.00234 | 110.4 | 110.52 | 106.2 | 5.14 | 109 | 1.78 | 0 | 24 | 2028 | 96 | 173 | 9235 | 352 | 36607 |
| 17 | 0.11931 | 0.00658 | 0.01667 | 0.00032 | 0.35 | 0.05012 | 0.0026 | 200.6 | 115.98 | 114.4 | 5.97 | 106.6 | 2.01 | 0 | 0 | 6677 | 329 | 409 | 14809 | 1113 | 123250 |
| 18 | 0.10491 | 0.00371 | 0.01686 | 0.0002 | 0.34 | 0.04726 | 0.00158 | 61.8 | 78.25 | 101.3 | 3.41 | 107.8 | 1.28 | 0 | 17 | 5088 | 236 | 660 | 38469 | 910 | 92867 |
| 19 | 0.10711 | 0.00416 | 0.01707 | 0.00021 | 0.32 | 0.04467 | 0.00164 | 0.1 | 13.97 | 103.3 | 3.82 | 109.1 | 1.32 | 0 | 0 | 3610 | 158 | 431 | 27237 | 597 | 65101 |
| 20 | 0.05001 | 0.00549 | 0.01662 | 0.0002 | 0.11 | 0.022 | 0.0024 | 0.1 | 0 | 49.5 | 5.31 | 106.3 | 1.26 | 0 | 0 | 3772 | 81 | 601 | 32671 | 658 | 69853 |
| Western Klaza Zone porphyry dyke - DDH KL-12-75 189-191m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| - | 0.07326 | 0.00194 | 0.01197 | 0.00011 | 0.35 | 0.04433 | 0.00113 | 0.1 | 0 | 71.8 | 1.83 | 76.7 | 0.67 | 0 | 10 | 9147 | 399 | 485 | 36628 | 2197 | 235218 |
| 2 | 0.07988 | 0.00195 | 0.01196 | 0.0001 | 0.34 | 0.04959 | 0.00117 | 175.9 | 53.97 | 78 | 1.83 | 76.6 | 0.64 | 7 | 0 | 5256 | 256 | 290 | 20110 | 1295 | 135282 |
| 3 | 0.0782 | 0.00189 | 0.01191 | 0.0001 | 0.35 | 0.04733 | 0.0011 | 65.2 | 55.14 | 76.5 | 1.78 | 76.3 | 0.62 | 0 | 0 | 5489 | 255 | 310 | 24373 | 1318 | 141857 |
| 4 | 0.07836 | 0.00221 | 0.01192 | 0.00012 | 0.36 | 0.04742 | 0.0013 | 69.6 | 64.35 | 76.6 | 2.08 | 76.4 | 0.74 | 27 | 22 | 8117 | 379 | 357 | 29685 | 1950 | 209572 |
| 5 | 0.0708 | 0.00299 | 0.01175 | 0.00017 | 0.34 | 0.04369 | 0.0018 | 0.1 | 0 | 69.5 | 2.83 | 75.3 | 1.06 | 32 | 16 | 6166 | 265 | 363 | 29922 | 1510 | 161600 |
| 6 | 0.07708 | 0.00209 | 0.01189 | 0.00011 | 0.34 | 0.04727 | 0.00124 | 62.4 | 61.84 | 75.4 | 1.97 | 76.2 | 0.7 | 26 | 7 | 6747 | 314 | 268 | 18457 | 1642 | 174712 |
| 7 | 0.07658 | 0.0014 | 0.01209 | 0.00007 | 0.32 | 0.04638 | 0.00081 | 17.3 | 40.8 | 74.9 | 1.32 | 77.5 | 0.47 | 18 | 12 | 10463 | 478 | 410 | 34578 | 2515 | 266434 |
| 8 | 0.07786 | 0.0017 | 0.01204 | 0.00009 | 0.34 | 0.04766 | 0.001 | 81.8 | 49.9 | 76.1 | 1.6 | 77.2 | 0.57 | 0 | 15 | 6838 | 321 | 322 | 23773 | 1661 | 174846 |
| 9 | 0.07456 | 0.0022 | 0.01198 | 0.00012 | 0.34 | 0.04383 | 0.00125 | 0.1 | 0 | 73 | 2.08 | 76.7 | 0.75 | 18 | 10 | 6340 | 273 | 296 | 23159 | 1479 | 162987 |
| 10 | 0.082 | 0.00258 | 0.01205 | 0.00013 | 0.34 | 0.04928 | 0.0015 | 161 | 69.69 | 80 | 2.42 | 77.2 | 0.84 | 9 | 6 | 5790 | 281 | 333 | 24996 | 1381 | 147866 |
| 11 | 0.07627 | 0.0021 | 0.01195 | 0.0001 | 0.3 | 0.04541 | 0.00121 | 0.1 | 29.72 | 74.6 | 1.98 | 76.6 | 0.63 | 19 | 3 | 5472 | 245 | 340 | 25093 | 1293 | 140995 |
| 12 | 0.07299 | 0.01077 | 0.01198 | 0.00057 | 0.32 | 0.04686 | 0.00682 | 41.7 | 315.76 | 71.5 | 10.19 | 76.7 | 3.6 | 0 | 20 | 580 | 26 | 80 | 6594 | 147 | 14916 |
| 13 | 0.08298 | 0.00366 | 0.01209 | 0.00018 | 0.34 | 0.04825 | 0.00206 | 111.4 | 97.75 | 80.9 | 3.43 | 77.5 | 1.16 | 0 | 0 | 2593 | 123 | 62 | 5205 | 598 | 66035 |
| 14 | 0.07909 | 0.00191 | 0.0119 | 0.0001 | 0.35 | 0.04831 | 0.00112 | 114.7 | 54.01 | 77.3 | 1.8 | 76.3 | 0.63 | 0 | 8 | 5443 | 259 | 258 | 21769 | 1320 | 140839 |
| 15 | 0.07854 | 0.00241 | 0.01177 | 0.00012 | 0.33 | 0.04744 | 0.00141 | 70.6 | 69.64 | 76.8 | 2.26 | 75.4 | 0.78 | 5 | 1 | 5609 | 262 | 265 | 20678 | 1344 | 146734 |
| 16 | 0.07708 | 0.00234 | 0.01191 | 0.00012 | 0.33 | 0.04674 | 0.00137 | 35.6 | 69.02 | 75.4 | 2.21 | 76.3 | 0.76 | 5 | 1 | 4816 | 222 | 191 | 14864 | 1159 | 124492 |
| 17 | 0.07665 | 0.00167 | 0.01215 | 0.00009 | 0.34 | 0.04598 | 0.00095 | 0.1 | 45.6 | 75 | 1.57 | 77.9 | 0.57 | 25 | 3 | 6977 | 316 | 302 | 23875 | 1661 | 176803 |
| 18 | 0.0739 | 0.00178 | 0.01213 | 0.0001 | 0.34 | 0.04496 | 0.00104 | 0.1 | 0 | 72.4 | 1.68 | 77.7 | 0.62 | 0 | 17 | 6020 | 267 | 293 | 20051 | 1453 | 152876 |
| 19 | 0.07548 | 0.0019 | 0.01208 | 0.0001 | 0.33 | 0.04594 | 0.00111 | 0.1 | 51.19 | 73.9 | 1.79 | 77.4 | 0.65 | 0 | 12 | 5154 | 233 | 383 | 27285 | 1245 | 131415 |
| tt | 0.07909 | 0.00169 | 0.01203 | 0.00009 | 0.35 | 0.04767 | 0.00097 | 82 | 48.46 | 77.3 | 1.59 | 77.1 | 0.56 | 0 | 0 | 7301 | 343 | 314 | 23695 | 1746 | 186893 |
| Western Klaza Zone porphyry dyke - DDH KL-14-185 189-191m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 1 | 0.07721 | 0.0022 | 0.01165 | 0.00012 | 0.361498244 | 0.0489 | 0.00135 | 142.9 | 63.64 | 75.5 | 2.07 | 74.6 | 0.74 | 4 | 1 | 8362 | 398 | 613 | 49708 | 2090 | 217369 |
| 2 | 0.07886 | 0.00193 | 0.01198 | 0.0001 | 0.341069312 | 0.0477 | 0.00113 | 83.6 | 55.92 | 77.1 | 1.81 | 76.8 | 0.64 | 0 | 0 | 9871 | 459 | 559 | 42674 | 2357 | 249407 |
| 3 | 0.0804 | 0.00289 | 0.01184 | 0.00015 | 0.352450201 | 0.04953 | 0.00173 | 172.9 | 79.7 | 78.5 | 2.72 | 75.9 | 0.95 | 0 | 13 | 3126 | 151 | 571 | 47198 | 760 | 79977 |


| Fraction | Isotopic Ratios |  |  |  |  |  |  | Isotopic Ages |  |  |  |  |  | Background corrected mean counts per second at specified mass |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | ${ }^{20 \mathrm{pbj} / 3 \mathrm{su}}$ | \% $1 \sigma$ | ${ }^{20 . p p} /{ }^{\text {ma }}$ | \% 1\% | rho | ${ }^{20 . p b} /$ mapp | \% 1\% | ${ }^{20 \mathrm{~Pb} / 2{ }^{3} \mathrm{U}}$ | $1{ }^{\circ}$ | ${ }^{20 \mathrm{mp}} /{ }^{2 \times \mathrm{U}}$ | $1{ }^{1}$ | ${ }^{20 \mathrm{Pbp} / \mathrm{mapp}}$ | $1{ }^{1}$ | 202 | 204 | 206 | 207 | 208 | 232 | 235 | 238 |
| 4 | ${ }^{0.07703}$ | ${ }^{0.0017}$ | ${ }^{0.01204}$ | ${ }^{0.00009}$ | 0.338709205 | ${ }^{0.04757}$ | ${ }^{0.00101}$ | 77.1 | 50.65 | 75.4 | 1.61 | 77.2 | 0.58 | 14 | 0 | 6750 | 313 | 283 | 24680 | 1646 | 16974 |
| 5 | 0.08021 | 0.00408 | 0.01212 | 0.00021 | 0.340631673 | 0.04918 | 0.00244 | 156.3 | 112.36 | 78.3 | 3.84 | 77.7 | 1.36 | 39 | 7 | 1249 | 59 | 186 | 14525 | 302 | 31219 |
| 6 | 0.08055 | 0.00223 | 0.01189 | 0.00011 | 0.334173119 | ${ }_{0}^{0.04931}$ | 0.00132 | 162.6 | 61.39 | 78.7 | 2.09 | 76.2 | 0.72 | 0 | ${ }^{10}$ | 5212 | 250 | 208 | 16508 | 1260 | 132781 |
| 7 | 0.07726 | 0.00154 | ${ }^{0.01209}$ | 0.00008 | 0.331969106 | 0.04702 | 0.0009 | 49.8 | 44.2 | 75.6 | 1.45 | 77.5 | 0.52 | 8 | 5 | 8588 | 393 | 350 | 27676 | 2064 | 215233 |
| 8 | 0.08345 | 0.00177 | 0.01209 | ${ }^{0.00009}$ | 0.350969424 | 0.05051 | ${ }_{0} 0.00103$ | 218.5 | 46.5 | 81.4 | 1.66 | 77.5 | 0.57 | 7 | 1 | 7115 | 350 | 301 | 22047 | 1700 | 178400 |
| 9 | 0.07668 | 0.00153 | ${ }_{0} 0.01203$ | ${ }^{0.00008}$ | 0.333284436 | ${ }^{0.04676}$ | 0.0009 | 36.8 | 44.6 | 75 | 1.44 | 77.1 | 0.52 | 0 | 0 | 8677 | 395 | 394 | 28650 | 2090 | 218657 |
| 10 | 0.0729 | 0.00203 | ${ }_{0} 0.01205$ | 0.00011 | ${ }_{0}^{0.32782127}$ | ${ }^{0.04424}$ | ${ }^{0.0012}$ | 0.1 | 0 | 71.4 | 1.92 | 77.2 | 0.71 | 13 | 1 | 6910 | 298 | 234 | 18684 | 1656 | 173922 |
| ${ }^{11}$ | ${ }^{0.07594}$ | 0.00163 | ${ }^{0.01207}$ | 0.00009 | 0.347390732 | ${ }_{0}^{0.04568}$ | ${ }^{0.00094}$ | 0.1 | 29.92 | 74.3 | 1.54 | 77.3 | 0.55 | 16 | 0 | 10163 | 452 | 662 | 52701 | 2415 | 255440 |
| 12 | 0.07951 | 0.00222 | ${ }_{0} 0.01172$ | 0.00011 | 0.33615057 | 0.0485 | ${ }^{0.00131}$ | 12.7 | 62.48 | 77.7 | 2.09 | 75.1 | 0.72 | 7 | 3 | 6853 | 324 | 707 | 49634 | 1651 | 177445 |
| 13 | 0.08379 | 0.00485 | 0.01194 | 0.00024 | 0.347262084 | 0.04941 | ${ }^{0.00278}$ | 167.4 | 126.55 | 81.7 | 4.54 | 76.5 | 1.52 | 0 | 0 | 2799 | 134 | 79 | 5193 | 652 | 71104 |
| 14 | 2.5523 | 0.13393 | ${ }^{0.19079}$ | ${ }_{0} 0.0025$ | 0.249711333 | ${ }^{0.09875}$ | ${ }^{0.00218}$ | 1600.6 | 40.58 | 1287.1 | 38.28 | 1125.7 | 13.53 | 0 | ${ }^{12}$ | 9579 | 922 | 2135 | 9061 | 146 | 15236 |
| 15 | 0.0809 | 0.0018 | 0.01187 | 0.00009 | 0.340775063 | ${ }_{0}^{0.04947}$ | ${ }^{0.00106}$ | 16.9 | 49.22 | 79 | 1.69 | 76.1 | 0.58 | 0 | 8 | 6820 | 328 | 310 | 23322 | 1647 | 174347 |
| 16 | 0.08017 | 0.00215 | 0.0121 | ${ }^{0.00011}$ | 0.338985201 | ${ }^{0.04833}$ | ${ }^{0.00125}$ | 115.5 | 59.94 | 78.3 | 2.02 | 77.5 | 0.7 | ${ }^{30}$ | 0 | 5068 | 238 | ${ }^{232}$ | 15379 | 1207 | 127180 |
| 17 | ${ }_{0} 0.07627$ | 0.00171 | ${ }_{0} 0.0118$ | 0.00009 | 0.340187333 | 0.04659 | 0.00101 | 28.1 | 50.32 | 74.6 | 1.61 | 75.6 | 0.57 | 46 | 0 | 7794 | 354 | 537 | 40609 | 1881 | 200631 |
| 18 | 0.07654 | 0.0022 | ${ }_{0} 0.01206$ | 0.00012 | 0.3461782 | ${ }_{0}^{0.04507}$ | 0.00125 | 0.1 | 14.02 | 74.9 | 2.07 | 77.3 | 0.75 | 0 | 0 | 4391 | 192 | 239 | 20370 | 1021 | 110565 |
| 19 | ${ }^{0.08083}$ | 0.00215 | ${ }^{0.01208}$ | ${ }^{0.00011}$ | 0.342341753 | ${ }^{0.04892}$ | ${ }^{0.00126}$ | 14.9 | 59.31 | 78.9 | 2.02 | 77.4 | 0.69 | 2 | 8 | 4718 | 225 | 478 | 37962 | 1128 | 118664 |
| 20 | 0.0783 | 0.00182 | 0.01185 | 0.0001 | 0.363054667 | 0.04909 | 0.0011 | 15.1 | 51.67 | 76.5 | 1.71 | 75.9 | 0.61 | 5 | ${ }^{24}$ | 8008 | 383 | 509 | 36882 | 1984 | 205312 |
| Western BRX Zone porphyry dyke - DDH KL-12-154 52.1-52.5m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 1 | ${ }^{0.07857}$ | 0.00308 | 0.01194 | ${ }^{0.00016}$ | 0.34 | ${ }_{0}^{0.04797}$ | ${ }^{0.00183}$ | 96.5 | 88.86 | 76.8 | 2.9 | 76.5 | 1.02 | 23 | 0 | 3149 | 149 | 192 | 15596 | 763 | 81217 |
| 2 | 0.12332 | 0.06648 | ${ }_{0} 0.01643$ | ${ }_{0}^{0.0003}$ | 0.35 | ${ }_{0}^{0.55707}$ | 0.00285 | 49.5 | 107.18 | 118.1 | 5.85 | 105 | 1.89 | 37 | 0 | 1241 | 69 | 110 | 5766 | 228 | 23282 |
| 3 | ${ }_{0} 0.10066$ | 0.00516 | ${ }^{0.01604}$ | ${ }^{0.00028}$ | 0.34 | ${ }_{0}^{0.44701}$ | ${ }^{0.00229}$ | 49.2 | ${ }^{12.8}$ | 97.4 | 4.76 | 102.6 | 1.76 | 22 | 0 | 4055 | 188 | 260 | 14533 | 751 | 77855 |
| 4 | ${ }^{0.08324}$ | 0.00309 | ${ }^{0.01209}$ | 0.00015 | ${ }_{0} .33$ | ${ }^{0.04793}$ | ${ }^{0.00172}$ | 94.6 | 83.76 | 81.2 | 2.9 | 77.5 | 0.99 | 45 | 0 | 5123 | 242 | 378 | 27894 | 1171 | 130489 |
| 5 | ${ }^{0.08827}$ | 0.00424 | 0.01198 | ${ }^{0.0002}$ | 0.35 | ${ }_{0}^{0.55256}$ | 0.00245 | 309.7 | 102.69 | 85.9 | 3.95 | 76.8 | 1.24 | ${ }^{13}$ | 0 | 1334 | 69 | 129 | 11351 | 315 | ${ }^{34303}$ |
| 6 | ${ }^{0.07564}$ | 0.00318 | 0.01199 | 0.00016 | 0.32 | ${ }_{0} 0.04462$ | 0.00182 | 0.1 | 20.42 | 74 | 3 | 76.9 | 1.02 | 0 | 0 | 2102 | 92 | 330 | 24144 | 492 | 53970 |
| 7 | 0.07502 | 0.00218 | 0.01207 | 0.00012 | 0.34 | ${ }_{0}^{0.04653}$ | 0.00131 | 25 | 65.09 | 73.5 | 2.06 | 77.3 | 0.75 | 17 | 11 | 4022 | 184 | 548 | 41025 | 990 | 102669 |
| 8 | ${ }_{0} 0.08131$ | 0.00165 | ${ }_{0} 0.01212$ | ${ }^{0.00008}$ | 0.33 | ${ }_{0}^{0.04803}$ | ${ }^{0.00092}$ | 100.5 | 44.61 | 79.4 | 1.55 | 77.7 | 0.53 | 0 | ${ }^{12}$ | 8726 | 413 | 371 | 27849 | 2045 | 221743 |
| 9 | ${ }_{0} 0.08134$ | 0.00209 | 0.01189 | 0.0001 | ${ }_{0} 033$ | ${ }_{0}^{0.04863}$ | ${ }^{0.0012}$ | 129.9 | 56.89 | 79.4 | 1.96 | 76.2 | 0.66 | 15 | 0 | 5058 | 242 | 506 | 37076 | 1200 | 131005 |
| 10 | 0.08249 | 0.00289 | ${ }^{0.01184}$ | 0.00014 | 0.34 | ${ }_{0}^{0.04933}$ | ${ }^{0.00166}$ | 163.8 | 77 | 80.5 | 2.71 | 75.9 | 0.92 | ${ }^{56}$ | 0 | 4163 | 202 | ${ }_{533}$ | 42260 | ${ }^{988}$ | 108256 |
| ${ }^{11}$ | ${ }^{0.07457}$ | 0.00284 | 0.01199 | ${ }^{0.00016}$ | 0.35 | ${ }^{0.04548}$ | ${ }^{0.00168}$ | ${ }^{0.1}$ | 56.52 | 73 | 2.68 | 76.8 | 0.99 | 35 | 3 | 3911 | 175 | 570 | 45074 | 946 | 100459 |
| 12 | ${ }_{0} 0.88073$ | 0.00267 | ${ }^{0.01206}$ | 0.00014 | 0.35 | ${ }_{0}^{0.44707}$ | 0.0015 | 52.5 | 74.57 | 78.8 | 2.51 | 77.3 | 0.87 | 2 | 0 | 2915 | 135 | 349 | 28126 | 674 | 74440 |
| 13 | ${ }_{0}^{0.0762}$ | 0.00309 | ${ }_{0} 0.01182$ | ${ }^{0.00016}$ | 0.33 | ${ }_{0}^{0.04648}$ | ${ }^{0.00183}$ | 22.6 | 91.94 | 74.6 | 2.92 | 75.8 | 1.03 | 5 | ${ }^{11}$ | 4247 | 195 | 447 | 35211 | 1028 | 110631 |
| 14 | 0.07933 | 0.00307 | ${ }^{0.01196}$ | 0.00016 | 0.35 | ${ }^{0.04814}$ | ${ }^{0.0018}$ | 106.2 | 86.2 | 77.5 | 2.89 | 76.6 | 1.02 | 0 | 0 | 4290 | 204 | 124 | 10926 | 1033 | 110482 |
| 15 | 0.0717 | 0.00225 | 0.01199 | 0.00012 | 0.32 | 0.04417 | ${ }^{0.00134}$ | 0.1 | 0 | 70.3 | 2.13 | 76.8 | 0.78 | 42 | 2 | 3632 | 158 | 499 | 42268 | ${ }^{888}$ | 93326 |
| 16 | 0.07925 | 0.00292 | ${ }_{0} 0.01209$ | 0.00015 | 0.34 | ${ }_{0}^{0.48803}$ | ${ }^{0.00171}$ | 100.9 | 82.15 | 77.4 | 2.75 | 77.5 | 0.97 | 54 | 0 | 3064 | 145 | 249 | 20693 | 737 | 78067 |
| 17 | ${ }^{0.07647}$ | 0.00262 | ${ }_{0} 0.01187$ | 0.00013 | 0.32 | ${ }_{0}^{0.44828}$ | ${ }_{0} 0.0016$ | 113.2 | 76.55 | 74.8 | 2.47 | 76.1 | 0.86 | ${ }^{13}$ | 0 | 2868 | 136 | 312 | 23742 | 718 | 74399 |
| 18 | 0.07375 | 0.00213 | ${ }_{0} 0.01207$ | ${ }^{0.00011}$ | 0.32 | ${ }_{0}^{0.04515}$ | 0.00125 | 0.1 | 18.28 | ${ }^{2} 2.3$ | 2.01 | 77.4 | 0.73 | 4 | 5 | 4228 | 188 | 230 | 19216 | 1027 | 107871 |
| 19 | $0^{0.07802}$ | 0.00175 | ${ }_{0} 0.01216$ | 0.00009 | 0.33 | ${ }_{0}^{0.04633}$ | 0.00098 | 14.7 | 49.77 | ${ }^{76.3}$ | 1.65 | 77.9 | 0.59 | 13 | 3 | 7029 | 321 | 364 | 30861 | 1656 | 178127 |
| 20 | 0.07852 | 0.00166 | 0.01215 | 0.00009 | 0.35 | ${ }^{0.04747}$ | 0.00094 | 72.5 | 47.14 | 76.7 | 1.56 | 77.9 | 0.56 | 14 | 4 | 7944 | 372 | 521 | 40287 | 1906 | 201352 |
| Central BRX Zone porphyry dyke - DDH KL-14-165 85.7-87.5m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 1 | 0.4214 | 0.02236 | ${ }^{0.05902}$ | ${ }^{0.00094}$ | 0.300159007 | ${ }^{0.05357}$ | 0.00215 | 35.8 | 88.13 | 357.1 | 15.97 | 369.7 | 5.73 | 0 | 0 | 4977 | 259 | 464 | 7052 | 250 | 25637 |
| 2 | ${ }^{0.07612}$ | 0.00177 | ${ }^{0.01198}$ | ${ }^{0.00009}$ | 0.323088841 | ${ }_{0}^{0.04639}$ | 0.00104 | 17.9 | 52.3 | 74.5 | 1.67 | 76.8 | 0.6 | ${ }^{20}$ | ${ }^{6}$ | 7759 | 350 | 588 | 44796 | 1869 | 196970 |
| 3 | ${ }^{0.08222}$ | 0.00191 | ${ }^{0.01198}$ | 0.0001 | 0.359324878 | 0.04959 | 0.00111 | 176 | 51.21 | 80.2 | 1.79 | 76.8 | 0.61 | 27 | 0 | 5991 | 289 | 556 | 42236 | 1428 | 152096 |
| 4 | 0.07761 | 0.00247 | ${ }^{0.01213}$ | 0.00013 | 0.336746648 | ${ }^{0.04787}$ | 0.00148 | 91.7 | 72.61 | 75.9 | 2.33 | 77.7 | 0.83 | 0 | 0 | 3283 | 153 | 450 | 33599 | 800 | ${ }_{82357}$ |
| 5 | 0.07443 | 0.00265 | ${ }_{0} 0.01183$ | ${ }^{0.00013}$ | 0.308646071 | ${ }^{0.0453}$ | ${ }^{0.00157}$ | 0.1 | 42.22 | 72.9 | 2.5 | 75.8 | 0.85 | 7 | 1 | 2898 | 128 | 169 | 15214 | 697 | 74560 |
| 6 | 0.08214 | 0.00448 | 0.01189 | ${ }^{0.00022}$ | 0.339248168 | 0.04969 | ${ }^{0.00264}$ | 180.3 | 119.42 | 80.2 | 4.21 | 76.2 | 1.42 | 10 | 11 | 1770 | 85 | 119 | 8716 | 423 | 45320 |
| 7 | 0.08011 | 0.00336 | 0.01215 | 0.00017 | 0.333595434 | 0.04669 | 0.0019 | 33.3 | 94.74 | 78.3 | 3.16 | 77.9 | 1.09 | 31 | 0 | 3059 | 139 | 102 | 9127 | 705 | 76636 |
| 8 | 0.103 | 0.00434 | ${ }^{0.01652}$ | ${ }^{0.00023}$ | ${ }^{0.33041921}$ | ${ }^{0.04609}$ | 0.00184 | 2.5 | 92.31 | 99.5 | 3.99 | 105.7 | 1.45 | 35 | 15 | 2680 | 120 | 399 | 20584 | 474 | 49377 |
| 9 | 0.07675 | ${ }_{0} 0.029$ | ${ }_{0} 0.01166$ | 0.00015 | 0.340465488 | 0.04738 | 0.00174 | 67.9 | 85.72 | 75.1 | 2.73 | 74.8 | 0.96 | 32 | 0 | 2691 | 124 | 188 | 13815 | 657 | 70253 |


| Fraction | Isotopic Ratios |  |  |  |  |  |  | Isotopic Ages |  |  |  |  |  | Background corrected mean counts per second at specified mass |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | ${ }^{207} \mathbf{P b} /{ }^{25} \mathrm{U}$ | \% 1 $\sigma$ | ${ }^{206} \mathrm{~Pb} /{ }^{2 \times \mathrm{U}} \mathrm{U}$ | \% 1 $\sigma$ | rho | ${ }^{207 P b / 205 P b}$ | \% 16 | ${ }^{207} \mathrm{~Pb} /{ }^{235} \mathrm{U}$ | 1\% | ${ }^{200} \mathrm{~Pb} /{ }^{23 \mathrm{sax}}$ | $1{ }^{\circ}$ | ${ }^{207 \mathrm{~Pb}} /{ }^{205} \mathrm{~Pb}$ | $1{ }^{\circ}$ | 202 | 204 | 206 | 207 | 208 | 232 | 235 | 238 |
| 10 | 0.0763 | 0.00204 | 0.01192 | 0.00011 | 0.345152323 | 0.04743 | 0.00122 | 70.1 | 60.85 | 74.7 | 1.92 | 76.4 | 0.68 | 21 | 0 | 4743 | 219 | 231 | 17288 | 1165 | 121141 |
| 11 | 0.07701 | 0.00265 | 0.01187 | 0.00014 | 0.34275087 | 0.04914 | 0.00164 | 154.5 | 76.43 | 75.3 | 2.5 | 76.1 | 0.91 | 0 | 0 | 3779 | 181 | 202 | 17852 | 953 | 97018 |
| 12 | 0.07456 | 0.00244 | 0.01183 | 0.00013 | 0.335795352 | 0.04725 | 0.0015 | 61.3 | 74.49 | 73 | 2.3 | 75.8 | 0.85 | 12 | 9 | 3181 | 146 | 376 | 29971 | 797 | 81920 |
| 13 | 0.07798 | 0.00247 | 0.01192 | 0.00013 | 0.344312964 | 0.04789 | 0.00147 | 92.9 | 72.21 | 76.2 | 2.33 | 76.4 | 0.83 | 6 | 0 | 3070 | 143 | 286 | 20322 | 745 | 78502 |
| 14 | 0.07638 | 0.00287 | 0.01189 | 0.00015 | 0.335743151 | 0.04759 | 0.00174 | 78.2 | 85.39 | 74.7 | 2.71 | 76.2 | 0.96 | 39 | 1 | 3602 | 167 | 282 | 22474 | 887 | 92353 |
| 15 | 0.10795 | 0.0055 | 0.01654 | 0.00029 | 0.344129933 | 0.05074 | 0.00245 | 229.1 | 107.96 | 104.1 | 5.04 | 105.7 | 1.84 | 39 | 12 | 2382 | 117 | 277 | 16336 | 442 | 43923 |
| 16 | 0.10907 | 0.00459 | 0.01642 | 0.00023 | 0.332849042 | 0.04976 | 0.00198 | 184 | 90.28 | 105.1 | 4.2 | 105 | 1.48 | 0 | 0 | 2615 | 126 | 304 | 16974 | 472 | 48563 |
| 17 | 0.11018 | 0.00446 | 0.01652 | 0.00022 | 0.328988371 | 0.04905 | 0.00188 | 150.4 | 87.4 | 106.1 | 4.08 | 105.6 | 1.4 | 7 | 13 | 2596 | 124 | 321 | 18549 | 457 | 47945 |
| 18 | 0.06905 | 0.00269 | 0.01201 | 0.00015 | 0.320597148 | 0.04333 | 0.00165 | 0.1 | 0 | 67.8 | 2.56 | 77 | 0.98 | 8 | 2 | 5098 | 215 | 853 | 65617 | 1265 | 129486 |
| 19 | 0.07623 | 0.00275 | 0.01193 | 0.00015 | 0.34853311 | 0.04807 | 0.00168 | 102.7 | 80.72 | 74.6 | 2.59 | 76.5 | 0.93 | 0 | 6 | 3122 | 146 | 345 | 26623 | 778 | 79838 |
| 20 | 0.08083 | 0.00276 | 0.01199 | 0.00014 | 0.34195767 | 0.04837 | 0.00159 | 117.3 | 75.89 | 78.9 | 2.59 | 76.8 | 0.89 | 0 | 4 | 3892 | 183 | 220 | 15896 | 921 | 99036 |
| Central Klaza Zone porphyry dyke - DDH KL-11-12 181-187m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| , | 0.07448 | 0.00364 | 0.01084 | 0.00018 | 0.339767244 | 0.04901 | 0.00234 | 148.3 | 108.21 | 72.9 | 3.44 | 69.5 | 1.16 | 5 | 20 | 2757 | 131 | 225 | 19019 | 718 | 77966 |
| 2 | 0.07481 | 0.0024 | 0.01105 | 0.00012 | 0.338506787 | 0.04945 | 0.00153 | 169.4 | 70.65 | 73.3 | 2.27 | 70.8 | 0.79 | 18 | 8 | 7257 | 349 | 600 | 53997 | 1900 | 201387 |
| 3 | 0.073 | 0.00244 | 0.01118 | 0.00013 | 0.347884102 | 0.04887 | 0.00158 | 141.8 | 74.04 | 71.5 | 2.31 | 71.7 | 0.83 | 2 | 11 | 8464 | 403 | 641 | 56851 | 2244 | 232176 |
| 4 | 0.0792 | 0.00515 | 0.01204 | 0.00027 | 0.344869851 | 0.04701 | 0.00297 | 49.5 | 144.68 | 77.4 | 4.85 | 77.1 | 1.69 | 0 | 6 | 2273 | 104 | 539 | 37604 | 534 | 57918 |
| 5 | 0.06699 | 0.0038 | 0.01122 | 0.00019 | 0.298529412 | 0.04307 | 0.0024 | 0.1 | 0 | 65.8 | 3.62 | 71.9 | 1.21 | 34 | 0 | 2368 | 99 | 341 | 34958 | 603 | 64751 |
| 6 | 0.07674 | 0.00495 | 0.01106 | 0.00024 | 0.336412954 | 0.04963 | 0.00313 | 177.8 | 140.63 | 75.1 | 4.67 | 70.9 | 1.52 | 29 | 5 | 1227 | 59 | 249 | 22588 | 314 | 34054 |
| 7 | 0.07187 | 0.00216 | 0.01121 | 0.00011 | 0.326498332 | 0.04655 | 0.00134 | 26 | 66.76 | 70.5 | 2.05 | 71.9 | 0.73 | 0 | 0 | 9272 | 420 | 752 | 58934 | 2379 | 253765 |
| 8 | 0.07572 | 0.00262 | 0.01121 | 0.00014 | 0.360937277 | 0.04947 | 0.00166 | 170.4 | 76.31 | 74.1 | 2.48 | 71.9 | 0.86 | 32 | 12 | 4320 | 208 | 289 | 22619 | 1118 | 118276 |
| 9 | 0.07565 | 0.00187 | 0.0111 | 0.00009 | 0.328009828 | 0.04915 | 0.00115 | 154.9 | 53.99 | 74 | 1.77 | 71.2 | 0.6 | 8 | 0 | 7332 | 351 | 1037 | 88488 | 1887 | 202730 |
| 10 | 0.07119 | 0.00202 | 0.01106 | 0.0001 | 0.318648954 | 0.04669 | 0.00127 | 33.4 | 62.68 | 69.8 | 1.91 | 70.9 | 0.67 | 0 | 0 | 5329 | 242 | 410 | 34305 | 1385 | 147894 |
| 11 | 0.0744 | 0.00413 | 0.01111 | 0.00022 | 0.356723324 | 0.05107 | 0.00278 | 243.8 | 120.58 | 72.9 | 3.91 | 71.2 | 1.38 | 0 | 2 | 2417 | 120 | 199 | 15090 | 657 | 66888 |
| 12 | 0.07431 | 0.00265 | 0.01095 | 0.00013 | 0.332912897 | 0.04698 | 0.00161 | 47.7 | 80.35 | 72.8 | 2.51 | 70.2 | 0.85 | 0 | 0 | 9223 | 422 | 920 | 80042 | 2310 | 258871 |
| 13 | 0.0736 | 0.00141 | 0.01114 | 0.00007 | 0.327998268 | 0.04705 | 0.00083 | 51.5 | 40.78 | 72.1 | 1.34 | 71.4 | 0.45 | 3 | 17 | 12350 | 566 | 1332 | 113347 | 3129 | 340633 |
| 14 | 0.07094 | 0.00151 | 0.01114 | 0.00008 | 0.337379766 | 0.04591 | 0.00091 | 0.1 | 39.81 | 69.6 | 1.43 | 71.4 | 0.5 | 11 | 5 | 9350 | 418 | 751 | 70332 | 2398 | 257988 |
| 15 | 0.07458 | 0.00142 | 0.01116 | 0.00007 | 0.329433591 | 0.04805 | 0.00084 | 101.7 | 40.73 | 73 | 1.34 | 71.5 | 0.46 | 0 | 18 | 12284 | 574 | 1185 | 97514 | 3137 | 338457 |
| 16 | 0.07397 | 0.0022 | 0.01135 | 0.00011 | 0.325859031 | 0.04655 | 0.00132 | 25.9 | 65.81 | 72.5 | 2.08 | 72.8 | 0.71 | 0 | 16 | 6118 | 277 | 1038 | 85866 | 1526 | 165778 |
| 17 | 0.07529 | 0.00177 | 0.01127 | 0.00009 | 0.339689892 | 0.0496 | 0.00109 | 176.5 | 50.52 | 73.7 | 1.67 | 72.2 | 0.57 | 9 | 0 | 7123 | 344 | 563 | 48264 | 1860 | 194460 |
| 18 | 0.07442 | 0.00207 | 0.01122 | 0.00011 | 0.352467557 | 0.04901 | 0.0013 | 148.1 | 60.98 | 72.9 | 1.96 | 71.9 | 0.68 | 3 | 7 | 7404 | 353 | 501 | 46832 | 1933 | 203052 |
| 19 | 0.07411 | 0.00165 | 0.01121 | 0.00008 | 0.320536318 | 0.04865 | 0.00101 | 130.9 | 48 | 72.6 | 1.56 | 71.9 | 0.54 | 10 | 0 | 7872 | 373 | 941 | 76164 | 2048 | 216055 |
| tt | 0.07248 | 0.00184 | 0.01109 | 0.00009 | 0.319676951 | 0.04785 | 0.00114 | 90.7 | 56.67 | 71.1 | 1.74 | 71.1 | 0.61 | 0 | 6 | 7918 | 368 | 728 | 60397 | 2072 | 219614 |
| Eastern BRX Zone porphyry dyke - DDH KL-12-134 218-221.5m |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 1 | 0.07253 | 0.00674 | 0.01231 | 0.00034 | 0.3 | 0.04388 | 0.00401 | 0.1 | 92.74 | 71.1 | 6.38 | 78.9 | 2.15 | 0 | 0 | 744 | 32 | 132 | 10868 | 178 | 18607 |
| 2 | 0.08244 | 0.00745 | 0.01201 | 0.00032 | 0.29 | 0.04826 | 0.00428 | 111.9 | 196.98 | 80.4 | 6.99 | 77 | 2.02 | 22 | 0 | 716 | 33 | 126 | 9651 | 166 | 18353 |
| 3 | 0.11223 | 0.00478 | 0.01718 | 0.00024 | 0.33 | 0.04928 | 0.00199 | 161.4 | 91.69 | 108 | 4.36 | 109.8 | 1.55 | 0 | 26 | 2445 | 118 | 347 | 19648 | 426 | 43790 |
| 4 | 0.0689 | 0.00647 | 0.0121 | 0.00033 | 0.29 | 0.04549 | 0.00424 | 0.1 | 182.37 | 67.7 | 6.15 | 77.5 | 2.08 | 21 | 0 | 530 | 23 | 57 | 5774 | 139 | 13490 |
| 5 | 0.08195 | 0.00849 | 0.01225 | 0.0004 | 0.32 | ${ }^{0.0503}$ | 0.00512 | 208.7 | 219.94 | 80 | 7.97 | 78.5 | 2.56 | 35 | 13 | 648 | 32 | 93 | 8795 | 157 | 16277 |
| 6 | 0.07507 | 0.00415 | 0.0122 | 0.00023 | 0.34 | 0.0457 | 0.00247 | 0.1 | 107.82 | 73.5 | 3.92 | 78.2 | 1.44 | 24 | 17 | 1405 | 63 | 236 | 16881 | 339 | 35437 |
| 7 | 0.08152 | 0.00661 | 0.01211 | 0.00032 | 0.33 | 0.05007 | 0.00399 | 198.5 | 174.99 | 79.6 | 6.21 | 77.6 | 2.07 | 16 | 0 | 806 | 39 | 94 | 8161 | 196 | 20473 |
| 8 | 0.11126 | 0.00587 | 0.01695 | 0.00029 | 0.32 | 0.04729 | 0.00237 | 63.3 | 115.64 | 107.1 | 5.36 | 108.3 | 1.85 | 0 | 4 | 1678 | 77 | 253 | 14575 | 283 | 30463 |
| 9 | 0.10709 | 0.00515 | 0.0168 | 0.00026 | 0.32 | 0.04667 | 0.00214 | 32.3 | 106.29 | 103.3 | 4.72 | 107.4 | 1.62 | 10 | 14 | 2180 | 99 | 262 | 15335 | 377 | 39906 |
| 10 | 0.06928 | 0.00672 | 0.01232 | 0.00038 | 0.32 | 0.04149 | 0.00394 | 0.1 | 0 | 68 | 6.38 | 78.9 | 2.39 | 19 | 1 | 1136 | 46 | 139 | 11505 | 270 | 28387 |
| 11 | 0.07806 | 0.00325 | 0.01231 | 0.00017 | 0.33 | 0.04595 | 0.00186 | 0.1 | 88.68 | 76.3 | 3.06 | 78.9 | 1.09 | 0 | 9 | 2373 | 107 | 131 | 9687 | 554 | 59314 |
| 12 | 0.07277 | 0.00798 | 0.01289 | 0.00045 | 0.32 | 0.04469 | 0.00481 | 0.1 | 171.16 | 71.3 | 7.55 | 82.6 | 2.88 | 21 | 15 | 753 | 33 | 141 | 10173 | 183 | 17971 |
| 13 | 0.07544 | 0.00993 | 0.01242 | 0.00051 | 0.31 | 0.04592 | 0.00594 | 0.1 | 278.95 | 73.8 | 9.38 | 79.6 | 3.24 | 0 | 3 | 737 | 33 | 161 | 10879 | 178 | 18258 |
| 14 | 0.08267 | 0.00654 | 0.01225 | 0.0003 | 0.31 | 0.04839 | 0.00375 | 118.4 | 173 | 80.7 | 6.13 | 78.5 | 1.89 | 45 | 0 | 792 | 37 | 121 | 10968 | 184 | 19914 |
| 15 | 0.10015 | 0.01412 | 0.01234 | 0.00061 | 0.35 | 0.05407 | 0.00737 | 373.8 | 280.9 | 96.9 | 13.03 | 79 | 3.91 | 0 | 0 | 657 | 34 | 105 | 8427 | 141 | 16402 |




Appendix 2. Pb isotopic analyses of igneous feldspar from the Cyprus porphyry and galenas from the Klaza and BrownMcDade clusters. Errors are given at the $1 \sigma$ level.

| Sample | ${ }^{206} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ | error | ${ }^{207} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ | error | ${ }^{208} \mathrm{~Pb} /{ }^{204} \mathrm{~Pb}$ | error | ${ }^{207} \mathbf{P b} /{ }^{206} \mathrm{~Pb}$ | error | ${ }^{208} \mathrm{~Pb} /{ }^{206} \mathrm{~Pb}$ | error |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Cyprus porphyry feldspar (97CH33-3) | 19.267 | 0.09 | 15.6602 | 0.04 | 39.0727 | 0.1 | 0.8128 | 0.085 | 2.028 | 0.028 |
| Brown-McDade - <br> Dickson Zone galena | 18.829 | 0.012 | 15.634 | 0.011 | 38.552 | 0.013 | 0.8303 | 0.005 | 2.0475 | 0.003 |
| Brown-McDade Flex Zone galena | 19.156 | 0.006 | 15.64 | 0.005 | 38.816 | 0.009 | 0.8165 | 0.003 | 2.0263 | 0.007 |
| Brown-McDade - <br> Heustis Zone galena | 19.186 | 0.002 | 15.672 | 0.002 | 38.92 | 0.002 | 0.8168 | 0.001 | 2.0285 | 0.001 |
| Klaza - Pika vein galena | 19.1485 | 0.01 | 15.685 | 0.01 | 38.9112 | 0.01 | 0.8191 | 0.008 | 2.0321 | 0.007 |
| Klaza-KL11-12-196 galena | 19.1859 | 0.01 | 15.7165 | 0.01 | 39.026 | 0.02 | 0.8192 | 0.006 | 2.0341 | 0.011 |
| Klaza - KL14-165 galena | 19.1518 | 0.01 | 15.68 | 0.01 | 38.9108 | 0.01 | 0.8187 | 0.005 | 2.0317 | 0.006 |

